

Mg 同位素示踪中国东部中生代 深部碳循环

段先哲^{1,2)}, 牛苏娟¹⁾, 李赛¹⁾, 李南²⁾, 孙浩然¹⁾, 郭聪¹⁾,

肖文舟^{1,2)}, 隋清霖^{1,2)}, 冯鹏^{1,2)}, 贺海洋^{1,2)}

1) 南华大学资源环境与安全工程学院, 湖南衡阳, 421001;

2) 稀有金属矿产开发与废物地质处置技术湖南省重点实验室, 湖南衡阳, 421001

内容提要: 地球深部碳循环对全球气候变化、生命探索和岩石圈演化的研究具有重要意义。中国东部岩石圈是地球深部碳循环的重要场所, 其减薄和破坏与深部碳循环密切相关。中生代太平洋板块俯冲是制约中国东部岩石圈减薄和破坏的关键, 对华北克拉通和华南板块的大规模金属成矿作用具有重要作用。笔者等系统阐述了 Mg 示踪地球深部碳循环原理, 例举了镁同位素示踪中国东部深部碳循环的实例, 论述了中生代俯冲的古太平洋板块所释放的碳酸盐熔体/流体与地幔相互作用, 是造成中国东部地幔具有普遍的轻 Mg 同位素组成的重要原因。此外, 指出轻镁同位素的多解性, 提出多同位素联合示踪是未来研究地球深部碳循环的趋势。

关键词: 深部碳循环; 中生代; 中国东部岩石圈; Mg 同位素; 岩石圈减薄与破坏; 古太平洋板块俯冲

碳循环分为大气圈、水圈、生物圈和土壤圈之间的短周期地球表层碳循环和表层—深部系统之间的长周期碳循环。地球表层碳约占地球总碳量的10%, 主要赋存于沉积碳酸盐、硅质页岩及油页岩中 (Falkowski et al., 2000), 而地球深部碳储量占地球总碳量的90%以上 (Dasgupta and Hirschmann, 2010; Yang Wei et al., 2012), 强烈影响着深部地幔的组成和结构。

地表碳(包含有机碳和无机碳)通过板片深俯冲作用进入地球深部系统, 其中部分碳伴随着岛弧火山脱气作用携至地表, 其他大部分残余碳稳定地滞留在深部地幔中, 这一过程构成了深部碳循环 (Zhang and Liou, 1994; Zhu Yongfeng and Ogasawara, 2002; Dasgupta et al., 2004; Yaxley and Brey, 2004; Kawamoto, 2006; Poli et al., 2009; Keshav and Gudfinnsson, 2010; 张立飞等, 2017; Plank et al., 2019; Hilton et al., 2020; Wu Hailin et al., 2022)。地球深部碳循环的研究对全球气候变化、生命探索和地幔演化机制具有重要意义 (Hoffman and Nehrkom, 2009; Hyde et al., 2000;

Hoffman and Schrag, 2002; Armstrong et al., 2019; Manthilake et al., 2021)。

中国东部岩石圈自古生代以来发生了减薄和破坏, 从太古代厚的(>200 km)、冷的、难熔型的古老地幔转变为新生代薄的(<80 km)、热的、饱满型的新生地幔, 中生代是其减薄和破坏的关键时期; 中国东部岩石圈这种减薄和破坏机制可能与华北和华南中生代众多大规模金属矿床的形成有着密切的联系, 因而受到了国内外学者的广泛关注 (Fan Weiming and Menzies, 1992; Griffin et al., 1998; Xu Yigang et al., 2002; Wu Fuyuan et al., 2003; Zhang Hongfu et al., 2003; Rudnick et al., 2004; Zheng Jianping et al., 2004; Niu Yaoling, 2005; Liu Yongsheng et al., 2005; 李曙光, 2015, 2017; Li Shuguang and Wang Yang, 2018; Liu Jin et al., 2020; Duan Xianzhe et al., 2020; Chang Zeguang et al., 2021; Ji Zheng et al., 2021; Wang Chen et al., 2021; Sun Pu et al., 2021)。

中国东部岩石圈是深部碳循环的重要场所, 它的减薄和破坏与地球深部碳循环密切相关, 反映了



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作者简介: 段先哲, 男, 1985年生, 博士, 副教授, 主要从事岩石/矿床地球化学等方面研究; Email: duanxianzhe@126.com。

深部碳循环的地球动力学效应 (Huang Jian et al., 2015a; Li Shuguang et al., 2017; 李曙光和汪洋, 2018)。因此, 研究中国东部中生代深部碳循环对于制约中国东部岩石圈减薄和破坏机制有着重要意义。然而, 对于中国东部岩石圈破坏机制和减薄机理、深部碳的赋存形式和储量、板块俯冲作用的参与程度、岩石圈破坏的范围和程度等仍然存在较大的争议 (Follett et al., 2014; Liu Dong et al., 2015; Tappe et al., 2017; 陈春飞, 2018; 刘勇胜等, 2019; 沈骥等, 2019; 汤艳杰和张宏福, 2019)。

近年来, 岩石地球化学和同位素示踪研究在识别中国东部岩石圈物质组成方面取得了重大进展 (Huang Fang et al., 2011; 李曙光, 2015, 2017; Huang Jian et al., 2016a, d; Li Shuguang et al., 2017; 宗克清和刘勇胜, 2018; Tian Hengci et al., 2018; Wang Xiaojun et al., 2018)。前人依次通过模拟计算、实验及两者结合的方式得到一些研究成果。随着非传统稳定金属同位素的发展, 发现其示踪碳酸盐金属离子与地幔存在明显差异, 这对研究深部碳循环有重要意义。Mg 同位素是理想的示踪深部碳循环工具, 其分馏作用在早期太阳系演化、变质岩和地幔岩的平衡温度、大陆风化、植物生长、酶合成和古气候演变等诸多科学问题上发挥了制约作用。笔者等系统阐述了镁同位素示踪再循环碳酸盐对中国东部中生代深部碳循环形成制约, 论述了中国东部岩石圈减薄和破坏与深部碳循环的动力学效应及中国东部岩石圈 Mg 同位素研究进展, 以期为中国东部岩石圈减薄与破坏机制的进一步研究提供参考和依据。

1 同位素示踪

地表和深部“碳的命运”是关注的热点课题。近年来, 众多学者通过同位素示踪的方法对俯冲碳的循环进行研究。

1.1 C/Ca 同位素

碳同位素被广泛应用于研究深部碳循环, 碳同位素可以有效区分有机碳 OC ($\delta^{13}\text{C}$ 值小于 -15‰) 和无机碳 IC ($\delta^{13}\text{C}$ 值大于 -10‰ ; Deines, 2002; Cartigny, 2005), 约 95% 的火山作用释放的大部分都是无机碳 (例如壳源碳、板块俯冲再循环碳酸盐和地幔原生碳), 而碳同位素示踪无机碳无效, 并且在岩浆去气过程中会发生显著分馏 (Javoy et al., 1982; Matthey, 1991; Graham et al., 2018)。前人提出二价金属同位素 (Ca, Mg) 受去气分馏作用影响

