



南天山—北山—索伦—长春缝合带的 性质与演化

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内容提要: 南天山—北山—索伦—长春缝合带作为古亚洲洋的最终闭合位置, 其形成与演化特征一直以来都是中亚造山带相关研究的焦点与热点问题。对于该缝合带形成时代以及俯冲极性等方面的研究, 有利于揭示中亚造山带的增生与演化历史, 为古亚洲洋构造演化模型的建立提供理论支持。笔者等依据南天山—北山—索伦—长春缝合带内的大地构造背景、构造岩石组成、闭合方式和闭合时代的差异, 自西向东将其分为4段: ①南天山缝合带位于缝合带西段, 形成于塔里木板块向北俯冲与哈萨克斯坦—伊犁地块发生拼贴的过程中, 根据高压变质年龄、钉合岩体以及不整合盖层等证据来综合分析, 其闭合时代应为晚石炭世; ②北山缝合带位于缝合带中段, 形成于敦煌地块和阿拉善地块向北俯冲与北部图瓦—蒙古板块发生拼贴的过程中, 根据带内蛇绿岩的年代学证据限定其闭合时代应为早—中二叠世。阿拉善地块北缘的两条蛇绿岩带作为北山缝合带与索伦—长春缝合带之间的连接带, 分别代表了古亚洲洋在该区域闭合时形成的缝合带和弧后盆地, 其形成时代应当为中二叠世—晚二叠世早期; ③索伦—长春缝合带位于缝合带中—东段, 古亚洲洋在该地区同时发生了南北两侧的双向俯冲, 两侧地块在中二叠世—早三叠世完成拼贴; ④长春—延吉缝合带形成于中三叠世前后华北板块与佳木斯—兴凯地块的俯冲增生过程中, 其较西侧索伦—长春缝合带的形成时间(270~250 Ma)晚20~30 Ma。因此长春—延吉缝合带与索伦—长春缝合带的形成时代与构造背景存在显著的差异, 不属于其东延部分。在上述分析基础上, 笔者等认为古亚洲洋沿南天山—北山—索伦—长春缝合带自西向东发生了4个阶段的演化过程, 闭合时代自西向东逐渐变年轻, 整个过程从晚石炭世一直持续到了三叠世, 其中长春—延吉缝合带记录了古亚洲洋和古太平洋构造域叠加与转换的地质过程。

关键词: 中亚造山带; 南天山—北山—索伦—长春缝合带; 分段演化; 俯冲极性; 构造域转换

中亚造山带是全球规模最大的增生型造山带(Windley et al., 2007), 它西起里海, 东临西太平洋北部, 位于西伯利亚克拉通、塔里木克拉通和华北克拉通之间, 跨度达5500 km, 最宽处超过了1100 km(Windley et al., 1990; Şengör et al., 1991; Xiao Wenjiao et al., 2003; Kröner et al., 2007; Jian Ping et al., 2013; 杨高学等, 2021)。中亚造山带的演化是一个漫长而又复杂的过程, 向前可追溯至~1 Ga(Khain et al., 2002), 并一直持续到~250 Ma(Windley et al., 2007; Xiao Wenjiao et al., 2018)。甚至有学者认为直至白垩纪才沿狼山—阴山一线最终闭合(吕洪波等, 2018)。近些年来, 许多学者对中亚造山带进行了详尽的研究, 在构造细分、沉积学、年代学、地球化学、岩浆作用、变质作用、构造变形作

用等方面取得了显著的进展(Khain et al., 2003; Wilde, 2003; Demoux et al., 2009; Zhou Jianbo et al., 2010a; Wu Fuyuan et al., 2011a; Kröner et al., 2013; Xiao Wenjiao et al., 2013; Degtyarev et al., 2015; Buriánek et al., 2017; Chen Long et al., 2021)。然而由于缺乏整体性的讨论, 学者们对于古亚洲的闭合特征仍然有诸多争议。南天山—北山—索伦—长春缝合带作为中亚造山带最南端的缝合带, 它对于限定古亚洲洋的形成和演化特征来说有着非常重要的意义。笔者等依据大地构造背景以及闭合时代的不同, 将该缝合带自西向东分为了4段, 并依次对其闭合特征进行讨论, 从而为揭示中亚造山带的增生与演化历史、构建古亚洲洋的构造演化模型提供理论支持。

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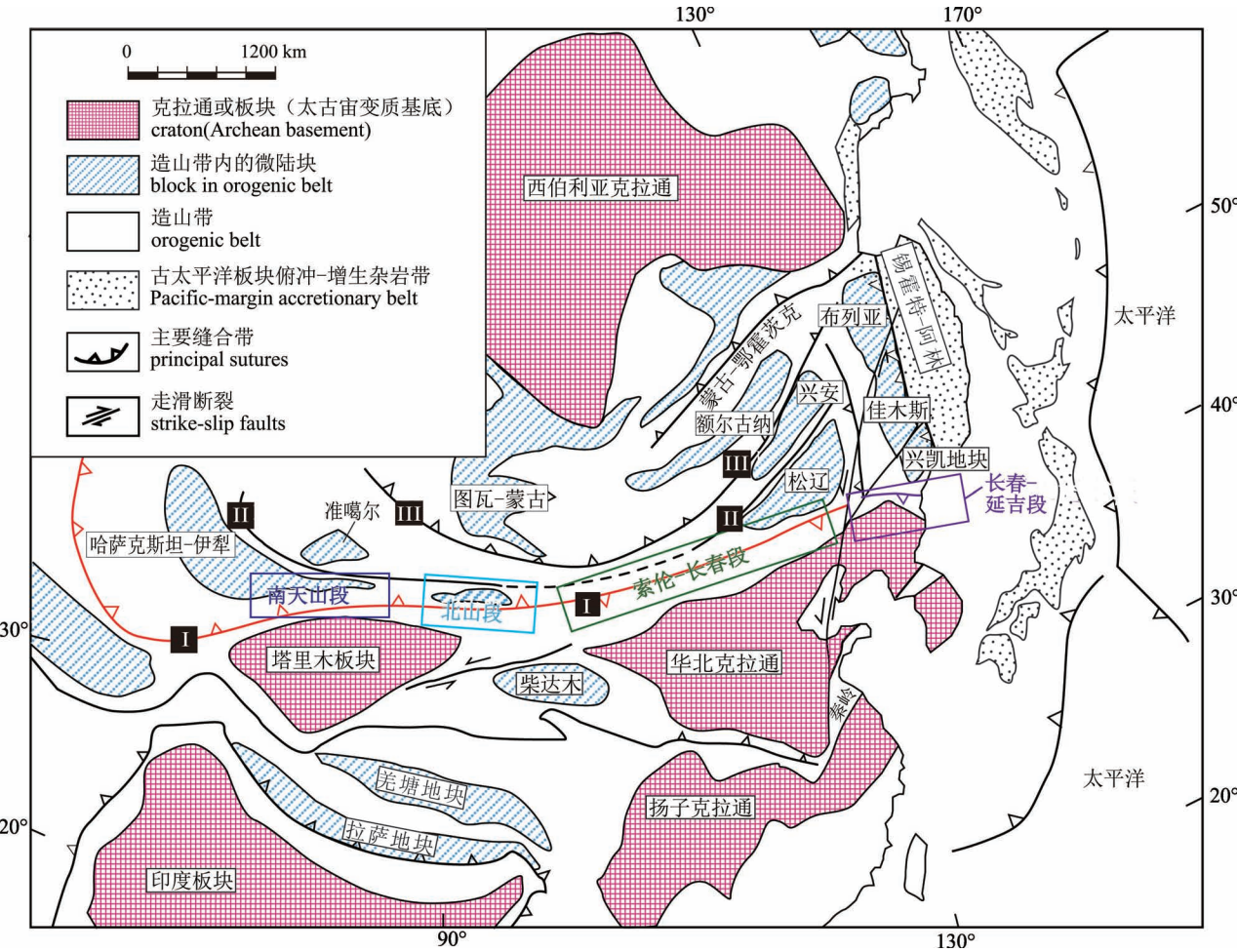


