

班公湖—怒江缝合带西段物玛地区原“晚古生代”混杂岩块中发现中生代化石



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内容提要: 在班公湖—怒江缝合带西段铁杂—日雍构造混杂岩带物玛地区, 原划为上石炭统拉嘎组地层中新发现了中生代孢粉化石 14 种及未定种, 这些化石的发现为将该地层修订为下白垩统多尼组提供了古生物学依据。上石炭统拉嘎组被修订后, 研究区班公湖—怒江缝合带不再存在拉嘎组混杂在中生代地层中的现象。同时本次研究在铁杂—日雍构造混杂岩带以南的冈底斯—腾冲地层区物玛分区原划为上石炭统拉嘎组中, 发现中生代孢粉化石 10 种及未定种, 并在其灰岩夹层中发现晚侏罗世珊瑚化石 6 种及未定种, 同时在原划分为中二叠统下拉组中也发现晚侏罗世珊瑚化石。根据化石时代和区域地层对比, 拉嘎组和下拉组分别被修订为上侏罗统萨波直不勒组和吐卡日组。至此, 物玛分区的上侏罗统以萨波直不勒组和吐卡日组为代表, 与其北侧的班公湖—怒江地层区上侏罗统可进行对比。本文以班公湖—怒江缝合带西段铁杂—日雍构造混杂岩带为例, 证实了混杂岩带开展古生物学研究非常必要。

关键词: 西藏; 班公湖—怒江缝合带; 萨波直不勒组; 吐卡日组; 珊瑚; 孢粉

前人认为地层混杂是班公湖—怒江缝合带西段的普遍现象。班公湖—怒江缝合带(以下简称班—怒带)是地质学家关注的大地构造单元之一, 带内的蛇绿岩在西藏内部断续延伸超过 1200 km, 牵涉到许多重大地质问题的讨论(Girardeau et al., 1984; 夏代祥, 1985; 张旗, 1995; 常承法等, 1973; Yin An and Harrison, 2000; 曹圣华等, 2005; 陈玉禄等, 2006; 史仁灯, 2007; Wang Weiliang et al., 2008; 孙立新等, 2011; Pan Guitang et al., 2012; 黄启帅等, 2012; 宋扬等, 2013; 吴小双等, 2014; Yuan Yajuan et al., 2015; Chen Shengsheng et al., 2015; Huang Qishuai et al., 2015; Wu Guichun et al., 2019; 雷传扬等, 2019; 卢雨潇等, 2019)。从最近的研究看, 地质学家更强调了蛇绿混杂岩的概念, 认为班—怒带由超镁铁质—镁铁质岩片、古生代—中生代地层和微陆块构成, 古生代地层以大小不等的构造岩片形式保存在蛇绿岩和中生代构造混杂岩中(王保弟等, 2007; 耿全如等, 2011; 王立全等, 2013; 潘桂棠

等, 2013; Zhu Dicheng et al., 2013; Zhou Tao et al., 2013; 宋扬等, 2014)。地层混杂现象被赋予了重要的构造地质学意义, 朱同兴等(2013)指出班—怒带西段南侧侏罗系木嘎岗日岩群为一套变形较为强烈的、含大量古生代外来岩块的类复理石沉积, 不同岩性、不同时代、不同大小和规模、不同形成环境的外来岩块与基质一起经后期强烈的构造改造, 形成一套总体无序局部有序的构造地层体, 是班公湖—怒江结合带碰撞造山过程的产物。

然而, 铁杂—日雍构造混杂岩带的地质混杂认识却受到古生代混杂岩块时代修订的质疑。铁杂—日雍构造混杂岩带位于班—怒带西段的改则县物玛乡(图 1)。四川省地质调查院(2005)以俄雄—罗仁淌断裂带为界在这一区域划分出班公湖—怒江地层区和冈底斯—腾冲地层区物玛分区(图 2a), 并认为班公湖—怒江地层区存在地层被强烈构造活动混杂的现象, 基质是侏罗系木嘎岗日岩群, 外来岩块主要是上石炭统拉嘎组和中二叠统下拉组, 因此研究