较小, Ca 同位素在示踪过程中也显示出极大潜力 (Huang Shichun et al., 2010, 2011; Kang Jinting et al., 2016; Zhao Xinmiao et al., 2017; Zhu Hongli et al., 2018)。例如, Huang Shichun 等 (2011) 观察到一些夏威夷玄武岩的 $\delta^{44/40}\text{Ca}$ 与 $^{87}\text{Sr}/^{86}\text{Sr}$ 成负相关, 认为地幔遭受了古老沉积碳酸盐的俯冲。而 Ionov 等 (2019) 认为 $\delta^{44/40}\text{Ca}$ 值并不能简单归因于再循环碳酸盐, 西伯利亚克拉通含交代方解石的地幔包体研究发现其与饱满地幔同位素组成几乎没有差别 ($\delta^{44/40}\text{Ca} = 0.81\text{‰} \sim 0.83\text{‰}$)。在高精度同位素测试方面, 质谱测量过程中 ^{40}Ar 强烈干扰且 ^{40}Ca 和 ^{42}Ca (或 ^{44}Ca) 同位素丰度存在很大差异, 因此 Ca 同位素示踪对象要有选择性。目前采用的同位素示踪方法的有 Ca、Zn、Mg、Li 等同位素结合示踪, 也取得了相关成果, 本文主要叙述关于 Mg 同位素的相关内容。

1.2 Mg 同位素

镁是宇宙中第九丰富和地球上第四丰富的元素, 广泛分布于碳酸盐和硅酸盐矿物中。>99.9% 镁分布在地幔橄榄石和辉石中, 0.031% 主要赋存在地壳辉石和橄榄石中, 总体上从地幔到洋壳, 镁的含量是逐渐降低的, 水圈中镁 $<0.001\text{‰}$ (Liu Fang et al., 2017a)。镁作为主量元素, 有 ^{24}Mg 、 ^{25}Mg 和 ^{26}Mg 3 个稳定同位素, 在自然界中丰度分别是 78.99%、10.00% 和 11.01%, ^{24}Mg 和 ^{26}Mg 之间的相对质量差高达 8%, 巨大的质量差使其足以在地质过程中发生显著的质量分馏, 这也是镁同位素成为理想地质示踪剂的条件之一。相比 Ca 同位素而言, 海相碳酸盐岩和地幔之间的 Mg 同位素组成差值可达 5‰, 同位素测试方法更加简洁、高效。因此, Mg 同位素是迄今深部碳循环应用最广泛的示踪剂。

$\delta^{26}\text{Mg}$ 、 $\delta^{25}\text{Mg}$ 计算公式 (Galy et al., 2001) 如下:

$$\delta^{26}\text{Mg} = \left\{ \frac{[n(^{26}\text{Mg})/n(^{24}\text{Mg})]_{\text{sample}}}{[n(^{26}\text{Mg})/n(^{24}\text{Mg})]_{\text{DSM-3}}} - 1 \right\} \times 1000\text{‰}$$

$$\delta^{25}\text{Mg} = \left\{ \frac{[n(^{25}\text{Mg})/n(^{24}\text{Mg})]_{\text{sample}}}{[n(^{25}\text{Mg})/n(^{24}\text{Mg})]_{\text{DSM-3}}} - 1 \right\} \times 1000\text{‰}$$

DSM-3 为标样。

1.3 表生过程中的 Mg 同位素组成

表生圈层 (水圈、洋/陆壳) Mg 同位素组成差异明显 ($\delta^{26}\text{Mg} = -5.5\text{‰} \sim +1.8\text{‰}$) (Teng Fangzhen, 2017)。Teng Fangzhen 等 (2010b) 对美国南卡的辉绿岩岩墙的风化剖面进行了研究, 发现随着剖面深度越接近地表岩石, 其风化程度越强, 风化残余物越

富集重 Mg 同位素组成, 轻 Mg 同位素优先释放到了海洋中。海水有均一的镁同位素组成, 其 $\delta^{26}\text{Mg}$ 平均值为 $-0.83\text{‰} \pm 0.07\text{‰}$ (2SD, $n = 25$) (Ling Mingxing et al., 2011; 孙剑等, 2012)。大部分海水是由流经地表经大陆风化作用的河水和洋壳经海水蚀变溶解的 (Tipper et al., 2006b; Brenot et al., 2008), 不同经纬度和深度的海水与其盐度和温度无关 (Foster et al., 2010; Ling Mingxing et al., 2011)。河流水作为海水的一个主要来源, 其 $\delta^{26}\text{Mg}$ 平均值是 -1.1‰ , 具有比海水更轻的镁同位素组成 (Tipper et al., 2006a; 何学贤等, 2008)。控制河流水的镁同位素组成因素之一是基岩种类, 流经硅酸岩河水相比原始硅酸岩更富集轻 Mg 同位素 (Tipper, 2006a; 高庭, 2016), 而流经碳酸盐河流水 ($\delta^{26}\text{Mg} = -1.7\text{‰} \sim -1.37\text{‰}$) 比流经硅酸盐河流水 ($\delta^{26}\text{Mg} = -1.1\text{‰} \sim -0.77\text{‰}$) 具有更低的 $\delta^{26}\text{Mg}$ 值 (Brenot et al., 2008), 这与碳酸盐 $\delta^{26}\text{Mg}$ 值较低是相一致的。Teng Fangzhen 等 (2016) 对海洋沉积物 (Site 144 和 Site 543) 的 Mg 同位素组成 ($\delta^{26}\text{Mg} = -0.76\text{‰} \sim +0.52\text{‰}$) 开展研究, 发现富含碳酸盐岩 (Site 144) 有更轻的镁同位素组成。Wang Shuijiong 等 (2015a) 对英国湖区开展研究, 发现含碳酸盐岩的南部的 Mg 同位素组成低于不含碳酸盐岩的北部, 有残留相的南部湖区和北部有一致的 Mg 同位素组成 (较重 Mg 同位素), 若玄武岩源区中有再循环碳酸盐混入, $\delta^{26}\text{Mg}$ 值与 Mg/Al 、 $\text{CaO}/\text{Al}_2\text{O}_3$ 和 CaO/TiO_2 呈负相关性 (Wang Shuijiong et al., 2015a)。

1.4 地幔中的 Mg 同位素组成

地幔具有均一的 Mg 同位素组成 ($\delta^{26}\text{Mg} \approx -0.25\text{‰}$, 平均值与球粒陨石 ($\delta^{26}\text{Mg} = -0.28 \pm 0.06\text{‰}$) 相似 (Teng Fangzhen et al., 2010a)。全球分布的洋中脊玄武岩 (MORB)、洋岛玄武岩 (OIB) 和大陆玄武岩的 Mg 同位素组成有均一性 ($\delta^{26}\text{Mg} = -0.23 \pm 0.19\text{‰}$, 柯珊等, 2011)。