图 1 中东亚主要构造分区示意图(据 Zhou Jianbo et al. , 2018 修改)

Fig. 1 Schematic tectonic map showing the main subdivisions of the central—eastern Asia (modified from Zhou Jianbo et al. , 2018)

主要缝合带:I—南天山—北山—索伦—长春缝合带 (SBSC 缝合带); II—北天山—索伦—贺根山—黑河缝合带 (NSHH 缝合带); III—额尔齐斯—南蒙古—头道桥—新林缝合带 (ESTX 缝合带)

Suture zones: I—South Tianshan Mts. —Beishan Mts. —Solonker—Changchun suture (SBSC suture); II—North Tianshan Mts. —Solonker—Hegen Mts. —Heihe suture (NSHH suture); III—Erqis—South Mongolian—Toudaoqiao—Xinlin suture (ESTX suture)

1 区域构造单元

中亚造山带同世界上其他著名造山带一样,不同学者给予了不同的划分方案。笔者等为了便于论述,在结合前人划分结果的基础上 (Zhou Jianbo et al. , 2018),自北向南将其划分为四个主要的拼合块体 (图 1):图瓦—蒙古和额尔古纳块体;准噶尔和兴安块体;哈萨克斯坦—伊犁、北山和松辽块体;以及环太平洋的佳木斯—兴凯 (布列亚) 块体,他们之间存在着 3 条大型缝合带,从北到南依次为:额尔齐斯—南蒙古—头道桥—新林缝合带 (ESTX 缝合带)、北天山—索伦—贺根山—黑河缝合带 (NSHH 缝合带) 和南天山—北山—索伦—长春缝合带

(SBSC 缝合带)。南天山—北山—索伦—长春缝合带位于中亚造山带最南端,是古亚洲洋最终闭合的位置所在。

依据大地构造背景以及闭合时代的差异,又可以将南天山—北山—索伦—长春缝合带划分为 4 段 (图 1),自西向东依次为:南天山缝合带,北山缝合带,索伦—长春缝合带和长春—延吉缝合带。

2 南天山缝合带

南天山缝合带位于中亚造山带西南缘,呈近东西向延伸,为哈萨克斯坦—伊犁板块与塔里木板块之间长期俯冲—拼贴的产物,是一条由前寒武纪微陆块、古生代岛弧、高压变质岩石和增生楔等共同构

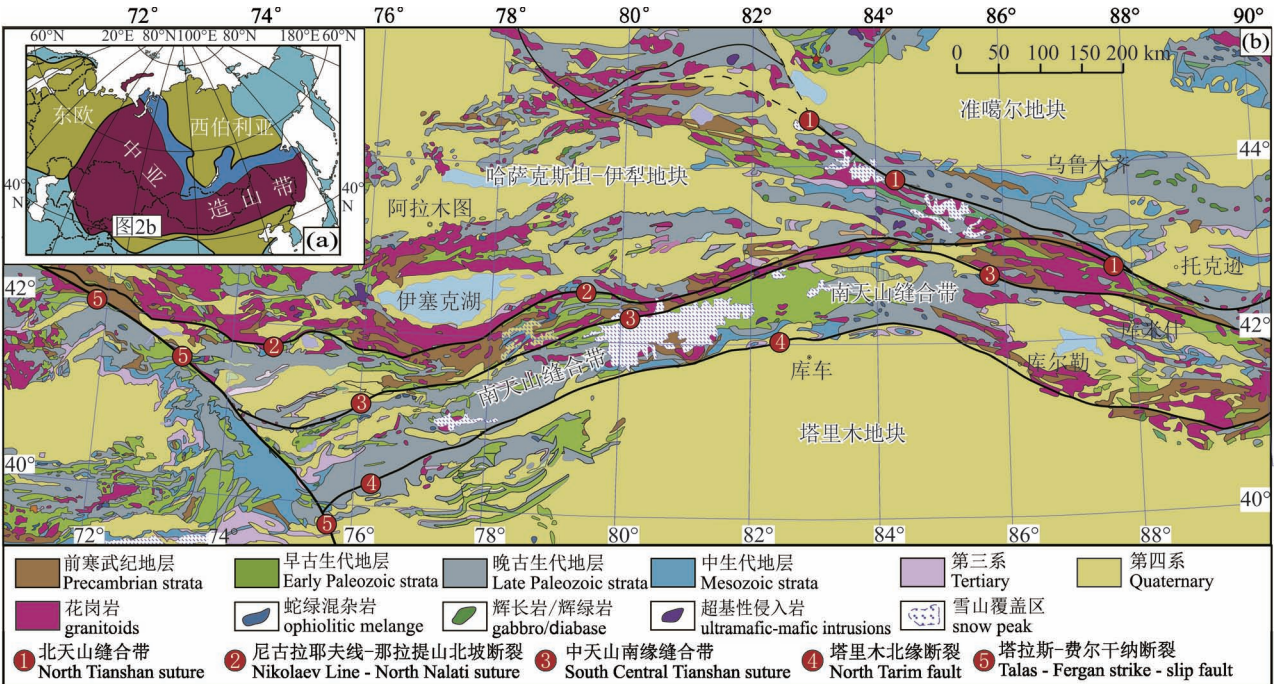


图 2 中亚造山带构造示意图,标注了南天山缝合带的位置(a, 据 Yakubchuk, 2004; Gao Jun et al., 2011 修改)和南天山缝合带及邻区地质图(b, 据高俊等,2009; Gao Jun et al., 2011 修改)

Fig. 2 Tectonic sketch map of the Central Asian Orogenic Belt and the location of South Tianshan Suture (a, modified from Yakubchuk, 2004; Gao Jun et al., 2011) and the geological map of the South Tianshan Suture (b, modified from Gao Jun et al., 2009&, 2011)

