

川西新元古代苏雄组层型剖面 SHRIMP 锆石 U-Pb 年龄及其地质意义

卓皆文^{1,2)}, 江卓斐^{1,2)}, 江新胜^{1,2)}, 王剑^{1,2)}, 蔡娟娟³⁾, 熊国庆^{1,2)},
陆俊泽^{1,2)}, 崔晓庄^{1,2)}, 刘建辉⁴⁾

1) 中国地质调查局成都地质调查中心, 成都, 610081;

2) 国土资源部沉积盆地与油气资源重点实验室, 成都, 610081;

3) 中国地质科学院研究生部, 北京, 100037; 4) 北京离子探针中心, 北京, 102206

内容提要:对采自川西新元古代苏雄组层型剖面中上部的英安岩进行 SHRIMP 锆石 U-Pb 定年, 获得了两组谐和年龄值: $838 \pm 9 \text{ Ma}$ ($n = 9, \text{MSWD} = 0.80$) 和 $780 \pm 12 \text{ Ma}$ ($n = 5, \text{MSWD} = 1.2$)。其中 $780 \pm 12 \text{ Ma}$ 解释为苏雄组英安岩的喷发时限, 代表康滇裂谷盆地强烈扩张时期的构造岩浆热事件; 而 $838 \pm 9 \text{ Ma}$ 则可能代表康滇裂谷盆地开启前的构造岩浆热事件。综合研究认为, 苏雄组位于康滇裂谷半地堑盆地的调节带, 为华南新元古代幕式岩浆活动的产物; 火山活动和沉积记录均表明, 扬子周缘发育与地幔柱相关的裂谷盆地系统, 且华南可能是 Rodinia 超大陆裂解的重要组成单元。

关键词:苏雄组; 英安岩; SHRIMP 锆石 U-Pb 年龄; 新元古代; 构造岩浆热事件

扬子周缘新元古代岩浆活动非常强烈, 其喷发时代与构造属性对于重建华南在 Rodinia 超大陆中的位置具有极为重要的意义 (Li Xianhua et al., 2003; Li Zhengxiang et al., 2003; Wang Jian and Li Zhengxiang, 2003; Cui Xiaozhuang et al., 2015)。一般认为, 苏雄组为扬子西缘新元古代康滇裂谷盆地的充填物 (图 1a), 主要分布于小相岭、大相岭、甘洛和西昌螺髻山等地, 主体为一套双峰式火山岩 (刘鸿允等, 1981, 1991; Li Xianhua et al., 2002)。然而, 限于测年技术的精度, 以往获得的苏雄组年龄变化范围大 ($900 \sim 600 \text{ Ma}$) (刘鸿允等, 1981; 董榕生与李建林, 1981; 张盛师, 1985; 马国干等, 1989; 刘鸿允等, 1991), 可靠性欠佳。Li Xianhua 等 (2002) 首次对苏雄组底部流纹岩进行了年龄测定, 获得了 $803 \pm 12 \text{ Ma}$ 的 SHRIMP 锆石 U-Pb 年龄; 谷志东等 (2014a) 对四川威远地区威 117 井花岗岩基底岩芯进行了 SHRIMP 锆石 U-Pb 测年, 获得了 $794 \pm 11 \text{ Ma}$ 的年龄数据, 罗志立 (1986) 认为花岗岩可与苏雄组对比。尽管如此, 仍需更多可靠年龄数据对苏

雄组的喷发时限进行精确限定, 进而认识华南的大地构造属性及重建其在 Rodinia 超大陆中的位置。最近, 笔者在川西甘洛地区的苏雄组层型剖面中上部采集了 1 个英安岩样品 (LGP-20N1, N $29^{\circ}07'10''$, E $102^{\circ}51'21''$), 进行了 SHRIMP 锆石 U-Pb 同位素测定, 为确定苏雄组双峰式火山活动的时限和研究扬子的幕式构造岩浆热事件提供了新的年代学资料; 并结合扬子周缘新元古代火山活动与沉积记录, 探讨了华南的大地构造属性。

1 地质背景

苏雄组分布于扬子西缘康滇新元古代裂谷盆地中段宝兴—汉源—西昌—甘洛—盐边一线 (四川省地质矿产局, 1997), 被认为是华南新元古代大火山岩省 (LIPs) 之一 (Ernst et al., 2008; Li Zhengxiang et al., 2008), 是康滇裂谷盆地调节带的核心部位 (卓皆文等, 2013), 其喷发时代为新元古代中期 (Li Xianhua et al., 2002)。最新研究表明, 苏雄组可与开建桥组、澄江组、牛头山组、陆良组上段、莲沱组进

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作者简介: 卓皆文, 男, 1980 年生。博士, 高级工程师, 主要从事沉积学与大地构造研究。电话: 028-83231381; Email: zhuojiewen_1980@163.com。