地幔 Mg 同位素组成 ($\delta^{26}\text{Mg} \approx -0.25 \pm 0.07\text{‰}$, Teng Fangzhen et al., 2010a) 与沉积碳酸盐岩的 Mg 同位素组成 ($\delta^{26}\text{Mg} = -5.31\text{‰} \sim -1.09\text{‰}$, Young and Galy, 2004; Tipper et al., 2006a; Brenot et al., 2008; Hippler et al., 2009) 存在明显差异, 因而幔源岩浆岩的轻镁同位素特征可以指示沉积碳酸盐。例如, 中国东部白垩纪到第四纪时期的玄武岩具有轻 Mg 同位素和低 $^{87}\text{Sr}/^{86}\text{Sr}$ 值, 表明是沉积碳酸盐岩与地幔相互作用的结果 (Yang Wei et al., 2012)。

1.5 板块俯冲过程中的 Mg 同位素组成

俯冲沉积物中的沉积碳酸盐是碳主要的存在形式。大洋板块从形成到运送至俯冲带期间与海水相互作用产生沉积物, 新鲜的大洋玄武岩和橄榄岩发生蚀变, 蚀变洋壳中大部分碳酸盐并不会在俯冲脱水过程中被释放, 而是被携带到地幔深部, 从而显著改变地幔的物理化学性质 (如固相线, 地震波速, 密度等) (Alt and Teagle, 1999; Thomson et al., 2016)。

Wang Shuijiong 等 (2012) 和 Wang Xuance 等 (2015b) 报道了南非和西非克拉通金伯利岩中榴辉岩 (CO_2 含量并不高) 具有低 $\delta^{26}\text{Mg}$ 值 (可到 -1.38‰), 认为这些俯冲进入地幔的榴辉岩的低 $\delta^{26}\text{Mg}$ 值特征可能继承了蚀变洋壳的特征, 也可能是在俯冲过程中与沉积碳酸盐岩发生同位素交换反应的结果; Huang Jian 等 (2016a) 同样在榴辉岩的研究中也认为板块俯冲变质脱水过程不导致显著的 Mg 活动和分馏; Wang Shuijiong 等 (2014a) 研究了新元古代扬子板块北缘和大别山造山带不同变质程度 (变质程度由浅到深分别为绿片岩、斜长角闪岩和榴辉岩) 和原岩 (玄武质变质岩), 得出随着变质程度的加深, MgO 和 Mg 同位素含量无变化, 岩石含水量下降, 板块俯冲变质脱水过程不发生显著的 Mg 同位素分馏; Wang Shuijiong 等 (2014b) 考察研究了苏鲁超高压变质带的荣成 UHP 大理岩及其包裹的碳酸盐化榴辉岩, 发现普通榴辉岩与其原岩的 $\delta^{26}\text{Mg}$ 与地幔值 ($\delta^{26}\text{Mg} = -0.25 \pm 0.07\text{‰}$) 一致, 但包裹在大理岩中的碳酸盐化榴辉岩有更低的 $\delta^{26}\text{Mg}$ 值 ($\delta^{26}\text{Mg} = -0.6 \sim -1.9\text{‰}$)。Huang Jian 等 (2015b) 发现东太平洋洋脊钻探 1256D 岩芯样品与地幔有相似的 Mg 同位素组成, 表明蚀变玄武岩与新鲜玄武岩有相似的 Mg 同位素组成。

值得指出的是, 深海橄榄岩发生蛇纹石化蚀变 ($t = 200^\circ\text{C}$), 其 Mg 同位素为一般地幔组成。当深海橄榄岩发生更低温的海底风化作用时, 轻 Mg 同位素被海水带走, 导致深海橄榄岩富集较重的 Mg 同位素组成 (Liu Xin et al., 2017b)。蚀变深海橄榄岩脱水可产生富 MgO 和具有较重的 Mg 同位素组成流体 (Teng Fangzhen et al., 2016)。

1.6 地幔部分熔融和岩浆过程中的 Mg 同位素组成

碳酸盐岩被俯冲到地幔深处并且在俯冲过程中可能经历了脱气反应或者部分熔融过程 (Wyllie, 1989; Wallace and Green, 1998; Hammouda, 2003;

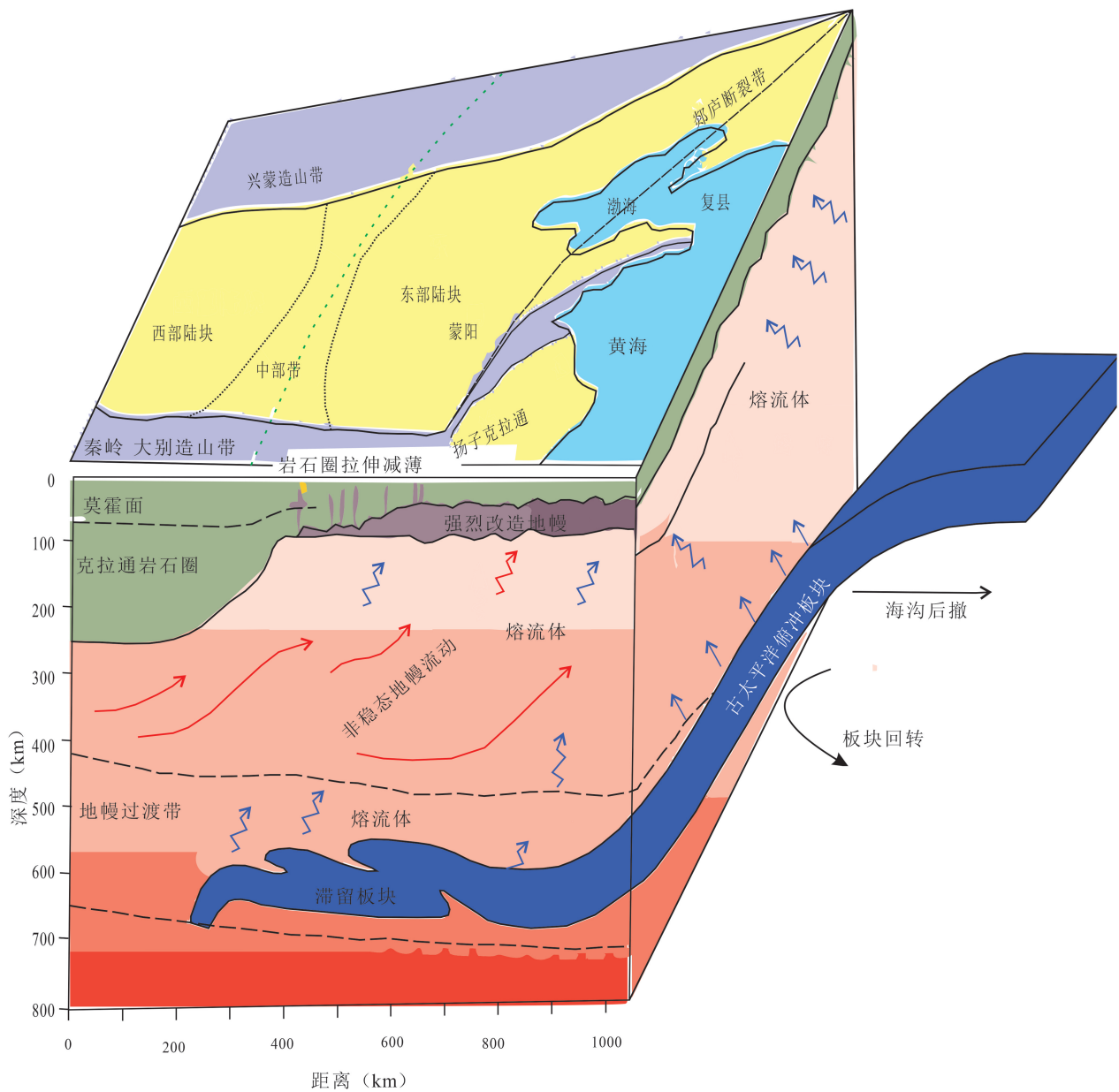


图 1 中生代中国东部岩石圈地幔改造与古太平洋板块俯冲示意图(改自汤艳杰,2021)