成的构造杂岩带(高俊等, 2009; Gao Jun et al., 2011; Zhou Jianbo et al., 2018)。南天山缝合带是古亚洲的分支—南天山洋俯冲消亡的产物,以保存大量的榴辉岩和蓝片岩等高压变质岩为特征,它属于南天山—北山—索伦—长春缝合带的西段,代表了古亚洲洋在该区域的最终闭合位置(Gao Jun et al., 2003; 艾永亮等,2005; John et al., 2008; 左国朝等,2008)。

目前学者们对于南天山洋的最终闭合时间存在有争议。部分学者认为其最终闭合于晚泥盆世—早石炭世(Xia Linqi et al., 2008; Xu Xueyi et al., 2013; 白建科等,2015),然而缝合带内 430~330 Ma 蛇绿岩的发现(徐学义等,2006; Jiang Tuo et al., 2014; Li Chao et al., 2015)表明南天山洋在晚志留世—早石炭世期间还未完全关闭。有部分学者认为南天山洋的演化一直持续到了早三叠世,他们对南天山北缘高压变质带内的榴辉岩中的锆石进行了 SHRIMP 定年,获得了 230~220 Ma 的锆石边部年龄值(张立飞等,2005; Zhang Lifei et al., 2007),以及李曰俊等(2002)在南天山西侧发现了被认为来自于晚二叠世的放射虫化石。然而也有学者认为较年

轻的锆石的出现可能是由于锆石在后期流体作用下继续生长所造成的(Rasmussen, 2005; Jong et al., 2008),不能代表碰撞造山事件的真正年龄(高俊等,2009)。张立飞等(2013)也随后将南天山榴辉岩的峰期变质年龄修正为 320.0 ± 3.7 Ma,与区域内广泛报道的 310~320 Ma 的榴辉岩变质年龄达成了共识(Su Wen et al., 2010; Klemd et al., 2011; Li Qiuli et al., 2011)。舒良树等(2007)认为晚二叠世放射虫化石由于保存较差,无法准确进行种类鉴定,同样无法作为南天山洋闭合时代判定的依据。

还有部分学者认为南天山洋闭合时代为晚石炭世。上二叠统陆相磨拉石的相关报道(李向东等, 2000; 卢华复等,2001)以及古地磁证据(Bazhenov et al., 2003; Wang Bo et al., 2007)证明塔里木板块和哈萨克斯坦—伊犁地块在早二叠世就已经拼合在了一起,自二叠纪开始南天山的区域构造背景由汇聚碰撞变为区域性扩张(李锦轶等,2006),南天山洋的闭合时代应当不晚于二叠世。南天山周围地区还分布着广泛的二叠纪陆相磨拉石沉积记录(王作勋等,1990; 王宝瑜等,1994; Gao Jun et al., 1998),表明南天山地区在二叠纪期间已经由海相

转变为了陆相。分布于南天山地区和塔里木北缘的石炭纪浅海碳酸盐地层沉积被二叠纪硅质碎屑—火山沉积所覆盖 (Han Yigui et al., 2016), 记录了晚石炭世至二叠纪南天山洋残余洋盆向湖泊河流相过渡的过程。Gao Jun 等 (2011) 报道了在南天山中部发现的年龄为 285 Ma 的花岗岩岩脉, 其切割了南天山洋在俯冲过程中所形成的高压—低温变质围岩, 地球化学数据显示其形成于碰撞后地壳折返过程中, 表明南天山洋的闭合应当发生在二叠纪之前。南天山的阿吾子地区在峰期榴辉岩相变质之后发生了广泛的异剥钙榴岩化 (Li Xuping et al., 2007), 后者的退变质年龄为 291 Ma (高俊等, 2006), 限定了南天山缝合带形成于早二叠世之前。在南天山高压变质岩中获得的 335~310 Ma 的白云母 Ar—Ar 坪年龄 (高俊等, 2006) 和 313~302 Ma 的白云母全岩等时线年龄 (Klemd et al., 2005), 代表了高压变质岩折返到浅构造层次 (蓝片岩相) 的时间。笔者等支持南天山洋在晚石炭世闭合的观点, 随后南天山地区进入到了后造山演化阶段。

对于南天山洋的俯冲极性存在有两种观点: 一些学者根据区域性构造特征如不整合等主张哈萨克斯坦—伊犁板块发生了南向俯冲 (Charvet et al., 2011; Wang Bo et al., 2011; Lin Wei et al., 2013; Alexeiev et al., 2015); 另外一些学者则主张塔里木板块发生了北向俯冲 (Windley et al., 1990; Gao Jun et al., 2008; Han Baofu et al., 2011; Xiao Wenjiao et al., 2013; Han Yigui et al., 2016)。区域地质特征显示, 在泥盆世—晚石炭世期间, 塔里木板块北缘主要为被动大陆边缘环境, 以沉积岩为主, 罕见弧岩浆作用 (Xiao Wenjiao et al., 2013); 而同时期北侧的伊犁—中天山地块内发育有广泛的弧岩浆作用 (Han Yigui et al., 2016) 和大面积的增生杂岩 (高俊等, 2009), 并随后发生了 LP—HT 麻粒岩相变质作用 (Lü Zeng et al., 2012; Zhang Lifei et al., 2013), 这是塔里木板块北向俯冲的产物。因此我们认为在南天山洋闭合过程中, 塔里木板块发生了北向俯冲。

3 北山缝合带

北山缝合带属于南天山—北山—索伦—长春缝合带中段, 东侧以巴丹吉林沙漠为界与兴蒙造山带相接, 西侧以星星峡断裂为界与中天山相接, 北侧为图瓦—蒙古板块, 南侧为敦煌地块, 东南部为阿拉善地块 (图 3; 刘雪亚等, 1995; Song Dongfang et al.,

2014; 牛亚卓等, 2018; 辛后田等, 2020)。北山缝合带最初被认为是中天山的东延部分 (李春昱, 1980; 刘雪亚等, 1995), 然而现在许多学者认为北山是一个独立的地质单元, 是不同类型和不用时期岩石混杂在一起而形成的杂岩体 (Xiao Wenjiao et al., 2010a; Song Dongfang et al., 2013a; Liu Qian et al., 2015)。北山缝合带从寒武纪开始便开始了漫长的俯冲增生造山过程 (Xiao Wenjiao et al., 2010a; Song Dongfang et al., 2015), 具有多旋回复合造山带的特点 (龚全胜等, 2002; 聂凤军, 2002)。