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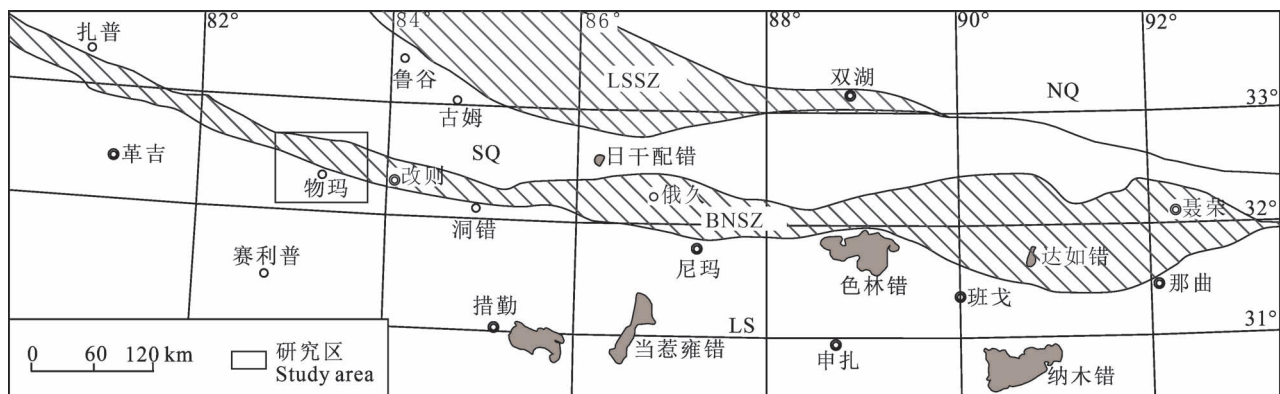


图1 班公湖—怒江缝合带西段示意图(据潘桂棠等, 2013 修改)

Fig. 1 Schematic diagram of the west section of Bangonghu–Nuijiang suture zone (modified after Pan Guitang et al. , 2013&)

LSSZ—龙木错—双湖缝合带;BNSZ—班公湖—怒江缝合带;NQ—北羌塘地块;SQ—南羌塘地块;LS—拉萨地块

LSSL—Longmucuo-Shuanghu suture zone; BNSZ—Bangonghu-Nuijiang suture zone; NQ—North Qiangtang block;