Fig. 1 Schematic diagram of the transformation of the lithospheric mantle beneath Eastern China and the subduction of the paleo-Pacific Plate in the Mesozoic(modified from Tang Yanjie, 2021&)

Ducea et al. , 2005; Doucelance et al. , 2014; Collins et al. , 2015; Zhong Yuan et al. , 2017; 陈春飞, 2018)。在地幔部分熔融、中—基性岩浆分异和变质过程中 Mg 同位素几乎不分馏(Teng Fangzhen et al. , 2007; 朱祥坤等, 2013), 未遭受陆壳混染的玄武质岩浆同位素组成可直接代表其地幔源区的同位素组成(Teng Fangzhen et al. , 2007; Handler et al. , 2009; Bourdon et al. , 2010; Liu Sheng’ ao et al. , 2010; Teng Fangzhen et al. , 2010b; Huang Fang,

2011; Philip et al. , 2011; Yang Wei et al. , 2012)。中国东部岩石圈在遭受板块俯冲过程中会发生部分熔融, 中高温的深部地幔会向较低温的浅部上涌减压产生玄武质岩浆, 但并不产生显著的 Mg 同位素分馏。玄武岩和地幔橄榄岩($\delta^{26}\text{Mg} \approx -0.25\text{‰}$)的 Mg 同位素组成有均一性和相似性, 接近球粒陨石的平均值($\delta^{26}\text{Mg} = -0.28 \pm 0.06\text{‰}$; Teng Fangzhen et al. , 2010a)。Liu Sheng’ ao 等(2011)发现花岗岩全岩和单矿物的 Mg 同位素组成

($\delta^{26}\text{Mg} = -0.21 \pm 0.07\text{‰}$)与地幔在误差范围内一致,富Mg矿物的分离结晶作用也不会导致显著Mg同位素分馏(Liu Sheng'ao et al., 2011; 柯珊等, 2011)。

2 中国东部岩石圈减薄与深部碳循环

中国东部主要由华南板块、华北克拉通和中国东北(即兴蒙地块)组成,其性质和厚度发生了强烈改变,是国内外学者研究的热门课题(图1)。古生代,华北克拉通有典型的克拉通型岩石圈地幔特征,主要成分是难熔的方辉橄榄岩,岩石圈厚度为200 km(郑建平和路凤香, 1999; Zheng Jianping et al., 2001; Zhang Hongfu et al., 2008; 吴福元等, 2008; Tang Yanjie et al., 2013)。中生代,华北克拉通和华南板块发生了大范围的岩浆活动和地壳形变,其稳定性遭受了不同程度破坏(郭锋等, 1997, 1998; Fan Weiming et al., 2000; Zheng Jianping et al., 2001; Li Xianhua et al., 2004; Wang Yuejun et al., 2008; 朱日祥等, 2012; Liu Chuanzhou et al., 2012)。新生代,华北克拉通和华南板块岩石圈厚度是60~80 km,岩石圈地幔在物质组成和性质上具有类似大洋型较年轻的地幔特征,由主量元素饱满、Sr—Nd同位素组成亏损的二辉橄榄岩组成(Fan Weiming and Menzies, 1992; Menzies et al., 1993; 汤艳杰等, 2019; Fan Weiming et al., 2000)。

关于减薄峰期大体分两个阶段,第一个减薄峰期是中生代中—晚期,这一阶段华北克拉通与整个东部岩石圈的剧烈破坏一致(Gao Shan et al., 2008; Meng Fanxue et al., 2015)。另一个减薄高峰是新生代晚期(25 Ma后),岩石圈厚度继续减薄到至今约70 km(Xu Yigang et al., 2004; Zhang Hongfu, 2005; Zheng Jianping et al., 2007)。对于减薄,李曙光(2015)认为应该要把再循环碳酸盐对地幔固相线的影响作为一个考虑因素。

中国东部岩石圈物理化学性质的强烈改变可能诱发了燕山期广泛的多金属矿化,形成了大规模的可有、稀有等多金属矿床(赵振华等, 1998, 2000; 黄诚和李晓峰, 2013; 毛建仁等, 2014)。因此,中生代对于研究华北和华南板块的成矿规律有重要意义。

众多学者研究了中国东部晚白垩世—新生代地幔橄榄岩包体含水量(Xia Qunke et al., 2013; Hao Yantao et al., 2012, 2016a, b; Li Yanqing et al., 2015; 徐义刚等, 2018),发现岩石圈减薄有不均一

性特征。典型和非典型克拉通岩石圈地幔的水含量分别为 $(119 \pm 54) \times 10^{-6}$ 和 $(78 \pm 45) \times 10^{-6}$,华北 $(34 \pm 34) \times 10^{-6}$ 和东北 $(47 \pm 32) \times 10^{-6}$ 岩石圈地幔的水含量低于该标准,表明其都发生了岩石圈减薄。华北克拉通岩石圈中西部水含量低于软流圈地幔,这与华北克拉通中西部受后期减薄影响小相一致,其厚度大约是~200 km。而其东部含水量远高于软流圈地幔,这显示出华北克拉通岩石圈显著减薄发生在最东部(新生的岩石圈地幔),其厚度80~120 km(渤海地区厚度60~70 km; Zheng Tianyu et al., 2006; Chen Bin et al., 2009; Chen Ling et al., 2014)。华南岩石圈含水量在中、新生代两个时代差别不大,水含量都较高为 $(90 \pm 45) \times 10^{-6}$,与岩石圈减薄发生在晚中生代特征相一致。

一些学者(如 Jiang Yaohui et al., 2009; Wang Xuance et al., 2011)认为华北克拉通的破坏和减薄是中—晚中生代太平洋板块俯冲所导致的。在板块俯冲过程中,大部分样品经历了地壳混染作用,无法鉴别其是否真的经历了再循环过程,因此几乎无法直接从地幔中找到再循环碳的可靠证据。