在北山缝合带内分布着四条蛇绿混杂岩带, 自北向南依次为: 红石山—百合山蛇绿混杂岩带、芨芨台子—石板井—小黄山蛇绿混杂岩带、红柳河—牛圈子—洗肠井蛇绿混杂岩带、辉铜山 (柳园)—帐房山蛇绿混杂岩带, 它们的形成时代差异较大, 引起了学者们对于北山缝合带演化特征的广泛讨论。左国朝等 (1990) 认为芨芨台子—石板井—小黄山一线分布的早古生代蛇绿混杂岩代表了早古生代缝合带的位置, 北山地区古洋盆在志留纪末至早泥盆世就已经闭合了。但是该构造带中辉长岩的最新同位素年龄研究显示该蛇绿岩形成年代为早石炭世 (Wu Tairan et al., 2011b; 李向民等, 2012), 且布格重力异常 (孟令顺等, 1995; 李治等, 2019) 和磁测数据 (余钦范等, 1995) 显示其可能是一条深大断裂而不是缝合带, 因此该带可能与晚古生代后碰撞运动时期的陆内伸展作用有关 (Xia Linqi et al., 2003, 2004, 2005)。红柳河—牛圈子—洗肠井蛇绿混杂岩带内辉长岩年龄为 536~520 Ma (Cleven et al., 2015; Shi Yuruo et al., 2018), 杨合群等 (2008, 2010) 和孙立新等 (2017) 认为该带代表了早古生代北山洋最终的闭合位置, 同时也认为较为年轻的红石山—百合山、柳园—帐房山混杂岩带并不代表缝合带的位置, 它们可能形成处于初始小洋盆形成之前“红海型”海槽环境下 (杨合群等, 2010)。然而从北山周围地区的大地构造演化特征来看, 北山西部的南天山缝合带和东部的索伦—长春缝合带均被认为是晚古生代—早中生代的板块拼贴的产物 (Gao Jun et al., 2003; Xiao Wenjiao et al., 2003, 2008, 2010b; Zhang Lifei et al., 2005), 北山洋在早古生代完全闭合的观点则难以解释这一奇怪的现象, 因此现在越来越多的学者认为早古生代并不是北山地区古大洋的最终拼合年代, 北山缝合带是多期次碰撞拼合的产物 (Xiao Wenjiao et al., 2010a; Song Dongfang et al., 2013a, b; Guo Qianqian et al.,

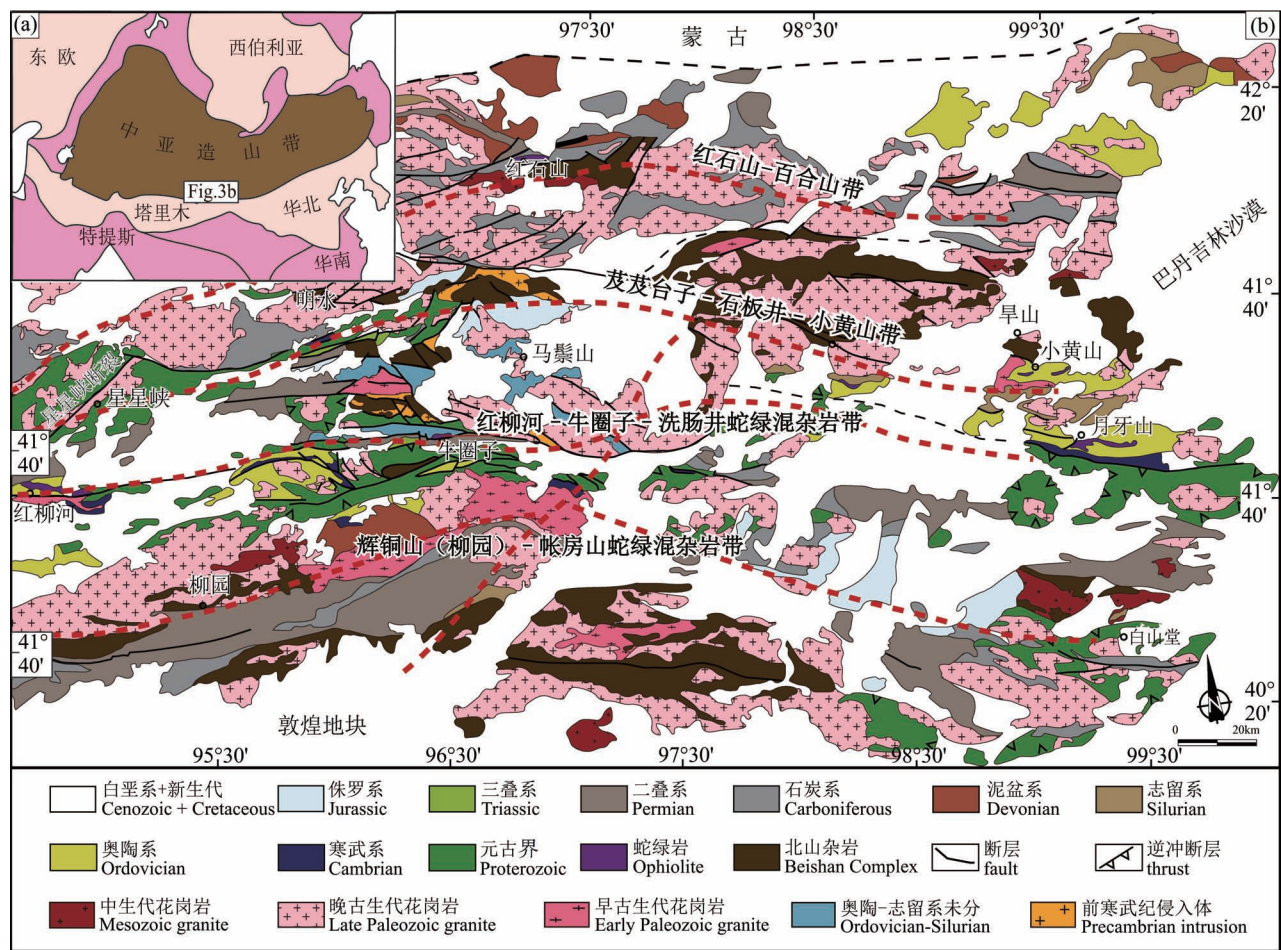


图 3 中亚造山带构造示意图,标注了北山缝合带的位置(a, 据王国强等,2021 修改)和北山缝合带区域地质图(b,据 Song Dongfang et al. , 2014 修改)

Fig. 3 Tectonic sketch map of the Central Asian Orogenic Belt and the location of the Beishan Suture (a, modified from Wang Guoqiang et al. , 2021&) and the geological map of the Beishan Suture (b, modified from Song Dongfang et al. , 2014)

2014; Tian Zhonghua et al. , 2014)。我们认为较晚形成的红石山—百合山蛇绿杂岩带和辉铜山(柳园)—帐房山蛇绿杂岩带可能代表着古亚洲洋在北山地区的最终闭合位置。

红石山—百合山蛇绿岩带位于北山北侧,黄增保等(2006a)对红石山—百合山带内的变质基性岩进行了测定,其地球化学特征显示其形成于洋中脊环境,来源于亏损的软流圈地幔,且具有密集的重力梯度带和局部磁力高异常,符合分割两个大陆缝合带的地球物理特征(何世平等,2002),地理上该带南侧为华南生物地理区,北侧为安加拉植物区(王元龙等,2001),符合板块缝合带的地理学特征。红石山蛇绿岩带普遍被认为形成于晚石炭世—二叠纪(龚全胜等,2002; 魏志军等,2004; 黄增保等,2006a, b),红石山—百合山地区蛇绿岩混杂岩中的