SQ—South Qiangtang block; LS—Lhasa block

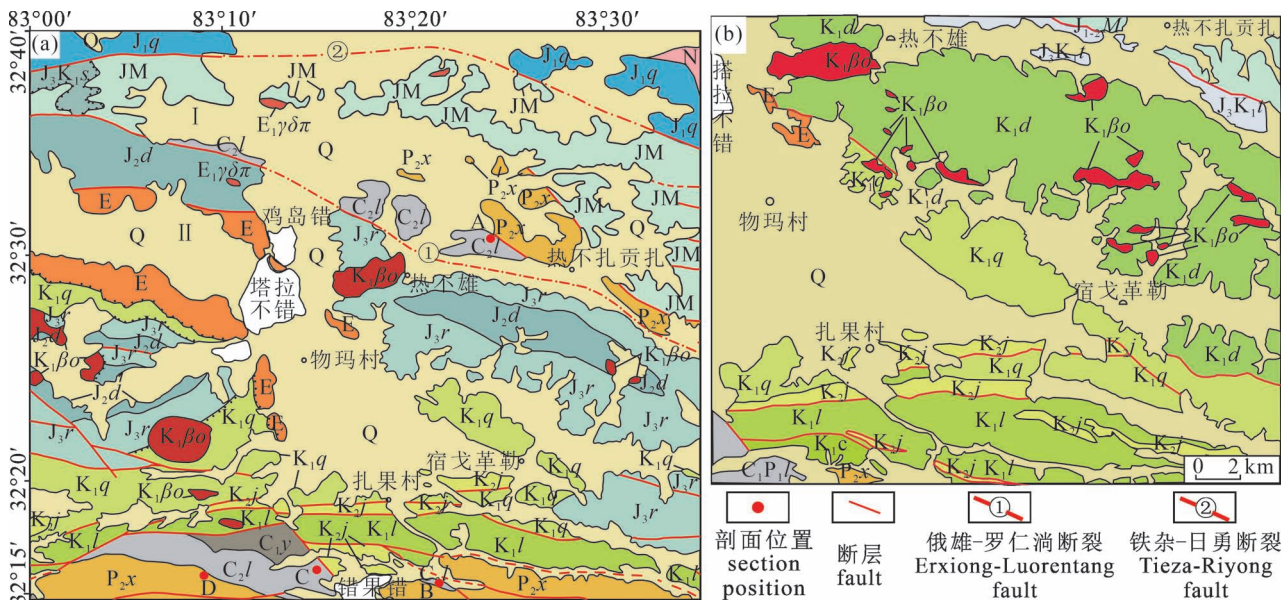


图2 班公湖—怒江缝合带西段物玛地区地质简图(a,据四川省地质调查院^①,2006;b,据张志平等,2019)

Fig. 2 Geological map of Wuma area in the western part of Bangonghu–Nujiang suture zone(a, according to Sichuan

Geological Survey Institute^①, 2006#; b, according to Zhang Zhiping et al. , 2019&)

C₁y—永珠组; C₂l—拉嘎组; P₂x—卜拉组; JM—木嘎岗日岩群; J₁q—曲色组; J₂d—多仁组; J₃r—日松组; J₃K₁t—吐卡日组; J₃K₁s—沙木罗组; K₁q—去申拉组; K₁c—错果错岩段; K₁d—多尼组; K₁l—郎山组; K₂j—竟柱山组; E—古新统; N—新近系; Q—第四系; E₁γδπ—花岗闪长斑岩; K₁β₀—中粒石英闪长岩; I—班公湖—怒江地层区; II—冈底斯—腾冲地层区; a 图与 b 图均表示同一区域

C₁y—Yongzhu Formation; C₂l—Laga Formation; P₂x—Xiala Formation; JM—Mugagangri Group; J₁q—Quse Formation; J₂d—Duoren Formation; J₃r—Risong Formation; J₃K₁t—Tukari Formation; J₃K₁s—Shamuluo Formation; K₁q—Qushenla Formation; K₁c—Cuoguocuo Member; K₁d—Duoni Formation; K₁l—Langshan Formation; K₂j—Jingzhushan Formation; E—Paleocene; N—Neogene; Q—Quaternary; E₁γδπ—Granodiorite porphyry; K₁βo—Medium grained quartz diorite; I—Bangonghu—Nujiang stratigraphic zone; II—Gangdise—Tengchong stratigraphic area; Both a and b represent the same region

区内班公湖—怒江地层区被赋予了构造地质学意义,铁杂—日雍构造混杂岩带也被视为班—怒带在本区通过的重要证据之一。然而,纪占胜等

(2011)、刘晓涛等(2020)和戴蕊等(2020)根据珊瑚化石研究将该带内被视为混杂岩块的中二叠统下拉组灰岩修订为上侏罗统吐卡日组,指出该区不存在

表1 班公湖—怒江缝合带西段地层划分

 整合接触
 不整合
 假整合
 地层缺失
 不明地层

由上可知,古生代混杂岩块的时代改变对铁杂—日雍构造混杂岩带的认识具有决定性影响,因此,地层时代的修订应当基于扎实的古生物学证据。铁杂—日雍构造混杂岩带内的下拉组时代修订有古生物学证据报道,但拉嘎组尚未见到古生物学证据的报道,因此,拉嘎组的古生物学证据是第一个需要解决的问题。另一个问题是物玛分区上侏罗统地层问题。四川省地质调查院(2005)认为物玛分区的上侏罗统为碎屑岩相的日松组(表1)。然而,张志平等(2019)和孔垂鹏等(2019)根据日松组中发现双壳类化石和锆石测年结果,将其修订为下白垩统多尼组。日松组修订为多尼组后,物玛分区的上侏罗统