3 俯冲碳的证据

3.1 实验模拟

Wyllie等(1975)模拟了地幔橄榄岩的部分熔融,采用 CO_2 —CMS(CaO — MgO — SiO_2)体系在>70 km处模拟出了富碳酸盐的岩浆。Dalton等(1998)在相同系统下模拟并发现3~7 GPa压力下碳酸盐化二辉橄榄岩的固相线熔体成分相当于碳酸岩。Grassi等(2010, 2011)模拟了碳酸盐化泥质岩固相线与俯冲温压的关系,发现伴随着固相线温度的降低产生了富钾的碳酸盐(Ca — Fe — Mg)熔体(6~9 GPa是第一个熔融区)。众多实验研究测定了部分熔融过程中碳酸盐种属固相转化P—T界限(Sato and Katsura, 2001; Morlidge et al., 2006; Tao Renbiao et al., 2014),这一结果与Mg—Sr同位素制约玄武岩源区的碳酸盐种属(菱镁矿主要+白云石少量)相一致。Dasgupta等(2007, 2013),朱日祥和郑天愉(2009),Huang Jian和Xiao Yilin(2016d),Li Shuguang等(2017),李曙光和汪洋(2018)用数值方法和实验模拟中国东部碳酸盐熔体(高T—P),认为大多数的碳酸盐在太平洋板块俯冲过程中留存下来并进入了深部地幔。

3.2 计算模拟

Yang Wei等(2012)通过计算模拟发现纯硅酸

盐熔体交代难以形成低 $\delta^{26}\text{Mg}$ 的地幔源区,必须少量的碳酸岩熔体才能形成,从而认为华北克拉通中、新生代玄武岩的低 $\delta^{26}\text{Mg}$ 值很可能源自再循环碳酸盐对玄武岩源区的交代作用。Huang Jian 和 Xiao Yilin (2016d) 和 Li Shuguang 等 (2017) 通过计算模拟和同位素相结合的方法研究了中国东部岩石圈自白垩世之后的碱性玄武岩,得出玄武岩的地幔源区碳酸岩种属主要是白云石+菱镁矿,地幔源区混入 <10% 的再循环碳酸盐就能产生中国东部 <110 Ma 玄武岩的轻 Mg 同位素组成 (Huang Jian and Xiao Yilin, 2016d; 田恒次等, 2017), 但其忽略了俯冲蚀变洋壳和泥质沉积物这一影响因素。因此仅靠简单的模拟计算方法可能会造成结果单一的特点,若要全面反映俯冲过程中的深部碳循环,探究深部碳循环过程还需要借助其他手段。

3.3 计算模拟—实验相结合

柯珊等 (2011) 分别用 DAR 模型、AF 模型模拟计算和实验相结合验证 Mg 同位素的分馏,得到分馏结果不一致,认为可能是模拟或实际观测的 Mg 同位素分馏未达到平衡或者白云岩实际形成与理论过程不同 (甯濛等, 2018)。因而,通过实验和计算模拟结合可能不太准确表达出地质历史过程,甚至可能有错误引导性,需要结合非传统性同位素才能反演出更准确的深部碳循环过程。

3.4 碳酸盐熔—流体交代岩石圈地幔

碳酸盐熔体是一种重要的地幔交代介质和深部碳循环作用的指示物,长期的碳酸岩熔体交代作用可能是碳储库形成的因素之一,也是造成中国东部岩石圈地幔组成和性质改变的重要因素 (Dai Baozhang et al., 2008; Jiang Yaohui et al., 2009; Wang Xuance et al., 2011; Li Xiyao et al., 2014; Huang Jian et al., 2015a; 杨金豹等, 2015)。深部碳循环过程中大量地表的碳酸盐 (或者 CO_2) 或者碳酸盐流体被俯冲进入地幔,极大降低地幔熔融的固相线,从而诱发深部地幔部分熔融,中国东部岩石圈的减薄与破坏被碳酸盐交代作用和部分熔融过程所制约。Li Shuguang 等 (2017) 发现中国东部碱性玄武岩有典型的碳酸盐熔体特征,如高 $\text{CaO}/\text{Al}_2\text{O}_3$ 比值、低 FeO/MnO 值和显著的 K、Pb、Hf、Ti 等微量元素负异常,证实了岩石圈地幔大多数发生了周边俯冲板块的含碳酸盐熔流体的交代作用。中生代,初始的碳酸盐熔体交代—熔融作用弱化了岩石圈地幔的强度,为中国东部岩石圈的最终破坏提供了关键前提条件。不同比例碳酸盐熔体改造了岩石圈的

再循环地壳物质,大规模的碳酸盐化交代—部分熔融造成了岩石圈的减薄与破坏。

碳酸盐化洋壳部分熔融产生碳酸盐熔体在地幔过渡带 410 ~ 660 km,具有低 $\text{Gd}_\text{N}/\text{Yb}_\text{N}$ 值, LREE 富集和 HREE 亏损特征,其向上渗透并交代底部岩石圈地幔可形成碳酸盐化橄榄岩,随着太平洋板块的俯冲和东部岩石圈的减薄与破坏,碳酸盐交代熔融深度 <130 km 和逐渐靠近 70 km (Huang Jian et al., 2015b)。

4 中国东部岩石圈轻 Mg 同位素组成

中国东部是环太平洋火山带的边缘组成部分,是目前世界上最活跃的构造—岩浆区域之一 (Zhou Xinghua and Armstrong, 1982),也存在一条巨大的 Mg 同位素异常区,与滞留在地幔过渡带中的太平洋俯冲板片范围一致 (Fukao et al., 1992; Huang Jinli and Zhao Dapeng, 2006)。中国东部地震层析成像表明—晚中生代,太平洋板块已经开始俯冲至中国东部 (Huang Jinli and Zhao Dapeng, 2006; Zheng Yongfei, 2012)。大量幔源火山岩的镁同位素研究表明,俯冲板块携带的碳酸盐被折返至上地幔,在地幔过渡带形成碳酸化橄榄岩,成为各种玄武岩的来源,可能揭示出中国东部是一个巨大的再循环碳库。

玄武岩的 Mg 同位素分馏为探索中国东部深部碳循环提供了约束。Yang Wei 等 (2012) 最先研究了华北克拉通中—新生代玄武岩的 Mg 同位素组成,发现辽西义县 >120 Ma 的玄武岩具有地幔类似的 Mg 同位素组成 ($\delta^{26}\text{Mg} = -0.31\text{‰} \sim -0.25\text{‰}$),而阜新和太行山 <110 Ma 的玄武岩 $\delta^{26}\text{Mg}$ 值 ($\delta^{26}\text{Mg} = -0.60\text{‰} \sim -0.42\text{‰}$) 低于地幔,其地球化学特征表现为 MORB 型的 Sr—Nd—Pb 同位素、高的 Ce/Pb 和 Nb/U 比值和类似地幔 HIMU 端元的高 U/Pb 和 Th/Pb 值,因此认为低 $\delta^{26}\text{Mg}$ 值可能与俯冲洋壳的含碳酸岩熔体交代地幔有关。Huang Jian 等 (2015b) 报道了华南新生代碱性玄武岩表现出典型碳酸盐熔体特征 (如 K、Pb、Hf、Ti 负异常),具有轻镁同位素组成特征 ($\delta^{26}\text{Mg} = -0.60\text{‰} \sim -0.35\text{‰}$),认为再循环碳酸盐熔体交代地幔造成了玄武岩的轻 Mg 同位素组成。李曙光 (2015) 认为中国东部地幔部分熔融产生的岩浆 Mg 同位素组成要轻于地幔,很可能是由于碳酸盐的交代作用造成。Tian Hengci 等 (2016) 对中国东北五大连池和二克山玄武岩展开研究 ($\delta^{26}\text{Mg} = -0.57\text{‰} \sim -0.46\text{‰}$),其富集 EM-I 型 Sr—Nd—Pb 同位素,如高 $\text{CaO}/\text{Al}_2\text{O}_3$, Ba/Rb 等