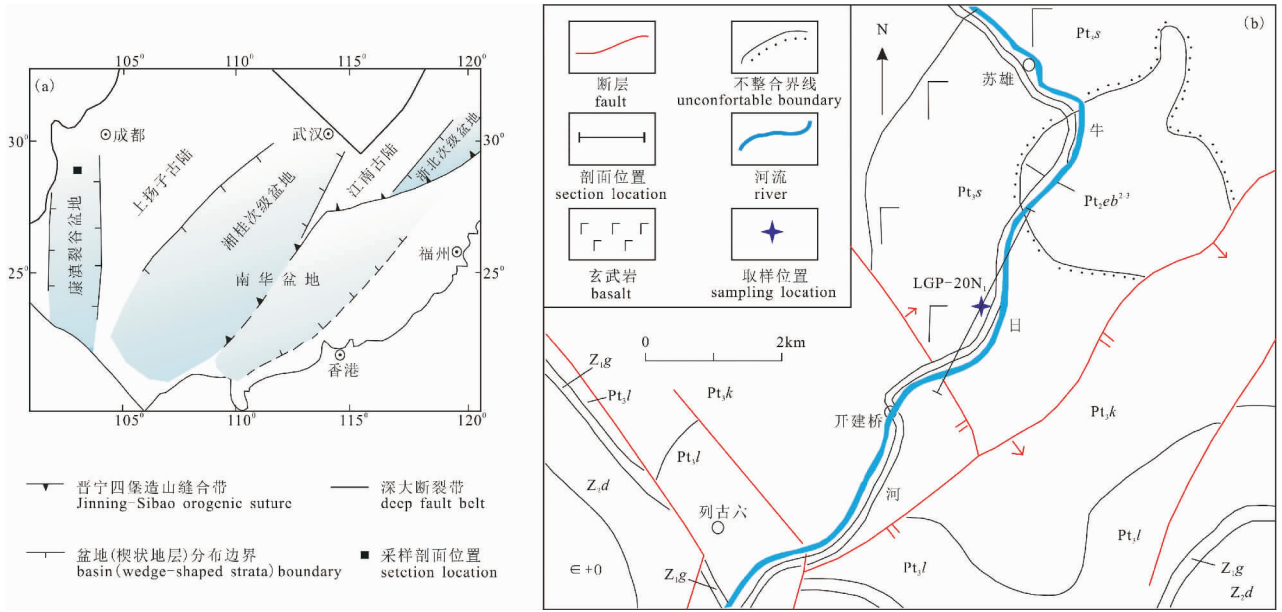


图1 华南新元古代裂谷次级盆地分布图(a, 据 Wang Jian 与 Li Zhengxiang, 2003 修改)和研究区区域地质背景图(b, 据卓皆文等, 2015 修改)

Fig. 1 Neoproterozoic rift sub-basins in South China (a, modified after Wang Jian and Li Zhengxiang, 2003) and the regional geological setting of studied area (b, modified after Zhuo Jiewen et al., 2015&)

Pt_2eb^{2-3} —中元古界峨边群第二、三段; Pt_3s —新元古界苏雄组; Pt_3k —新元古界开建桥组; Pt_3l —新元古界列古六组;

Z_1g —下震旦(埃迪卡拉)统观音崖组; Z_2d —上震旦(埃迪卡拉)统灯影组; $\mathbb{C} + 0$ —寒武系和奥陶系

Pt_2eb^{2-3} —No. 2 and 3 Member of the Ebian Group, Mesoproterozoic; Neoproterozoic; Pt_3s —Suxiong Formation, Pt_3k —Kaijianqiao Formation, Pt_3l —Lieguliu Formation; Z_1g —Lower Sinian (Ediacaran) Guanyinya Formation; Z_2d —Upper Sinian (Ediacaran) Dengying Formation; $\mathbb{C} + 0$ —Cambrian and Ordovician

行对比(江新胜等, 2012; 陆俊泽等, 2013; 崔晓庄等, 2014, 2015; 卓皆文等, 2013, 2015)。

苏雄组层型剖面位于四川西部甘洛县苏雄, 沿牛日河至开建桥; 下与中元古代峨边群呈角度不整合接触, 上与开建桥组呈断层接触(图1b, 2)。在区域上, 苏雄组厚数百米至上万米, 与开建桥组在横向上呈相互消长关系(杨暹和与陈远德, 1981; 刘鸿允等, 1991; 卓皆文等, 2015)。笔者对苏雄组层型剖面进行了实测, 测得剖面实际厚度为787 m。苏雄组可分为三段: 第1~19层为下段, 厚470 m, 以紫灰色、灰绿色厚层状火山碎屑岩为主, 夹灰绿色凝灰岩和暗紫色玄武岩、紫灰色英安岩, 底部为10 m厚的灰色块状砾岩; 第20~21层为中段, 厚282 m, 以灰白色层状流纹岩和紫灰色厚层块状英安岩等酸性熔岩为主; 第22~24层为上段, 厚35 m, 主要为灰绿色厚层块状沉凝灰岩和深灰色致密块状玄武岩(图2)。本文研究样品采自层型剖面第20层, 为紫灰色厚层块状英安岩(图2)。

2 样品及分析方法

英安岩呈紫灰色, 厚层块状(图3a)。镜下观察具斑状结构, 斑晶含量10%~15%, 主要矿物为长石和石英, 副矿物为磁铁矿和锆石; 长石主要为钾长石, 部分酸性斜长石, 多呈板条状、柱粒状; 石英以细中粒浑圆状为主, 表面干净, 个别有熔蚀边; 基质具隐晶—微粒结构, 含量85%~90%, 主要由长英质微粒和隐晶硅质组成(图3b)。

将样品粉碎至60目以下, 先用人工重砂方法富集锆石, 然后在双目镜下精选锆石。将挑好的锆石颗粒粘在双面胶上, 固定在透明的环氧树脂中, 和树脂一起打磨抛光, 直至露出锆石的内部以适合分析。对抛光后的锆石进行反射光、透射光显微照相和阴极发光(CL)图像分析, 选择代表性的锆石颗粒和区域进行U—Th—Pb同位素分析。