物玛地区有两个地层古生物学问题亟待解决。

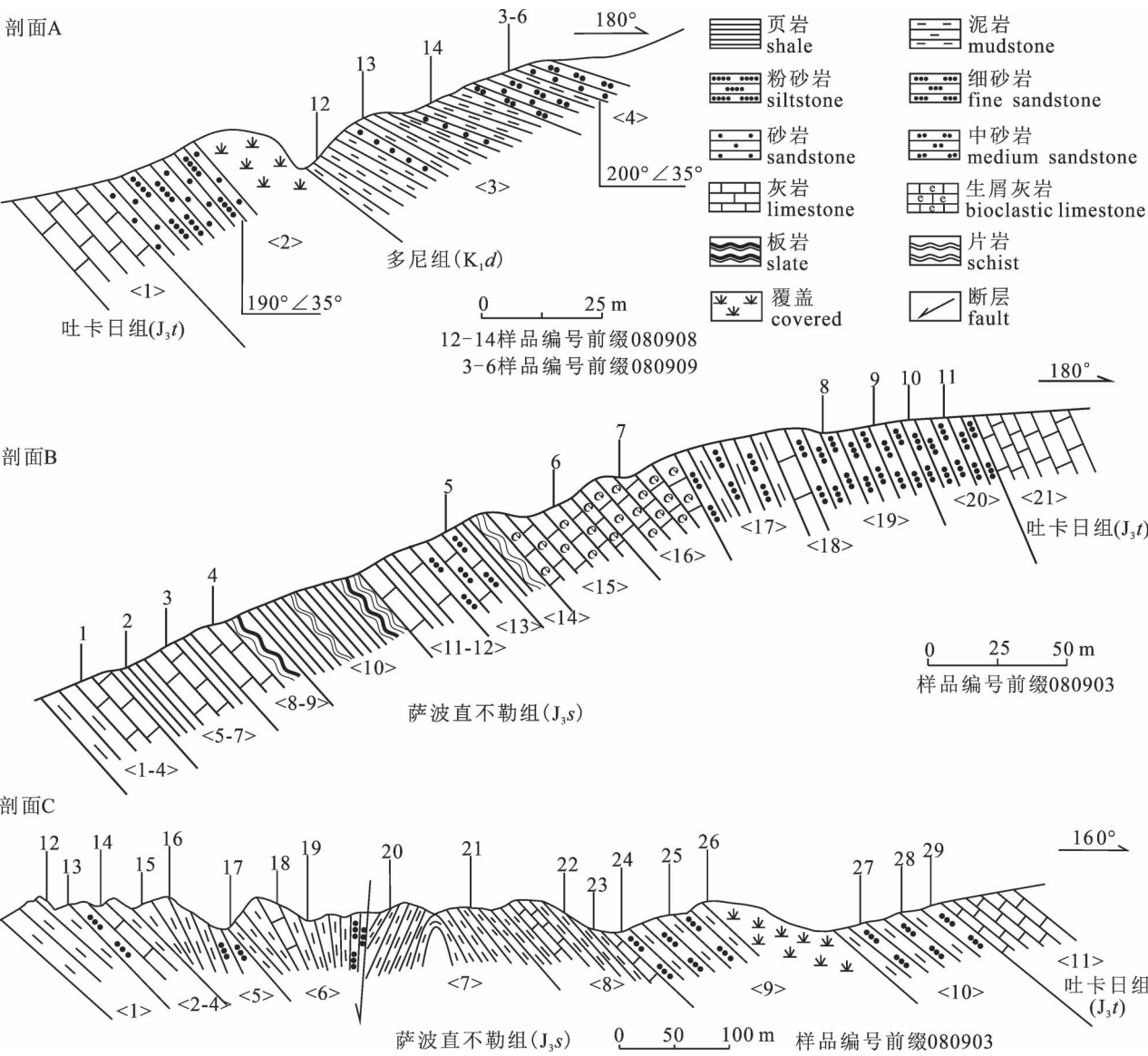


图 3 西藏改则县物玛乡实测剖面 A、B、C
Fig. 3 Section A, B, C of Wuma Township, Gaize County, Tibet