有典型碳酸盐岩交代特征。Hu Yan 等(2016)报道了华北汉诺坝石榴子石辉石岩有低 Ti/Eu 和轻的 Mg 同位素组成($\delta^{26}\text{Mg} = -1.37\text{‰} \sim -1.47\text{‰}$),认为是再循环碳酸盐熔体与橄榄岩反应的结果。Li Shuguang 等(2017)认为太平洋板片携带大量沉积碳酸盐俯冲到了软流圈地幔,造成了中国东部晚白垩纪和新生代玄武岩的低 $\delta^{26}\text{Mg}$ 值特征。因此,上述研究表明中国东部大陆玄武岩的低 $\delta^{26}\text{Mg}$ 值很可能是地幔源区卷入再循环碳酸盐的结果(图 2)。

综合地震层析影像资料和前人对 Mg 同位素研究表明,俯冲至中国东部岩石圈并滞留在地幔过渡带的太平洋板块,其释放的碳酸盐熔体与地幔相互作用,很可能是造成东部地幔 Mg 同位素组成偏轻的重要原因(Xiao Yan et al., 2013; Huang Jian et al., 2015b)。因此,俯冲的太平洋板块在中国东部岩石圈减薄与破坏过程中发挥着重要作用。

除中国东部的“轻镁同位素异常”,世界各地均发现轻镁同位素的幔源岩岩浆并将其解释为与地幔源区中再循环碳酸盐有关(Liu Dong et al., 2015; Liu Fang et al., 2017a; Wang Shuijiong et al., 2016b; Wang Xiaojun et al., 2018; Ke Shan et al., 2016; Dai Liqun et al., 2017; Tian Hengci et al.,

2018; Hoang et al., 2018; Kim et al., 2019)。例如,Wang Shuijiong 等(2016b)报道了具有 HIMU 特征的新西兰板内玄武岩的 Mg 同位素组成($-0.47\text{‰} \sim -0.06\text{‰}$),认为低 $\delta^{26}\text{Mg}$ 异常是再循环的碳酸盐洋壳(榴辉岩)加入地幔源区的结果。Tian Hengci 等(2018)研究了较轻 Mg 同位素组成的腾冲玄武岩($-0.51\text{‰} \sim -0.45\text{‰}$),将低 $\delta^{26}\text{Mg}$ 值异常归因于超临界流体溶解白云石的结果。Hong 等(2018)报道了越南中部和南部板内晚新生代玄武岩的轻 Mg 同位素组成($-0.62\text{‰} \sim -0.28\text{‰}$),认为其源区为橄榄岩和再循环碳酸盐化榴辉岩的混合源区。Wang Xiaojun 等(2018)发现属于 EM1 端元的经典 Pitcairn 群岛地区玄武岩具有低 $\delta^{26}\text{Mg}$ 值异常,认为其源区来自俯冲含碳酸盐沉积物改造的深部地幔。

值得指出的是,近年来研究发现低 $\delta^{26}\text{Mg}$ 值存在多解性。除再循环碳酸盐外,部分榴辉岩和碳酸盐化榴辉岩也具有低的 $\delta^{26}\text{Mg}$ 值(Wang Shuijiong et al., 2012; Wang Xuance et al., 2015b)。月岩高 Ti 玄武岩的钛铁矿堆晶作用也同样显示出低 $\delta^{26}\text{Mg}$ 值(Sedaghatpoura et al., 2013)。另一方面,成因与俯冲板块有关的岛弧火山岩却有着与地幔相似的轻 Mg 同位素组成(Teng Fangzhen et al., 2016; Li

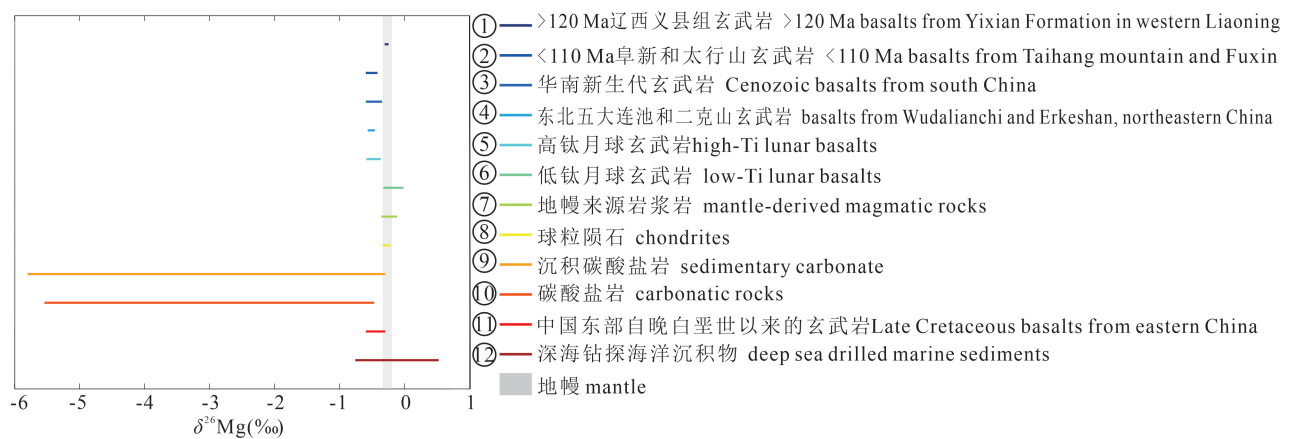


图 2 部分储库镁同位素组成

Fig. 2 Part of the reservoir magnesium isotope composition

①②据 Zhang Hongfu and Zheng Jianping, 2003; Tang Yanjie et al., 2006; Yang Wei and Li Shuguang, 2008; ③据 Huang Jian et al., 2015a; ④据 Tian Hengci et al., 2016; ⑤据 Sedaghatpoura et al., 2013; ⑥据 Huang Jian et al., 2015a; ⑦据 Teng Fangzhen et al., 2010a; ⑧据 Sun and McDonough, 1989; ⑨据 Tian Hengci et al., 2016; ⑩据 Huang Jian et al., 2015a; ⑪据 Galy et al., 2002; Young and Galy, 2004; Tipper et al., 2006a; Pogge et al., 2008; Higgins and Schrag, 2010; Jacobson et al., 2010; Pokrovsky et al., 2011; Wombacher et al., 2011; ⑫据 Teng Fangzhen et al., 2016