SHRIMP 锆石U—Pb分析在中国地质科学院地质研究所北京离子探针中心SHRIMP II上完成。在分析过程中, 一次离子流强度为4 nA, 一次离子流

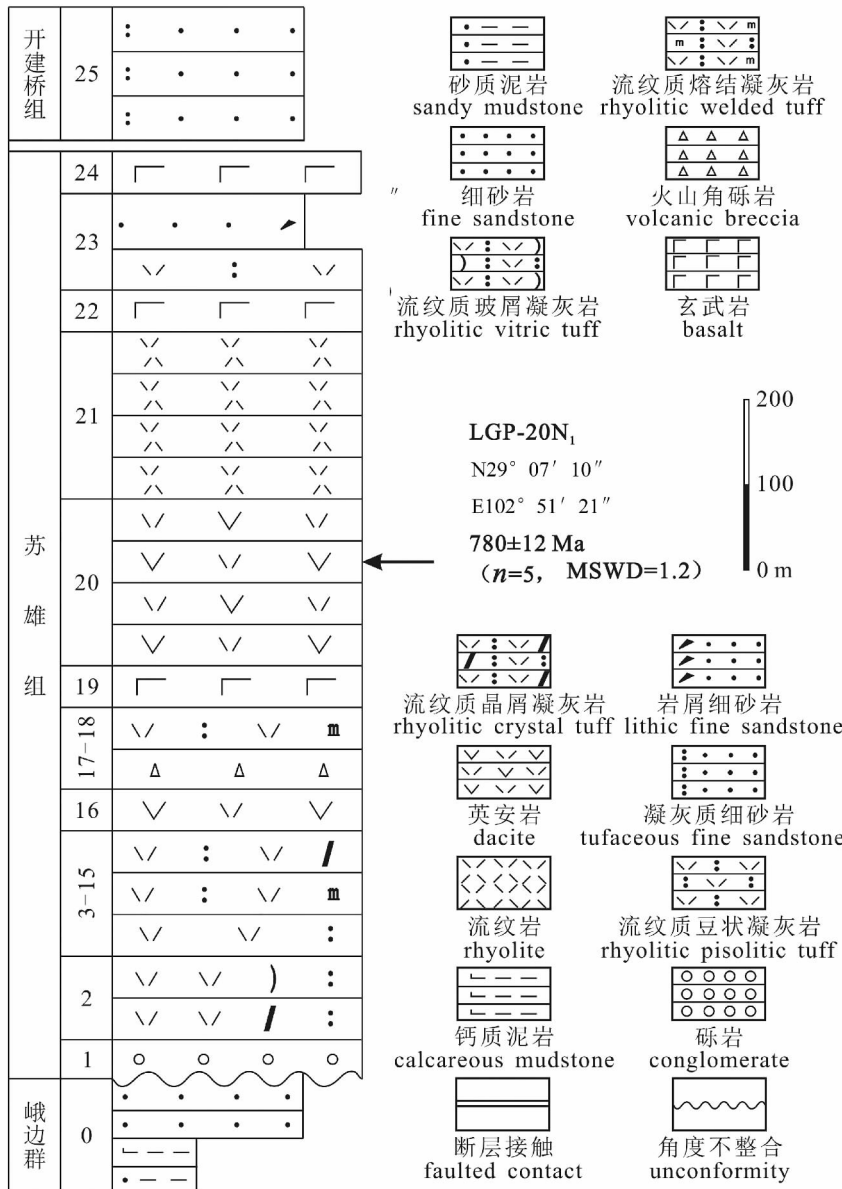


图 2 川西新元古界苏雄组层型剖面柱状图及采样位置

Fig. 2 The sample position and stratotype section of the Neoproterozoic Suxiong Formation in western Sichuan province

束斑为 20 ~ 30 μm 。每个测量点由 5 组扫描获得。标样为 M257 (U 含量为 840×10^{-6}) (Nasdala et al., 2008) 和 TEM (年龄为 417 Ma) (Black et al., 2003), 分别用于 U 含量和 $n(^{206}\text{Pb})/n(^{238}\text{U})$ 年龄值校正。为了监测仪器的稳定性和同位素分馏行为, 每隔 2 ~ 3 个未知样品测量即进行一次标样测试 (TEM), 详细的 SHRIMP 分析流程按照 Compston 等 (1984)、Williams 与 Claesson (1987) 所设计的流程进行。数据处理和年龄计算采用程序 SQUID 1.02 和 ISOPLOT 3.0 (Ludwig, 2003); 衰变常数使用

Steiger 与 Jager (1977) 的推荐值; 普通铅校正使用直接测定 ^{204}Pb 方法 (Compston et al., 1984), 其组成用 Stacey-Kramers 模式给出的相应时间地壳 Pb 同位素组成平均值 (Stacey and Kramers, 1975)。本文数据表中 (表 1) 所列数据均为同一测量点连续 5 次分析的平均值, 相对误差为 1σ 。样品年龄采用 $n(^{206}\text{Pb})/n(^{238}\text{U})$ 表面年龄的平均值, 误差置信度为 95%。

3 分析结果

阴极发光影像显示, 样品 LGP-20N1 的锆石晶体为自形长柱状, 长宽比为 2: 1 ~ 4: 1, 晶面光滑, 晶棱平直, 具清晰的震荡环带结构, 为典型岩浆型锆石 (图 4)。本文共选择 17 颗锆石进行了 U—Th—Pb 同位素分析, 结果列于表 1。