斜长花岗岩年龄为 297.3 ± 1.5 Ma(牛文超等,2019),橄榄辉长岩的锆石加权平均年龄为 281.8 ± 2.6 Ma(Ao Songjian et al. , 2010),附近坡北地区蛇绿杂岩锆石年龄为 $280 \sim 270$ Ma(姜常义等,2006; 李华芹等,2006),可见红石山地区较为年轻的蛇绿杂岩主要集中在早二叠世。根据造山带中最年轻的蛇绿混杂岩代表了洋盆最终的闭合位置与闭合时间理论(李继亮,2004; 肖文交等,2019),红石山洋的闭合时间应当为早—中二叠世。

辉铜山(柳园)—帐房山蛇绿岩带位于北山南侧,其中柳园地区辉长岩锆石 U-Pb 加权平均年龄为 286 ± 2 Ma (Mao Qigui et al. , 2011), 272.7 ± 4.4 Ma 和 291.4 ± 4.9 Ma(Zhang Yuanyuan et al. , 2011),上覆流纹岩的年龄为 277 ± 4 Ma(Wang et al. , 2016b),柳园枕状玄武岩内发现有煌斑岩侵入

体,其侵入时间为 240~220 Ma,代表了岩浆岩源区由亏损地幔向富集地幔的转变(刘畅等,2006),柳园地区所有火山岩均被上二叠统一三叠统砂岩不整合覆盖(Wang Yu et al., 2016b),因此该蛇绿岩带同样也被认为形成于早—中二叠世(Xiao Wenjiao et al., 2010a, 2015; Mao Qigui et al., 2011)。北山地区在晚石炭世同时存在有红石山洋和柳园洋两个古洋盆(Xiao Wenjiao et al., 2010a),古亚洲洋北山段的完全闭合发生在早—中二叠世。

阿拉善地块位于北山缝合带东南部,其通常被认为是华北板块的西段。在阿拉善地块北缘发育有两条主要的蛇绿混杂岩带:恩格尔乌苏混杂岩带与查干楚鲁混杂岩带,他们分别沿恩格尔乌苏断裂带和巴丹吉林断裂带呈北东向延伸,是连接北山缝合带和索伦—长春缝合带的过渡带(Xiao Wenjiao et al., 2018),前者被认为代表了古亚洲洋在该地区主要缝合线的位置,后者则代表着弧后盆地(Zheng Rongguo et al., 2014),二者对于约束古亚洲洋的闭合特征有着重要的意义。恩格尔乌苏蛇绿混杂岩带主要由碳酸盐岩基质以及辉长岩、玄武岩等镁铁—超镁铁质岩构造块体所组成,从玄武岩中获得的锆石 U-Pb 年龄为 302 ± 14 Ma。查干楚鲁蛇绿混杂岩带主要由微碎屑岩基质以及辉长岩、超镁铁质岩构造块体所组成,从辉长岩样本中获得的锆石 U-Pb 年龄为 275 ± 3 Ma (Zheng Rongguo et al., 2014)。结合从恩格尔乌苏蛇绿混杂岩带中发现的中二叠世晚期放射虫化石证据(谢力等,2014),以及区域内指示碰撞后伸展构造环境的中—晚二叠世钙碱性花岗岩证据(Chen Yan et al., 2021; Hui Jie et al., 2021),我们认为古亚洲洋在阿拉善地块北缘的俯冲闭合活动可能持续到了中二叠世—晚二叠世早期。

据前人研究结果显示,古亚洲洋北山段的闭合是南侧敦煌地块与阿拉善地块发生了向北的俯冲所导致的。在北山地区自南向北依次发育有牛圈子—洗肠井蛇绿岩带、勒巴泉变质杂岩和公婆泉群火山碎屑岩,它们代表着敦煌地块向北俯冲过程中形成的洋壳残片、弧前增生杂岩和岩浆弧(Song Dongfang et al., 2014, 2015)。宋东方等(2018)观察到北山地区的韧性变形事件自北向南逐渐变年轻,这是地体在北向俯冲过程中,其边缘不断发生增生作用所导致的。此外,北山地区的地球物理剖面显示该地区的大部分逆冲断层均为向南推覆的趋势(Xiao Wenjiao et al., 2010a),同样表明在北山洋的闭合

过程中,南侧陆块发生了向北的俯冲。

4 索伦—长春缝合带

索伦—长春缝合带位于南天山—北山—索伦—长春缝合带的中—东段,它被三个构造单元所包围,北侧为西伯利亚板块,南侧为华北板块,东侧为东北陆块群(图 4; Natalin, 1991; Şengör et al., 1993; Jahn et al., 2000; Li Jinyi, 2006),它西侧与北山缝合带相接,东部可一直延伸至东北陆块群之内,北侧为一条长达数千千米的超岩石圈深大断裂带:西拉木伦河断裂带,南侧通过一条增生杂岩带与华北板块相接(李益龙等, 2009; Xiao Wenjiao et al., 2003)。该地区邻近古亚洲洋构造域和古太平洋构造域的交汇部位,对于两大构造域的启动、终止和转换等方面的研究来说都极为重要。

对于该地区缝合带的具体位置,目前人们存在有争议。部分学者认为贺根山—黑河断裂带代表了该地区的最终缝合带,古亚洲洋沿该断裂带最终发生闭合(Zhang Yunping et al., 1989; Mueller et al., 1991; Zorin, 1999; Kravchinsky et al., 2002)。然而东北地区发生了广泛的 500 Ma 左右的泛非期高级变质事件,而类似的变质事件在华北板块和华南板块中没有出现,中国东北地区的构造属性不同于华北和华南板块的构造属性。贺根山—黑河断裂带作为松辽地块和兴安地块之间的断裂带,应当只是东北陆块群内部的断裂带,而不具有板块缝合带的属性(周建波等, 2012; Zhou Jianbo et al., 2013)。在贺根山蛇绿岩带南侧发现的海相地层(尚庆华, 2004)和在西拉木伦河北岸发现的远洋古生物化石(王玉净等, 1997)也能证明真正缝合带的位置应当在更南的索伦—长春一带。板块缝合带还应当具有生物种群分割线的特征,这一点在索伦缝合带两侧的生物分布特征中得到了体现(孙跃武等, 2018)。因此大部分学者认为索伦—长春缝合带才是华北板块和西伯利亚板块间的缝合带(Wang Quan et al., 1986; Xiao Wenjiao et al., 2003; Wu Fuyuan et al., 2007; Jian Ping et al., 2008; Miao Laicheng et al., 2008; Chen Bin et al., 2009; Zhang Xiaohui et al., 2009)。

对于索伦—长春缝合带的闭合时间同样存在争议,具体包括泥盆—早石炭世(Tang Kedong, 1990; 邵济安, 1991; Hong Dawei et al., 1995; 徐备等, 1997)、二叠—三叠纪(Şengör et al., 1993; Robinson et al., 1999; Badarch et al., 2002; Xiao Wenjiao et