①② from Zhang Hongfu and Zheng Jianping, 2003; Tang Yanjie et al., 2006; Yang Wei and Li Shuguang, 2008; ③ from Huang Jian et al., 2015a; ④ from Tian Hengci et al., 2016; ⑤ from Sedaghatpoura et al., 2013; ⑥ from Huang Jian et al., 2015a; ⑦ from Teng Fangzhen et al., 2010a; ⑧ from Sun and McDonough, 1989; ⑨ from Tian Hengci et al., 2016; ⑩ from Huang Jian et al., 2015a; ⑪ from Galy et al., 2002; Young and Galy, 2004; Tipper et al., 2006a; Pogge et al., 2008; Higgins and Schrag, 2010; Jacobson et al., 2010; Pokrovsky et al., 2011; Wombacher et al., 2011; ⑫ from Teng Fangzhen et al., 2016

Shuguang et al., 2017)。

5 低 $\delta^{26}\text{Mg}$ 值的多解性

5.1 岛弧火山岩的低 $\delta^{26}\text{Mg}$ 值

Teng Fangzhen 等 (2016) 和 Li Shuguang 等 (2017) 对环太平洋和 Martinique 岛弧火山岩进行系统的 Mg 同位素组成研究, 结果表明岛弧火山岩具有与地幔相似的 Mg 同位素组成。这原因之一可能是俯冲板片携带的再循环沉积碳酸盐岩在岛弧下发生板片脱水, 主要溶解的是富 Ca 的方解石, 因此产生的流体含很少量的 Mg, 因此, 不会明显的改变岛弧玄武岩的 Mg 同位素组成。另一个原因是来自蚀变洋壳和蚀变深海橄榄岩的流体, 可能具有比地幔重的 Mg 同位素组成。低 $\delta^{26}\text{Mg}$ 玄武岩是由大洋板块俯冲释放碳酸盐所致, 还是同位素轻钛铁矿在其地幔源中积累所致, 一直存在争议。

5.2 碳酸盐化榴辉岩的低 $\delta^{26}\text{Mg}$ 值

碳酸盐化榴辉岩和沉积碳酸盐具有类似的轻 Mg 同位素组成 (Wang Shuijiong et al., 2012, 2014a, 2015a; Huang Kangjun et al., 2016b), 蚀变洋壳经历高温高压变质可以形成榴辉岩, 碳酸盐化榴辉岩和再循环沉积碳酸盐交代产生的熔体表现出低 $\delta^{26}\text{Mg}$ 值, 用 Mg 同位素示踪深部碳循环可能会产生多解 (Wang Shuijiong et al., 2016b; Tian Hengci et al., 2018)。Wang Shuijiong 等 (2014b) 发现苏鲁超高压变质带的荣成 UHP 大理岩和其包裹的榴辉岩块, 独立的榴辉岩的 Mg 同位素组成与原岩一致 ($\delta^{26}\text{Mg} \approx -0.25 \pm 0.07\text{‰}$), 而包裹的碳酸盐化榴辉岩块有较轻的 Mg 同位素组成 ($\delta^{26}\text{Mg} \approx -0.6 \pm 1.9\text{‰}$), 该文指出碳酸盐化榴辉岩的极低 $\delta^{26}\text{Mg}$ 值异常是由榴辉岩变质所致。Wang Shuijiong 等 (2014b) 将不同的 Mg 同位素组成归结于板块俯冲过程中榴辉岩和大理岩之间发生了同位素交换反应, 认为再循环进入深部地幔的低 $\delta^{26}\text{Mg}$ 值物质有两类: 再循环碳酸盐的 $\delta^{26}\text{Mg}$ 值不可能低于 -2.5‰ ; 再循环碳酸盐化榴辉岩 (含碳量不高), 其 $\delta^{26}\text{Mg}$ 值可低至 -1.9‰ 。Huang Jinxian 等 (2016c) 在南非金伯利岩榴辉岩中再次验证了确实存在轻镁同位素组成, 解释为继承了蚀变洋壳的特点, 可能是板块俯冲过程中和沉积碳酸盐岩发生同位素交换所造成的结果。Liu Pingping 等 (2016b) 认为碳酸盐化榴辉岩部分熔融形成的熔体富集 SiO_2 , Na_2O 和 K_2O , 亏损 TiO_2 , 并与地幔橄榄岩反应形成辉石岩; 然而, 玄武

岩的 $\delta^{26}\text{Mg}$ 值并没有与 $(\text{Gd}/\text{Yb})_{\text{N}}$ 和 Fe/Mn 存在相关性, 以此排除混入碳酸盐化榴辉岩, 得出源区并没有混入碳酸盐化榴辉岩的结论, 同时指出 Zn 同位素可以区分再循环碳酸岩和再循环榴辉岩。

5.3 钛铁矿堆晶的低 $\delta^{26}\text{Mg}$ 值

除此之外, 钛铁矿堆晶作用也可造成低 $\delta^{26}\text{Mg}$ 值, Sedaghatpoura 等 (2013) 发现高钛和低钛月岩玄武岩表现出显著的 Mg 同位素分馏 ($\delta^{26}\text{Mg} = -0.61\text{‰} \sim -0.02\text{‰}$), 但是月球不存在板块俯冲作用, 与碳酸盐交代作用无关, 认为钛铁矿堆晶作用可能与高 Ti 月岩玄武岩低 $\delta^{26}\text{Mg}$ 值 ($\delta^{26}\text{Mg} = -0.6\text{‰}$) 有关。另外, 地幔岩浆上升时因温度不同, 扩散引起的同位素动力学分馏程度十分显著 (Richter et al., 2008), 由于热扩散和化学扩散的发生, 造成玄武岩轻 Mg 同位素组成异常 (Liu Sheng'ao and Li Shuguang, 2019), 这一问题一度引来争议。前人提出了排除源区钛铁矿堆晶作用的方法, 认为若存在钛铁矿堆晶, TiO_2 会与 Nb/Ta 和 $\delta^{26}\text{Mg}$ 形成负相关关系 (Huang Jian et al., 2015a, Tian Hengci et al., 2016)。

6 结论与展望

地球深部碳循环对全球气候变化、生命探索和岩石圈演化的研究具有重要意义, Mg 同位素是其有效的示踪工作。笔者等系统阐述了镁同位素示踪深部碳循环的原理, 例举了镁同位素示踪中国东部深部碳循环的实例, 论述了中生代俯冲的古太平洋板块所释放的碳酸盐熔体/流体与地幔相互作用, 是造成中国东部地幔具有普遍的轻 Mg 同位素组成的主要原因。中国东部岩石圈的减薄与破坏体现了中国东部中生代深部碳循环的地球动力学效应。因此, 研究中国东部中生代深部碳循环有助于制约中国东部岩石圈减薄和破坏机制。