测点 2.1、4.1、8.1、13.1 和 16.1 的 U 含量为 $14 \times 10^{-6} \sim 203 \times 10^{-6}$, Th 含量为 $8 \times 10^{-6} \sim 110 \times 10^{-6}$, Th/U 比值为 0.53 ~ 0.73 (表 1), 均大于 0.4; 按照 Rubatto (2002) 的判别标准, 属于典型的岩浆成因锆石。其 U-Pb 年龄值在误差范围内具有很好的—致性, 均分布于谐和线上或附近, 表明这些锆石几乎没有 U 和 Pb 的丢失或加入, 数据可靠性高 (图 5)。这 5 个测点的 $n(^{206}\text{Pb})/n(^{238}\text{U})$ 年龄加权平均值为 $780 \pm$

12 Ma (95% 置信度, $\text{MSWD} = 1.2$) (图 5), 代表川西苏雄组层型剖面英安岩 (LGP-20N1) 锆石的结晶年龄, 即英安岩的喷发时限。

测点 3.1、5.1 ~ 7.1、10.1、11.1、14.1、15.1、17.1 的 U 含量为 $24 \times 10^{-6} \sim 286 \times 10^{-6}$, Th 含量为 $17 \times 10^{-6} \sim 221 \times 10^{-6}$, Th/U 比值为 0.41 ~ 1.61 (表 1), 均大于 0.4, 按照 Rubatto (2002) 的判别标准, 也为典型的岩浆成因锆石。

测试结果表明, 其 U-Pb 年龄在误差范围内具有很好的—致性, 均分布于谐和线上或附近 (图 5),



图3 川西新元古界苏雄组英安岩采样露头(a)与显微照片(b)

Fig. 3 The outcrop(a)and micro-photograph(b) of the dacite sample from the Neoproterozoic Suxiong Formation in western Sichuan province

F—长石; Q—石英; Mt—磁铁矿 F—feldspar; Q—quartz; Mt—magnetite

表明这些锆石 U 和 Pb 的封闭系统未遭受破坏,数据可靠性高。这9个测点的 $n(^{206}\text{Pb})/n(^{238}\text{U})$ 年龄的加权平均值为 $838 \pm 9 \text{ Ma}$ (95% 置信度, $\text{MSWD} = 0.8$) (图5), 代表康滇裂谷盆地开启前的构造岩

浆热事件。测点 12.1 的锆石具有明显的核边结构, 由于打点位置位于中心核部, 测得的数据可能代表继承年龄, 未参与计算。

表1 川西新元古界苏雄组英安岩 LGP-20N1 的锆石 U—Th—Pb 同位素分析结果

Table 1 Zircon U—Th—Pb isotopic analysis for the LGP-20N1 dacite from the Neoproterozoic Suxiong Formation in western Sichuan province

测点编号	元素含量				同位素比值						同位素年龄 (Ma)			
	$^{206}\text{Pb}^*$	U	Th	Th/U	$\frac{n(^{207}\text{Pb}^*)}{n(^{206}\text{Pb}^*)}$		$\frac{n(^{207}\text{Pb}^*)}{n(^{235}\text{U})}$		$\frac{n(^{206}\text{Pb}^*)}{n(^{238}\text{U})}$		$\frac{n(^{206}\text{Pb})}{n(^{238}\text{U})}$		$\frac{n(^{207}\text{Pb})}{n(^{206}\text{Pb})}$	
	($\times 10^{-6}$)				测值	$\pm \%$	测值	$\pm \%$	测值	$\pm \%$	测值	$\pm 1\sigma$	测值	$\pm 1\sigma$
1.1	19.70	198	133	0.7	0.0683	4.4	1.081	4.7	0.1148	1.5	700.7	10	877	91
2.1	12.00	110	78	0.73	0.0676	6.6	1.176	6.7	0.1263	1.6	766	11	855	140
3.1	30.30	248	126	0.53	0.0639	3.4	1.241	3.7	0.1409	1.4	850	11	739	73
4.1	16.00	138	70	0.53	0.0571	8.7	1.036	8.9	0.1315	1.6	796	12	497	190
5.1	34.00	286	179	0.65	0.0703	3.2	1.33	3.4	0.1372	1.3	829	10	938	65
6.1	17.50	142	221	1.61	0.075	7.9	1.45	8	0.1402	1.5	846	12	1068	160
7.1	25.80	219	126	0.59	0.0701	3.6	1.315	3.9	0.1361	1.4	822	11	931	74
8.1	1.69	14	8	0.6	0.061	84	1.08	84	0.1295	6.8	785	51	630	180
9.1	21.10	160	71	0.46	0.0768	6.6	1.59	6.8	0.1504	1.5	903	13	1116	130
10.1	3.13	24	17	0.72	0.076	22	1.5	22	0.1423	4	858	32	1101	440
11.1	13.30	108	43	0.41	0.051	25	0.98	25	0.1378	2.1	832	17	262	57
12.1	44.80	358	192	0.55	0.0677	2.7	1.356	3	0.1452	1.3	874	10	860	56
13.1	3.71	33	23	0.71	0.055	22	0.95	22	0.1252	2.6	761	19	422	49
14.1	30.90	254	123	0.5	0.0646	3.5	1.254	3.8	0.1409	1.3	850	11	761	74
15.1	13.40	109	86	0.82	0.077	13	1.48	13	0.1392	1.8	840	14	1129	260
16.1	22.90	203	110	0.56	0.0622	5.2	1.115	5.4	0.1299	1.6	787	12	682	110
17.1	26.30	216	116	0.55	0.0662	11	1.26	11	0.1378	1.6	832	12	814	230