叠世古亚洲洋东段仍然有安第斯型岩浆活动的痕迹,同时区域内缺乏三叠纪的沉积记录(内蒙古地质矿产局, 1991),这可能与三叠纪区域内的碰撞隆升活动有关。此外,种种证据表明在索伦—长春缝合带形成过程中,西段的形成时间要略早于东段;源区为增厚地壳的高 Sr/Y 碰撞花岗岩在西段的年龄(264~256 Ma, 柳长峰等, 2011)要早于东段(251~245 Ma, Li Shan et al., 2017);西段岩浆活动的年龄峰值要早于东段(吴迪迪等, 2021);沉积记录显示深海相沉积记录在东段地区持续的时间比西段久(赵英利等, 2016);古地磁数据亦显示在二叠纪晚期松辽地体和华北板块之间的纬度差在逐渐缩小(Ren Qiang et al., 2020)。因此,我们认为索伦—长春缝合带在中二叠世—早三叠世发生了闭合。

关于古亚洲洋东段的俯冲极性,目前主要有两种争议:一部分学者认为其发生了向南的俯冲,在缝合带北侧发育有双峰式火山作用(Li Jinyi, 2006; Jian Ping et al., 2008, 2010);另外一部分学者认为沿索伦—长春缝合带发生了双向俯冲,缝合带两侧的南、北增生带最终于晚古生代发生拼合(Xiao Wenjiao et al., 2003, 2015; Eizenhoefer et al., 2014; Eizenhofer and Zhao Guochun, 2018; Song Dongfang et al., 2021)。经过综合分析,我们支持第二种观点。首先,在研究区内没有明显的逆冲推覆和地壳增厚现象,也没有高压麻粒岩相变质作用的相关报道(Song Shuguang et al., 2016),该特征更加符合“软碰撞”模式,在该模式中不存在有被动大陆边缘,大洋岩石圈两侧均发育俯冲带(Draut and Clift, 2001, 2013)。其次,在缝合带两侧发育有大量的蛇绿混杂岩,记录了该地区的双向俯冲作用过程,但是,北侧蛇绿混杂岩的年龄多集中在石炭世(Miao Laicheng et al., 2007, 2008; Chen Bin et al., 2009; Zhang Shuanhong et al., 2014; Zhou Jianbo et al., 2015; Pei Shenghui et al., 2016; Song Shuguang et al., 2016),而南侧蛇绿混杂岩的年龄多集中在中二叠世—早三叠世(Miao Laicheng et al., 2007, 2008; Jian Ping et al., 2010; Song Shuguang et al., 2016),二者之间相差大约 50 Ma,推测这可能是由于南向俯冲结束时间稍晚于北向俯冲所致。

5 长春—延吉缝合带

长春—延吉缝合带也被称之为长春—延吉增生杂岩带,它呈北西—南东向延伸,是佳木斯—兴凯地块与华北板块之间的一条典型的增生杂岩带(图 5;

Zhou Jianbo et al., 2013)。它处于中亚造山带的最东缘,既受到古亚洲洋构造域的影响(Şengör, 1996; Wu Fuyuan et al., 2002; Windley et al., 2007; Zhou Jianbo et al., 2013; Kröner et al., 2014),又受到了古太平洋构造域的叠加影响(Faure et al., 1992; Maruyama, 1997; 任纪舜等, 1999; Wilde et al., 2003; Zhou Jianbo et al., 2009, 2010c),是研究古亚洲洋和古太平洋构造域叠加转换过程的关键区域。

呼兰群杂岩、色洛河群杂岩、青龙村群杂岩和开山屯杂岩是沿长春—延吉缝合带自西向东出露的典型增生杂岩(周建波等, 2013, 2020; Zhou Jianbo et al., 2017),它们由不同时代和不同性质的岩石混杂而成(Wu Fuyuan et al., 2007, 2011a),以发育高压变质矿物组合为特征,能为我们提供板块俯冲活动的关键信息(周建波等, 2020)。呼兰群杂岩中的碎屑沉积物的年龄为 287~237 Ma(Wu Fuyuan et al., 2007; 张春艳等, 2009; Lu Linsu et al., 2011; 周建波等, 2013; Cao Jialin et al., 2019),其中最年轻的是从烟筒山红帘石片岩中获得的 237 ± 5 Ma 的原岩年龄(Cao Jialin et al., 2019),而该地区变质岩的变质年龄区间为 229~188 Ma(郗爱华等, 2003; 刘金玉等, 2010; 刘志宏等, 2016),以上数据表明呼兰群杂岩主要受二叠纪—早三叠世构造活动的影响。色洛河群杂岩中碎屑沉积物的年龄为 265~255 Ma(Zhang Yanbin et al., 2008; Zhou Jianbo et al., 2013),与板块俯冲有关的高镁安山岩的年龄为 252 ± 5 Ma(李承东等, 2007),表明色洛河群杂岩应当形成于晚二叠世—早三叠世。青龙村群变质岩的原岩年龄为 250~248 Ma(周建波等, 2013),开山屯杂岩中碎屑锆石的年龄为 292~234 Ma(周建波等, 2013),表明它们均形成于早三叠世。此外,在烟筒山红帘石片岩中发现的未变质长英质侵入体的加权平均年龄为 217 ± 5 Ma(Cao Jialin et al., 2019),限定了俯冲活动年龄的上限。增生杂岩带上覆大酱缸组地层中发育有指示陆相沉积的磨拉石建造(辛玉莲等, 2011),从中获得的最年轻锆石加权平均年龄为 225 ± 1 Ma(图 6b; Wang Bin et al., 2016a),表明长春—延吉地区的古洋盆在晚三叠世就已经关闭了,结合大酱缸组中发现的晚三叠世陆相植物化石(辛玉莲等, 2011),进而我们认为长春—延吉缝合带的形成时代应当在中三叠纪前后(220~240 Ma)。