以何为代表就成为一个需要解决的问题。本文在详细研究地层古生物基础上对上述两个问题的研究结果予以报道。

1 实测剖面描述

本次共研究了改则县物玛乡附近的 3 个剖面(图 2a;图 3),剖面 A 位于班公湖—怒江地层区内的甲不拉以南,起点坐标: N32°32′56. 4″, E83°20′1. 7″,海拔高度 4579 m。剖面 B 位于物玛分区的错果错东部,起点坐标: N32°15′49. 2″, E83°21′49. 2″,海拔高度 4757 m。剖面 C 位于错果错西部,起点坐标: N32°16′7. 8″, E83°16′50. 4″,海拔高度 4730 m。

3 个剖面中以碎屑岩为主的地层在 1 : 25 万物玛幅中被划分为上石炭统拉嘎组,碳酸盐岩为主的地层被划分为中二叠统下拉组(四川省地质调查院, 2005)。本次研究在 3 个剖面中均发现了孢粉化石(图 4),并且在剖面 B 的灰岩夹层、剖面 C 以南原划为中二叠统下拉组的灰岩(D 点)中采集到了珊瑚化石(图 5)。根据孢粉和珊瑚化石的鉴定结果和区域地层对比,我们对所测剖面的地层进行了岩石地层单位的重新划分。

1.1 剖面 A

下白垩统多尼组(原划为上石炭统拉嘎组)

未见顶

4. 黄褐色薄层细砂岩、中砂岩,偶夹灰黑色泥质粉砂岩。孢粉样品:080909-3、4、5、6,产孢粉化石
Cyclogranisporites sp., *Lophofriletes* sp.,
Acanthotriletes sp., *Aratrisporites* cf. *scobatus*
Klaus, *Classopollis* cf. *annulatus* (Verbitzka) Li,
Cycadopites sp., *Deltoidospora* cf. *perpusilla*
(Bolch.) Pocock, *Sphagnumsporites* sp.,
Aratrisporites sp., *Punctatisporites* sp., *Lophotriletes*
sp., *Apiculatisporis* sp., *Anaplanisprites* sp.,
Neoraistrickia taylorii Playford et Dettmann,
Neoraistrickia. sp., *Schizosporites* sp. 75.0 m
3. 灰黑色泥页岩夹极薄层砂岩。孢粉样品:
080908 - 12、13、14, 产 孢 粉 化 石:
Abietinaepollenites sp., *Aratrisporites paenulatus*
Playfor et Detmann, *Aratrisporites* sp., *Alisporites*
sp., *Anaplanisporites* sp., *Apiculatisporis* sp.,
Asseretospora sp., *Classopllis* cf. *annulatus*
(Verbitzka) Li, *Cyathidites* sp., *Gleicheniidites*
sp., *Granulatisporites* cf. *arenaster* Phillips et
Felix, *Kraeuselisporites* sp., *Neoraistrickia tayloxii*
Playford et Dettmann, *Plicatipollenites* sp. 42.0 m
2. 黄褐色、灰绿色中薄层中粗砂岩夹灰黑色泥质
粉砂岩、薄板状泥岩。孢粉样品:080908-10、11
40.0 m

整合

- 上侏罗统吐卡日组(原划为中二叠统下拉组)
1. 灰黑色厚层状灰岩 >50.0 m

1.2 剖面 B

上侏罗统吐卡日组(原划为中二叠统下拉组)