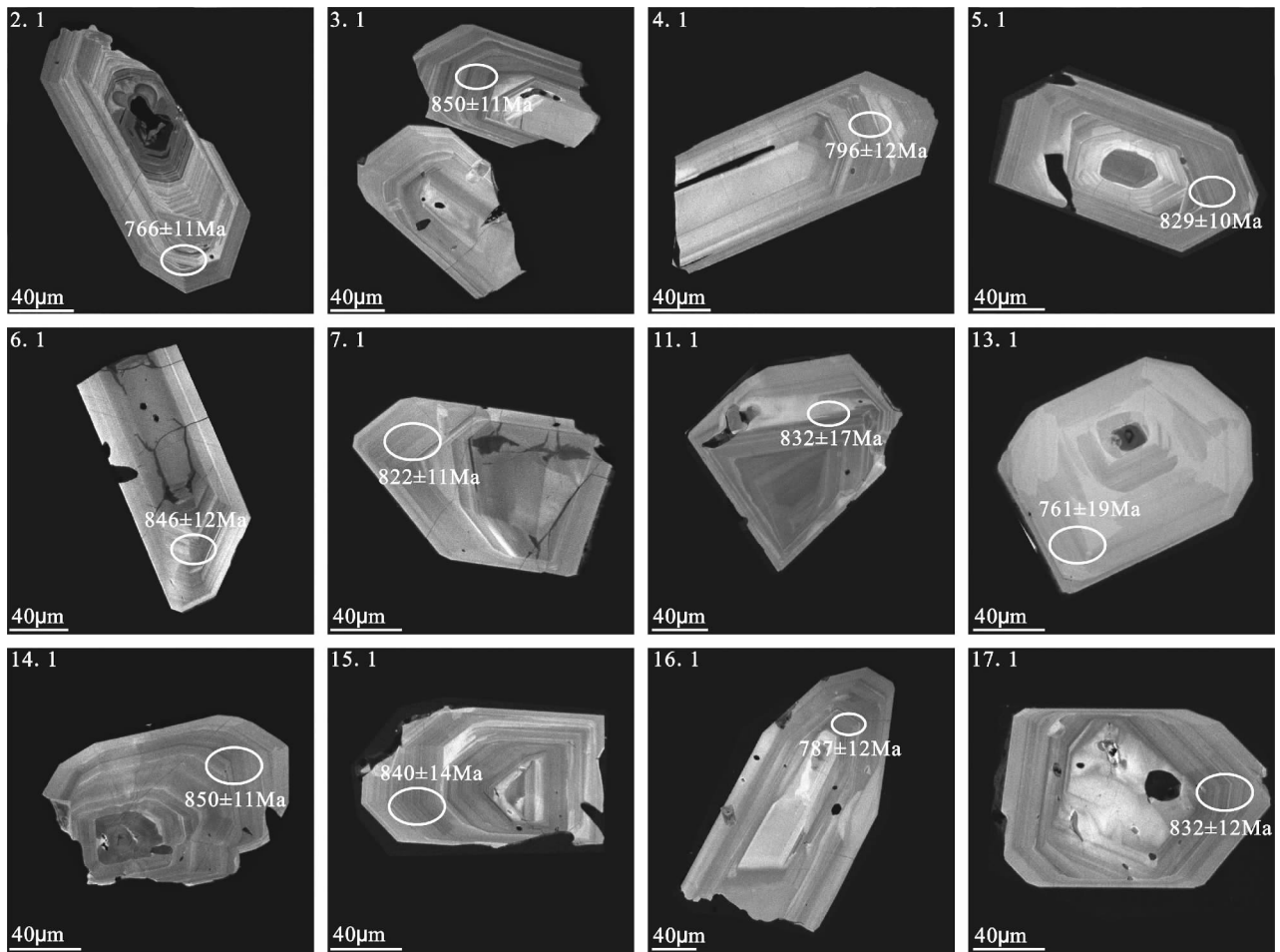


图 4 川西新元古界苏雄组英安岩锆石的阴极发光图像及相应年龄

Fig. 4 Zircon CL images and ages of the dacite from the Neoproterozoic Suxiong Formation in western Sichuan province

4 讨论

4.1 苏雄组的幕式喷发

Li Xianhua 等(2002)、谷志东等(2014a)在苏雄组底部分别获得了 803 ± 12 Ma、 794 ± 11 Ma 两个高精度年龄数据,结合本文在苏雄组中上部获得的 780 ± 12 Ma 的 SHRIMP 锆石 U-Pb 年龄,及殷继成(1984)测得苏雄组顶部玄武岩 759 Ma 的 K-Ar 年龄,表明苏雄组并非一次性喷发的产物,而是与华南新元古代的幕式构造岩浆热事件(825 Ma、800 Ma、780 Ma 和 755 Ma 四幕岩浆活动)(Ernst et al., 2008)密切相关。本文在扬子西缘新元古代康滇裂谷盆地获得苏雄组英安岩 780 ± 12 Ma 的 SHRIMP 锆石 U-Pb 年龄,应属于华南第二幕(800 Ma)构造岩浆热事件的延续。事实上,在扬子西缘广泛发生的 780 Ma 左右的岩浆活动(表 2),均是华南新元古代幕式构造岩浆热事件的表现。

对于苏雄组双峰式火山岩的成因,一直存在着不同的认识。刘鸿允等(1991)、Li Xianhua 等(2002)、谷志东与汪泽成(2014b)认为其与陆内裂谷有关,周朝宪等(1998)认为其形成于后造山环境。根据卓皆文等(2013)研究表明,新元古代康滇裂谷盆地属于大陆裂谷半地堑盆地,由主边界断裂、裂谷肩、裂谷基和调节带组成,川西苏雄组位于其调节带。调节带是裂谷盆地非常重要的组成部分,由一系列斜交裂谷轴的转换断层组成,是半地堑地极性反转的过渡地带;其特点表现为,地形上高于盆地基底,时间上晚于裂谷初始期,因此苏雄组缺乏第一幕(825 Ma)的岩浆活动,这点在前人于苏雄组及相当层底部获得的相关年代学资料(800 Ma 左右)得到了证明(Li Xianhua et al., 2002; 谷志东等, 2014a; 崔晓庄等, 2015; 卓皆文等, 2015)。