长春—延吉缝合带传统上被认为是索伦—长春

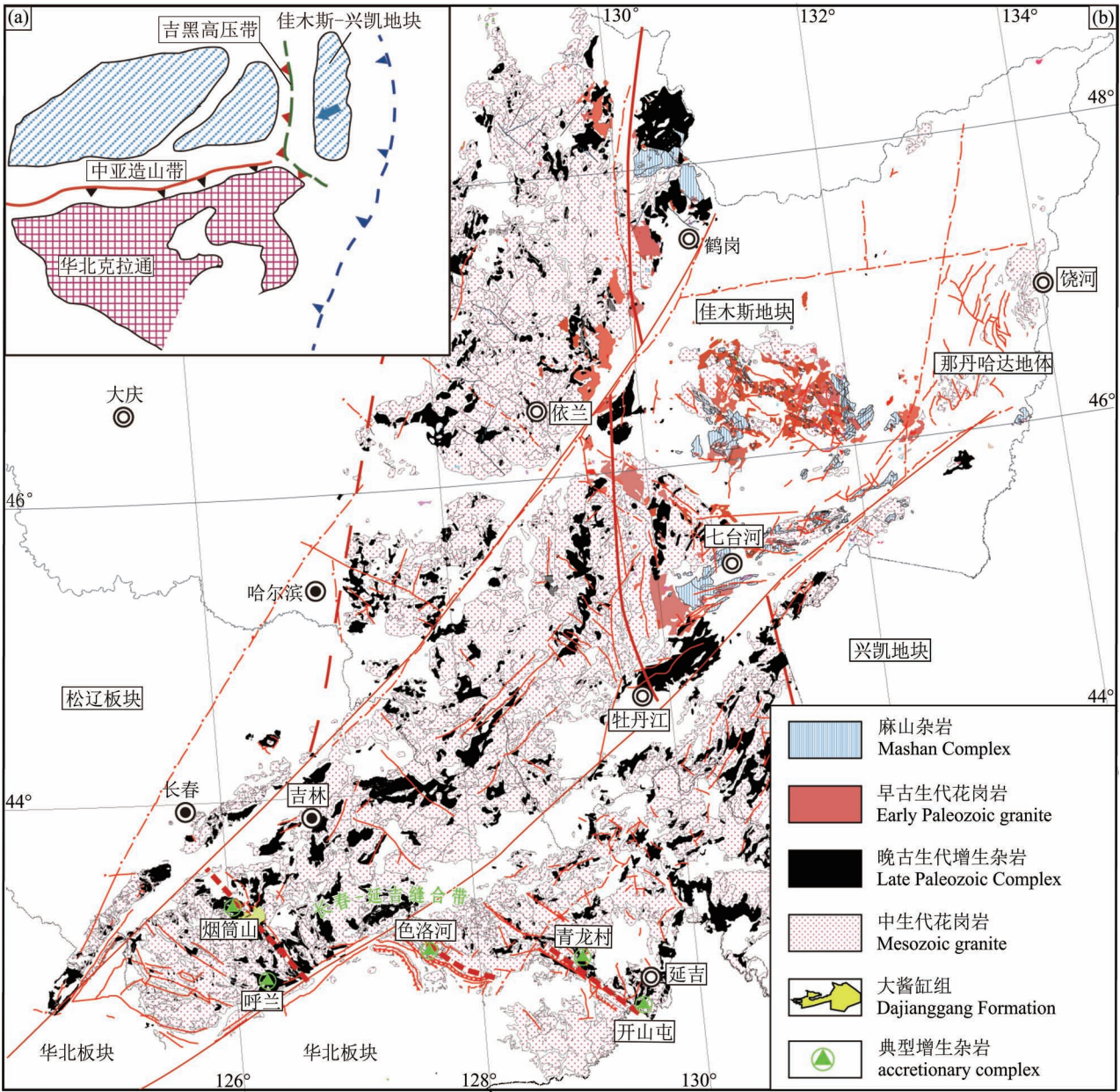


图 5 长春—延吉缝合带构造地质图,标注了主要增生杂岩的名称与位置(据周建波等,2013 修改)

Fig. 5 The regional geological map of the Changchun—Yanji Suture and showing the location of major accretionary complex (modified from Zhou Jianbo et al. , 2013&)

缝合带的东延部分(吉林省地质矿产局, 1988; 唐克东等,1995; 王玉净等,1997; Li Jinyi, 2006; Wu Fuyuan et al. , 2007),属于古亚洲洋闭合的产物。然而最新的研究观点认为,长春—延吉缝合带是佳木斯—兴凯地块受到古太平洋构造域影响,与华北板块发生拼贴的产物,其形成时代较西侧的索伦—长春缝合带的年龄(270~250 Ma)要晚 20~30 Ma (Zhou Jianbo et al. , 2009, 2010b, 2014),因此长春—延吉缝合带记录了古亚洲洋和古太平洋构造域

叠加与转换的地质过程(图 7a,b)。

6 结论

南天山—北山—索伦—长春缝合带位于塔里木—中朝地块与西伯利亚地块之间,为两大板块之间古亚洲洋的最终闭合位置所在,依据缝合带两侧的大地构造背景、构造岩石组成、闭合方式和闭合时代的差异,自西向东将该缝合带分为四段:南天山缝合带位于缝合带西段,形成于塔里木板块向北俯冲

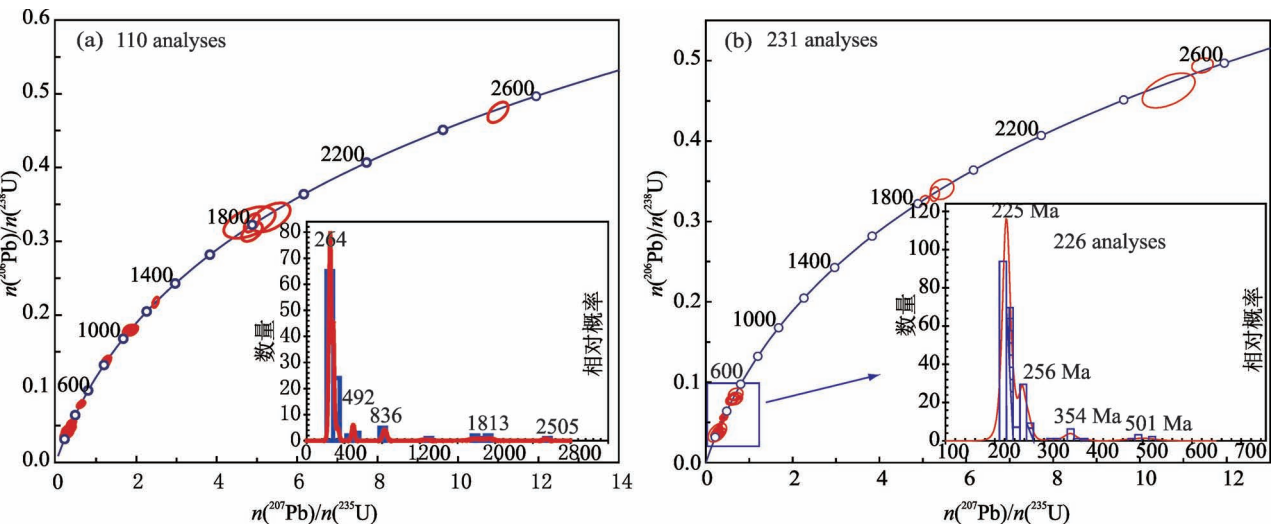


图 6 烟筒山杂岩锆石谐和年龄谱(a, 据 Cao Jialin et al. , 2019 修改)和大酱缸组磨粒石
谐和年龄锆石 U-Pb 年龄谐和图与概率图(b, 据 Wang Bin et al. , 2016a 修改)

Fig. 6 U-Pb concordia diagram for zircon data from the Yantongshan Complex (a, modified from Cao Jialin et al. , 2019) ;
Concordia diagram and relative probability plot of all concordant zircon U-Pb ages for the Dajianggang Formation (b, modified from
Wang Bin et al. , 2016a)