然而, 轻镁同位素组成存在多解性。除了再循环碳酸盐外, 岛弧火山岩、榴辉岩、钛铁矿堆晶均具有轻镁同位素组成, 因此, 需要研究探索不同同位素和 (或) 主微量元素特征联合示踪。例如, Liu Sheng'ao 等 (2016a) 研究了中国东部玄武岩 Zn 同位素组成, 发现 $\delta^{66}\text{Zn}$ 与 $\delta^{26}\text{Mg}$ 具有负相关性, 可以采用 Zn—Mg 同位素联合示踪地球深部碳循环。另一方面, 地幔深部碳循环机理和深部碳循环如何贡献地幔不均一性等科学问题仍未解决, 有待进一步开展研究。

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Mg isotope tracing the Mesozoic deep carbon cycle in eastern China

DUAN Xianzhe^{1, 2)}, NIU Sujuan¹⁾, LI Sai^{1, 2)}, LI Nan²⁾, SUN Haoran¹⁾, GUO Cong¹⁾,
XIAO Wenzhou^{1, 2)}, SUI Qinglin^{1, 2)}, FENG Peng^{1, 2)}, HE Haiyang^{1, 2)}

1) School of Resource & Environment and Safety Engineering, University of South China, Hengyang, Hunan, 421001;
2) Hunan Key Laboratory of Rare Metal Minerals Exploitation and Geological Disposal of Wastes, Hengyang, Hunan, 421001

Abstract: The Earth’s deep carbon cycle is of great significance for studying the global climate change, life exploration, and lithospheric evolution. The lithosphere beneath the eastern China is an important place for the deep carbon cycle, and its thinning and destruction are closely related to the deep carbon cycle. The subduction of the Pacific plate in the Mesozoic is the key to the thinning and destruction of the lithosphere in Eastern China, and plays a vital role in the large-scale metal mineralization in North China Craton and South China Block. This paper systematically elucidates the principle of Mg isotope tracing the Earth’s deep carbon cycle, gives an example of the deep carbon cycle in the eastern China, and discusses the interaction of the mantle with the carbonated melts/fluids released from the subducted paleo-Pacific plate during the Mesozoic, which is an important reason accounting for the commonly existed light Mg isotopic composition in the eastern China. Furthermore, the polysolvability of light magnesium isotopes is pointed out and the future trend of studying the Earth’s deep carbon cycle via multi-isotope joint tracing is also proposed.

Keywords: deep carbon cycle; Mesozoic; lithosphere beneath the eastern China; Mg isotope; lithospheric thinning and destruction; paleo-Pacific plate subduction

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First author: DUAN Xianzhe, male, born in 1985, Doctor, Associated Professor, mainly engaged in Petro-/Ore deposit geochemistry; Email: duanxianzhe@126.com

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中国地质学会印发《创新基地评选和管理办法(试行)》通知

为深入贯彻落实国家创新驱动发展战略,提高地质科技创新能力,促进地质科技成果转化和示范应用,服务国家经济社会发展。中国地质学会依托自身平台优势和行业特色,开展创新基地建设工作。根据中国科协《中国科学技术协会事业发展“十四五”规划(2021-2025年)》、《“科创中国”三年行动计划(2021-2023)》和《“科创中国”创新基地建设实施与管理办法(试行)》的要求,中国地质学会制定了《中国地质学会创新基地评选和管理办法(试行)》,并经中国地质学会第40届理事会第三十七次常务理事会议(通讯)审批通过,现予以印发,请遵照执行。

附件:《中国地质学会创新基地评选和管理办法(试行)》

第一章 总 则

第一条 为深入贯彻落实国家创新驱动发展战略,打造各类科技创新平台,提高地质科技创新能力,促进地质科技成果转化和应用示范,助力国家经济社会发展。根据中国科协《中国科学技术协会事业发展“十四五”规划(2021-2025年)》、《“科创中国”三年行动计划(2021-2023)》和《“科创中国”创新基地建设实施与管理办法(试行)》的要求,中国地质学会将依托自身平台优势和行业特色,建设一批科技创新平台——创新基地,并择优推荐申报中国科协“科创中国”创新基地。为规范中国地质学会创新基地(以下简称创新基地)建设工作,特制定本办法。

第二条 创新基地应坚持“面向世界科技前沿、面向经济主战场、面向国家重大需求、面向人民生命健康”的战略方向,围绕地质科技前沿和“卡脖子”关键核心技术,凝聚创新资源,加强国际合作,开展前瞻性研判、技术研发、成果转化及应用示范,破解地学领域的重大科技难题,助力新时代地质工作创新发展,为国家经济社会高质量发展提供重要地质科技支撑力量。

第三条 创新基地建设坚持“需求导向、协同攻关、开放融合、互信共享、共同发展”的原则。聚焦国家、行业和区域经济社会发展重大需求,汇聚一批有创新能力、创新激情、创新活力的优秀地质科技工作者,沉淀一批先导技术、产业需求、技术标准和数据资源,联合企业、高校、科研院所等单位的创新团队开展协同攻关创新活动,提升自主创新、成果转化与应用示范的能力,引领行业技术进步,助力国家经济社会发展。

第二章 建设任务

第四条 创新基地分为产学研协作、创新创业孵化、国际创新合作三种类型。

第五条 各类型创新基地的主要任务包括以下内容:

(一)产学研协作类。聚焦关键国家和行业内核心技术领域,开展前瞻性研判,组织团队集聚攻关,取得原创性成果,兼顾公益性与市场化属性,探索产学研可持续协作机制,促进科技成果转化,推动行业技术进步,服务国家经济社会发展;

(二)创新创业孵化类。融合产业界、学术界、创投界等创新要素,促进各领域优秀人才密切联系、交流碰撞,科学有序孵化优质科创企业,营造区域良好创新创业创造生态,服务科技创新创业力量有序壮大和良性扩散;

(三)国际创新合作类。引入境外亟需紧缺科技人才团队和成熟度高、具有实用价值和市场潜力的技术成果,通过合作研究、国际技术转移等形式,促进全球科创资源与区域重点产业和企业高效衔接、持续落地。

第六条 创新基地主要采取以下方式开展具体工作:

(一)坚持问题导向,广泛征集或企业发展过程中面临的技术需求、问题,张榜求贤,吸引有能力的科研单位或团队,开展重点创新任务、重大科技难题攻关,推动行业的技术进步;

(二)围绕国家、区域产业共性问题和企业急需解决的关键技术问题,推行从需求端到研发端的自下而上研发模式,联合企业开展项目技术合作;

(三)坚持目标导向,梳理产业发展共性难题,围绕产业发展需要,促使政府科研立项与产业发展实际紧密结合,加快成果转化速度,引导相关产业链上下游企业加强合作、共同发展;

(四)围绕产业发展需要,梳理科创项目清单,促使政府科研立项与产业发展实际紧密结合;

(五)坚持国内国际两个视野,通过甄选国内外具有实用价值和市场潜力的最新研究成果,组织开展技术和性能试验,最终导入到地方,促进产业发展;

(下转第 1411 页)