值得一提的是,最近 Jiang Zhuofei 等(2016)在川西甘洛开建桥组层型剖面顶部的凝灰质粉砂岩和

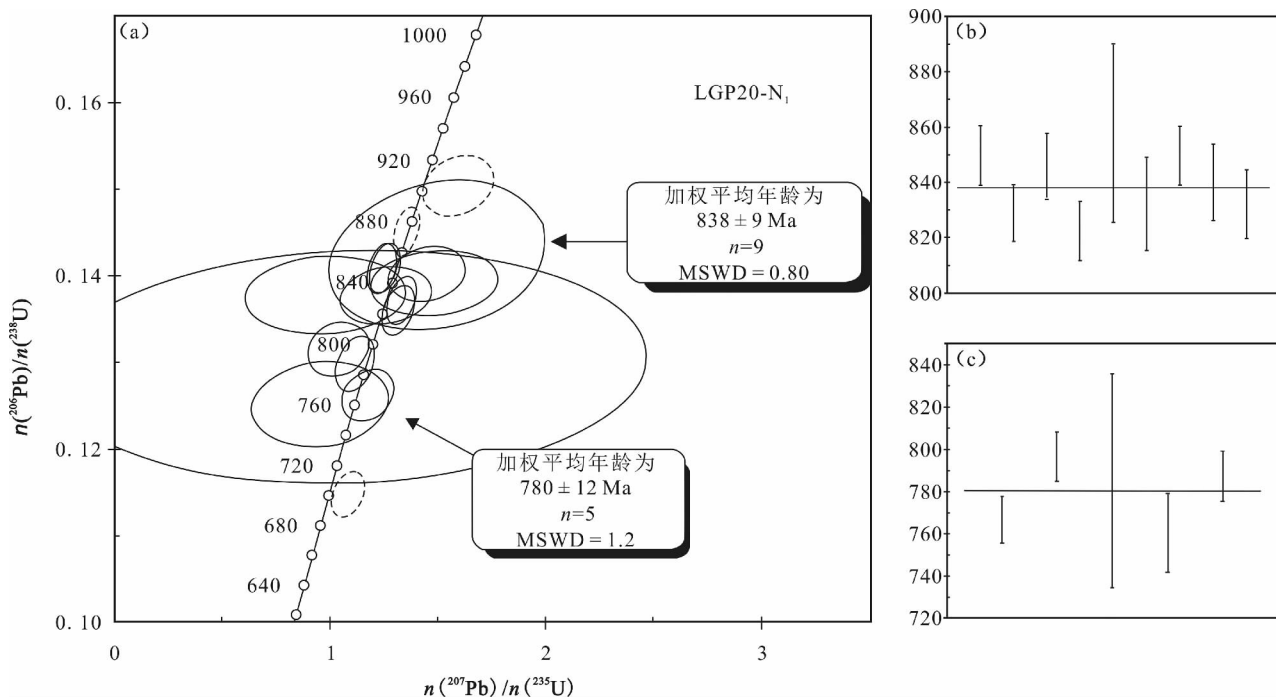


图5 川西新元古界苏雄组英安岩锆石U—Pb谐和图(a)和 $n(^{206}\text{Pb})/n(^{238}\text{U})$ 年龄图(b,c)

Fig. 5 Zircon U—Pb concordance (a) and $n(^{206}\text{Pb})/n(^{238}\text{U})$ ages (b,c)

of the dacite from the Neoproterozoic Suxiong Formation in western Sichuan province

凝灰岩分别获得 $721 \pm 7 \text{ Ma}$, $717 \pm 7 \text{ Ma}$ 的 LA-ICP-MS 锆石 U-Pb 年龄值, 认为开建桥组的顶界时限为 717 Ma。苏雄组和开建桥组为同时异相关系, 两者可以完全进行对比(卓皆文等, 2015)。因此, 苏雄组的幕式喷发(800 Ma, 780 Ma, 755 Ma, 720 Ma)可以与导致 Rodinia 超大陆裂解的地幔柱幕式岩浆活动(Ernst et al., 2008)相对应, 进一步表明其为华南新元古代幕式岩浆活动的产物。

4.2 华南大地构造属性

华南很好地保存了 850 ~ 720 Ma 的岩浆活动和沉积记录(Li Xianhua et al., 2003; Wang Jian and Li Zhengxiang, 2003)。然而, 针对这些广泛发育岩浆事件的构造背景还存在着很大的争议。Li Zhengxiang 等(2003, 2008)、Ling Wenli 等(2003)、Wang Jian 与 Li Zhengxiang(2003)、Ernst 等(2008)、Cui Xiaozhuang 等(2015)认为该岩浆活动形成于陆内裂谷环境, 与导致 Rodinia 超大陆裂解的地幔柱幕式岩浆活动有关; Zhou Meifu 等(2002, 2006)则认为其形成于岛弧环境, 与洋壳岩石圈向东插入扬子板块下有关; Wang Xiaolei 等(2004, 2006)认为其形成于碰撞后垮塌环境。

裂谷观点(Wang Jian and Li Zhengxiang, 2003;