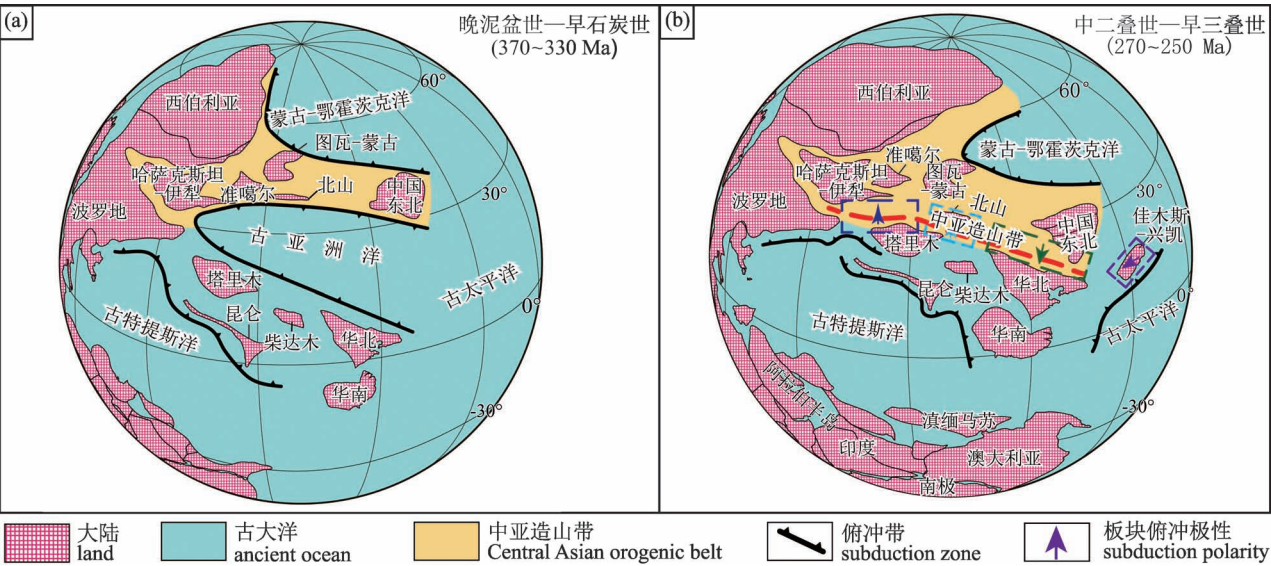


图 7 晚古生代—早中生代前后古亚洲洋与古太平洋构造域叠加转换的动力学模式(据 Zhou Jianbo et al. , 2018 修改)

Fig. 7 Dynamical model of the superposition and transition of the Paleo-Asian Ocean and Paleo-Pacific tectonic domain around
the Late Paleozoic—Early Mesozoic (modified from Zhou Jianbo et al. , 2018)

与哈萨克斯坦—伊犁板块发生拼贴的过程中,根据
高压变质年龄、钉合岩体以及不整合盖层等证据来
综合分析,其闭合时代应为晚石炭世;北山缝合带位
于缝合带中段,形成于敦煌地块与阿拉善地块向北
俯冲与图瓦—蒙古板块发生拼贴的过程中,根据带
内蛇绿岩的年代学证据限定其闭合时代应为早—中
二叠世。阿拉善地块北缘的两条蛇绿岩带连接了北

山缝合带和索伦—长春缝合带,代表了古亚洲在该
区域闭合时形成的缝合带和弧后盆地,其形成时代
应当为中二叠世—晚二叠世早期;索伦—长春缝
合带位于缝合带中—东段,古亚洲洋在该地区同时
发生了南北两侧的双向俯冲,两侧地块在中二叠世—
早三叠世完成拼贴;长春—延吉缝合带形成于中三
叠世前后华北板块与佳木斯—兴凯地块俯冲—增生

过程,较西侧索伦—长春缝合带的形成时间(270~250 Ma)晚 20~30 Ma,因此长春—延吉缝合带与索伦—长春缝合带形成时代与构造背景存在显著的差别。在总结了已有研究进展基础上,笔者等认为南天山—北山—索伦—长春缝合带自西向东发生了四阶段的演化过程,闭合时代自西向东逐渐变年轻,整个过程从晚石炭世一直持续到了三叠世(图 7a,b),其中长春—延吉缝合带记录了古亚洲洋和古太平洋构造域叠加与转换的地质过程。

致谢:评审专家和编辑审阅文稿并提出了详细的修改意见,作者谨致深切谢意。

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Nature and evolution of the South Tianshan Mountains—
Beishan Mountains—Solonker—Changchun Suture

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Abstract: The formation and evolution of the South Tianshan Mountains—Beishan Mountains—Solonker—Changchun Suture, as the final closure position of the Paleo-Asian Ocean, has been the focus of research of the Central Asian Orogenic Belt. Studies on the formation age and subduction polarity of this suture zone can help us to reveal the accretionary and evolutionary history of the Central Asian Orogenic Belt and provide theoretical support for the establishment of a tectonic evolutionary model of the Paleo-Asian Ocean. In this paper, the South Tianshan Mountains —Beishan Mountains —Solonker—Changchun Suture is divided into four segments based on the differences in geotectonic setting, rock composition, closure mode and closure era, and they are, from west to east:

the South Tianshan Suture, the Beishan Suture, the Solonker—Changchun Suture, and the Changchun—Yanji Suture. The South Tianshan Suture is located in the western section of this suture zone and is a result of the northward subduction of the Tarim block, which collapsed with the Kazakhstan—Yili block, and it was formed in the Late Carboniferous based on the evidences of high-pressure metamorphic age, intrusive dike and unconformity cover, etc. The Beishan Suture is located in the middle section, formed during the northward subduction of Dunhuang and Alxa blocks, which collapsed with the Tuva—Mongolia block, and it was formed in the Early—Middle Permian based on the chronological evidences of the ophiolites. The two ophiolite belts at the northern margin of the Alxa block, as the connecting belt between the Beishan Suture and the Solonker—Changchun Suture, represent the suture zone and the post-arc basin formed during the closure process of the Paleo-Asian Ocean, and they were formed in the middle Permian—early Late Permian. The Solonker—Changchun Suture is located in the middle—eastern section, where the Paleo-Asian Ocean underwent simultaneous southward and northward subduction in both directions, and the amalgamation of two sides was completed in the Middle Permian—Early Triassic. The Changchun—Yanji Suture was formed during the amalgamation between the North China Craton and the Jiamusi—Khanka block around the Middle Triassic, and is 20 ~ 30 Ma later than the formation age of the Solonker—Changchun Suture (270~250 Ma), Therefore, these two sutures have significantly different formation age and tectonic setting, and the Changchun—Yanji Suture is not part of the eastern extension of the Solonker—Changchun Suture. The Paleo-Asian Ocean evolved in four stages along the South Tianshan Mountains —Beishan Mountains —Solonker—Changchun suture, and the closure age gradually become younger from west to east, and the whole process lasted from the Late Carboniferous to Triassic, in which the Changchun—Yanji suture recorded the geological process of superposition and transformation between the Paleo-Asian Ocean tectonic domain and the Paleo-Pacific Ocean tectonic domain.

Keywords: Central Asian orogenic belt; South Tianshan Mountains —Beishan Mountains —Solonker—Changchun suture; segmental evolution; subduction polarity; tectonic domain conversion

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