Li Zhengxiang et al., 2003, 2008)认为, 华南新元古代的岩浆活动具有 3 个显著的特点: 双峰式、多幕式和广泛性。晋宁期约 830 ~ 745 Ma 的岩浆事件被认为是扬子对新元古代地幔柱作用的响应, 其中 830 ~ 795 Ma 和 780 ~ 745 Ma 分别对应 Rodinia 超大陆开始张裂和最终裂解两个演化阶段(Li Zhengxiang et al., 2003)。Li Zhengxiang 等(2008)进一步指出, 约 850 ~ 740 Ma 之间的剧烈裂谷岩浆活动是幕式(825 Ma, 780 Ma 和 750 Ma)地幔柱活动事件形成。Ernst 等(2008)将分布在全球不同地区的元古界火山岩浆岩群称为相应的大火成岩省, 并认为 825 Ma, 800 Ma, 780 Ma 和 755 Ma(可能还包括 720 Ma)岩浆事件与 Rodinia 超大陆裂解相关。

年代学研究表明, 在扬子西缘、北缘和东南缘, 均发育 780 Ma 的岩浆事件(表 2), 为华南裂谷盆地裂解高峰(800 Ma)的延续。且该岩浆事件与北美西部 780 Ma 的巨型岩墙群(Jefferson and Parrishi, 1989; Park et al., 1995)和澳大利亚南部 783 Ma 的 Boucaut 火山岩(Preiss, 2000)的时代也是一致的, 表明华南新元古代裂谷盆地的幕式岩浆活动和 Rodinia 超大陆裂解是同步的。

沉积学研究表明, 华南新元古代裂谷的沉降与

表 2 华南新元古代裂谷盆地系统的火成岩锆石 U-Pb 年龄统计表

Table 2 Zircon U-Pb ages about igneous rocks of Neoproterozoic rift basins in South China

构造部位	采样位置	岩石单元	样品岩性	测试方法	年龄 (Ma)	数据来源
扬子西缘	四川冕宁	康定群	花岗质片麻岩	SHRIMP	772 ± 15 773 ± 11	陈岳龙等,2004
	四川攀枝花	康定群	奥长花岗岩	SHRIMP	778 ± 11	杜利林等,2006
	四川泸定	康定群	花岗岩	SHRIMP	779 ± 6	林广春等,2006
	四川盐边	荒田组	玄武岩	SHRIMP	782 ± 53	杜利林等,2005
	四川米易	康定群	石英闪长岩	SHRIMP	775 ± 8	Li Zhengxiang 等,2003
	四川金河口	烂包坪组	凝灰岩	SHRIMP	779 ± 16	熊国庆等,2013
	四川喜德	登相营群	基性岩墙群	SHRIMP	774 ± 10	任光明等,2013
	四川甘洛	开建桥组	凝灰质岩	LA-ICP-MS	780 ± 3	Jiang Zhuofei 等,2016
	云南巧家	澄江组	凝灰岩	SHRIMP	785 ± 12	陆俊泽等,2013
	云南澄江	澄江组	凝灰岩	SHRIMP	781 ± 11	崔晓庄等,2013
扬子东南缘	贵州铜仁	鹅家坳组	凝灰岩	SHRIMP	782 ± 8 785 ± 8	汪正江等,2010
	贵州清镇	鹅家坳组	凝灰岩	SHRIMP	780 ± 9	汪正江等,2010
	浙江道林山	岩体	花岗岩	LA-ICP-MS SHRIMP	780 ± 5 775 ± 13	Wang Qiang 等,2010
	广西罗城	拱洞组	凝灰岩	SHRIMP	786 ± 6	高林志等,2013
	广西三江	长安组	凝灰岩	SHRIMP	778 ± 5	高林志等,2013
	贵州雷山	清水江组	凝灰岩	SHRIMP	774 ± 8	高林志等,2010
	贵州锦屏	隆里组	凝灰岩	LA-ICP-MS	785 ± 8	汪正江等,2013
	广西三江	拱洞组	凝灰岩	LA-ICP-MS	775 ± 13	汪正江等,2013
	广西合桐	拱洞组	凝灰岩	LA-ICP-MS	781 ± 5	崔晓庄等,2016
	湖南中部	砖墙湾组	凝灰岩	LA-ICP-MS	775 ± 3	马慧英等,2013
	贵州锦屏	清水江组	凝灰岩	SIMS	774 ± 5	Wang Xuance 等,2012
	贵州刚边	岩体	花岗斑岩	LA-ICP-MS	785 ± 4	陈建书等,2014
	贵州宰便	岩体	辉绿岩	TIMS	788 ± 3	陈建书等,2014
	安徽南部	井潭组	凝灰岩	LA-ICP-MS	779 ± 7	吴荣新等,2007
	安徽南部	井潭组	英安岩	SHRIMP	773 ± 7	吴荣新等,2007
	安徽石耳山	岩体	花岗类岩石	SHRIMP	779 ± 11	Li Zhengxiang 等,2003
	安徽石耳山	岩体	花岗类岩石	SHRIMP	778 ± 11	吴荣新等,2005
	安徽石耳山	岩体	花岗类岩石	LA-ICP-MS	776 ± 6 777 ± 11	吴荣新等,2005
浙江建德	志棠组	凝灰岩	SHRIMP	780 ± 10	尹崇玉等,2007	
扬子北缘	湖北三峡	莲沱组	凝灰岩	SHRIMP	787 ± 7	高维与张传恒,2009
	湖北大洪山	莲沱组	凝灰岩	SHRIMP	779 ± 12	Du Qiuding 等,2013
	湖北三都坪	莲沱组	凝灰质粉砂岩	SIMS U-Pb	772 ± 4 777 ± 4	Lan Zhongwu 等,2015
	甘肃碧口	碧口群	中酸性火山岩	SHRIMP	776 ± 13	闫全人等,2003
	陕西西乡	大石沟组	凝灰岩	LA-ICP-MS	789 ± 4	邓奇等,2013
	陕西毕叽沟	岩体	辉长岩	SHRIMP	782 ± 10	Zhou Meifu 等,2002
	陕西五堵门	岩体	英云闪长岩	LA-ICP-MS	789 ± 10	凌文黎等,2006
	陕西南郑	岩体	花岗斑岩	LA-ICP-MS	776 ± 6	夏林圻等,2009

充填过程具有强烈的阶段性,可划分为四个旋回 (Wang Jian and Li Zhengxiang, 2003; 江新胜等, 2012; 崔晓庄等, 2014; Wang Jian et al., 2015)。第 I 旋回始于 820 Ma, 以基性玄武岩喷发为主, 盆地基底快速沉降为特征, 沉积范围非常局限, 沉积体呈楔状, 底部为少量火山岩、火山碎屑岩和大量的洪冲积扇沉积, 主体部分为饥饿盆地深灰色页岩、碳质页

岩沉积。第 II 旋回始于 800 Ma, 以双峰式火山岩强烈喷发、沉积范围快速扩展为标志, 沉积体呈面状, 底部为厚层双峰式火山岩堆积, 主体为移地滨岸相冲积扇、河流、湖泊及滨海沉积; 780 Ma 和 755 Ma 为裂解高峰的延续。第 III 旋回的起始时间为 725 Ma, 以冰川作用开始为标志, 形成了大陆冰盖和盆地冰水沉积。第 IV 旋回沉积始于 635 Ma, 以冰川作

用结束,海平面上升,全面海侵作用为标志,形成了以碳酸盐岩为主的陆表海沉积。华南新元古代裂谷盆地的沉积一层序演化特征,与南澳、北美西部新元古代裂谷盆地的沉积一层序演化特征极为相似,且四个大的层序旋回基本可以相互对比(王剑,2000; Wang Jian and Li Zhengxiang,2001,2003)。

综上所述,扬子周缘发育与地幔柱相关的新元古代裂谷盆地系统,且华南可能是 Rodinia 超大陆裂解的重要组成单元。因此,本文研究结果支持华南处于 Rodinia 超大陆“核心”位置,是连接北美大陆和澳大利亚—南极大陆之间的桥梁(Li Zhengxiang et al.,1995,2008)。华南广泛发育的岩浆活动和沉积记录不支持岛弧模式和碰撞后垮塌模式。

5 结论

(1)川西苏雄组中上部英安岩的喷发时限为 780 ± 12 Ma。

(2)苏雄组位于康滇裂谷半地堑盆地调节带,为华南新元古代幕式岩浆活动的产物。

(3)火山活动和沉积记录均表明,扬子周缘发育与地幔柱相关的裂谷盆地系统,且华南可能是 Rodinia 超大陆裂解的重要组成单元。

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SHRIMP Zircon U-Pb Ages for the Stratotype Section of Neoproterozoic Suxiong Formation in Western Sichuan Province and Their Geological Significance

ZHUO Jiewen^{1,2)}, JIANG Zhuofei^{1,2)}, JIANG Xinsheng^{1,2)}, WANG Jian^{1,2)}, CAI Juanjuan³⁾,
XIONG Guoqing^{1,2)}, LU Junze^{1,2)}, CUI Xiaozhuang^{1,2)}, LIU Jianhui⁴⁾

1) Chengdu Center, China Geological Survey, Chengdu, 610081;

2) Key Laboratory for Sedimentary Basin and Oil and Gas Resources, Ministry of Land and Resources, Chengdu, 610081;

3) Graduate Faculty of Chinese Academy of Geological Science, Beijing, 100037;

4) Beijing SHRIMP Center, Beijing, 102206

Objectives: The age and tectonic setting of extensive Neoproterozoic volcanic activities around the margin of Yangtze Block were of great significance for reconstructing the position of South China Block in the Rodinia supercontinent. At present, only a few precise ages were obtained from the volcanoclastics of Suxiong Formation in the western margin of Yangtze Block. Therefore, a dacite sample for the stratotype section was collected from Suxiong Formation, with the purpose of determining the high-precise age in western Sichuan province.

Methods: The zircon SHRIMP U-Pb analysis was used to determine the age of the dacite sample from the Suxiong Formation, which was carried out by SHRIMP II in Beijing SHRIMP Center, Chinese Academy of Geological Sciences.

Results: Two precise ages of 838 ± 9 Ma and 780 ± 12 Ma were obtained.

Conclusions: Among them, the younger one, 780 ± 12 Ma, was interpreted as the eruption age of the dacite sample, which could represent the tectonic—magma event of the Kangdian rift basin during the period of strong extension, and the older one, 838 ± 9 Ma, represented the tectonic—magma event before the opening of the Kangdian rift basin. Through the comprehensive analysis, the Suxiong Formation was regarded as the product of episodic magmatic activity, locating at the accommodation zones of the Kangdian half graben. Consequently, a series of rift basins, which were related to the mantle plume that led to the breakup of Rodinia supercontinent, should be developed around the margin of Yangtze Block, while the South China Block could be a member of Rodinia supercontinent breakup, through the analysis of volcanic activities and sedimentary records.

Keywords: Suxiong Formation; dacite; SHRIMP zircon U-Pb age; Neoproterozoic; tectonic—magma event

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First author: ZHUO Jiewen, male, born in 1980, doctor, senior engineer, mainly engaged in sedimentology and tectonics. Email: zhuojiewen_1980@163.com

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