上覆地层未测

21. 灰—深灰色生物碎屑灰岩 >20.0 m
- 整合

上侏罗统萨波直不勒组(原划为上石炭统拉嘎组)

20. 灰白色、灰绿色、黄褐色砂岩,覆盖严重 15.0 m
19. 灰绿色薄层状泥质粉砂岩。孢粉样品:080903
-8、9、10、11。产孢粉化石:*Aratrisporites* sp.,
Cyclogranisporites sp., *Leavigatosporites* sp.,
Neoraistrickia sp. 58.0 m
18. 灰黑色砾屑灰岩。 3.0 m
17. 灰黑色—灰绿色泥岩夹灰褐色砂岩,覆盖严重
80.0 m
16. 灰白色、灰黑色厚层生屑灰岩。珊瑚样品:
080903-7。产珊瑚化石:*Cladophyllia turbinata*
(Gregory), *Thecosmilia* sp. 60.0 m
15. 灰黑色厚层生屑灰岩。珊瑚样品:080903-6。
产珊瑚化石:*Dermosmilia* cf. *laxa* Étallon,
Epistreptophyllum sp., *Stylina* sp., *Thamnasteria*
sp., *Thamnasteria mettensis* Milne - Edwards et

- Haime, *Thecosmilia* sp. 40.0 m
14. 灰黑色页岩 10.0 m
13. 砂屑灰岩 2.0 m
12. 灰绿、灰黑色页岩、片岩,向上递变为砂岩与页
岩互层,夹薄层砂屑灰岩。孢粉样品:080903-
5 4.0 m
11. 底部砾屑灰岩,向上递变为砂屑灰岩,顶平底
凸 0.8 m
10. 灰黑色、黄褐色页岩与片岩互层,向上递变为
片岩与板岩互层 30.0 m
9. 灰黑色页岩与黄褐色中砂质板岩、片岩互层
7.0 m
8. 灰黑色中层砾屑灰岩 0.4 m
7. 泥岩、页岩,风化后黄褐色、灰白色 10.0 m
6. 砂屑灰岩、生屑灰岩 0.3 m
5. 灰黑色页岩、片岩 10.0 m
4. 砂屑灰岩 0.2 m
3. 灰黑色页岩、片岩,覆盖严重,表面风化色为灰
绿色 8.0 m
2. 灰黑色粗屑灰岩 0.3 m
1. 黄褐色钙质泥岩、灰褐色页岩 0.5 m
- 未见底

1.3 剖面 C

上侏罗统吐卡日组(原划为中二叠统下拉组)

上覆地层未测

11. 灰—深灰色生物碎屑灰岩 >12.0 m
- 整合

上侏罗统萨波直不勒组(原划为上石炭统拉嘎组)

10. 灰黑色泥岩、砂岩互层。孢粉样品:080903-
25、26、27、28、29、30 380.0 m
9. 黄褐色中薄层状砂岩夹泥岩。孢粉样品:
080903-24 15.0 m
8. 灰黑色泥岩夹灰岩透镜体。孢粉样品:080903-
22、23 140.0 m
7. 灰黑色泥岩与薄层砂岩。孢粉样品:080903-
18、19、20、21。产孢粉化石:*Chasmatosporites*
sp., *Ciootiumspora* sp., *Classopllis* cf. *annulatus*
(Verbitzka) Li, *Dictyophyllidites* sp.,
Foveosporites sp., *Granulatisporites* sp. 150.0 m
6. 灰绿色泥岩夹灰岩透镜体。孢粉样品:080903-
17 40.0 m
5. 黄褐色、灰绿色泥岩夹薄层砂岩。孢粉样品:
080903-15、16 25.0 m
4. 灰绿色泥岩。孢粉样品:080903-14 20.0 m
3. 黄褐色砂屑灰岩 1.0 m
2. 黄褐色砂岩 10.0 m
1. 灰绿色泥岩。孢粉样品:080903-12、13 20.0 m
- 未见底

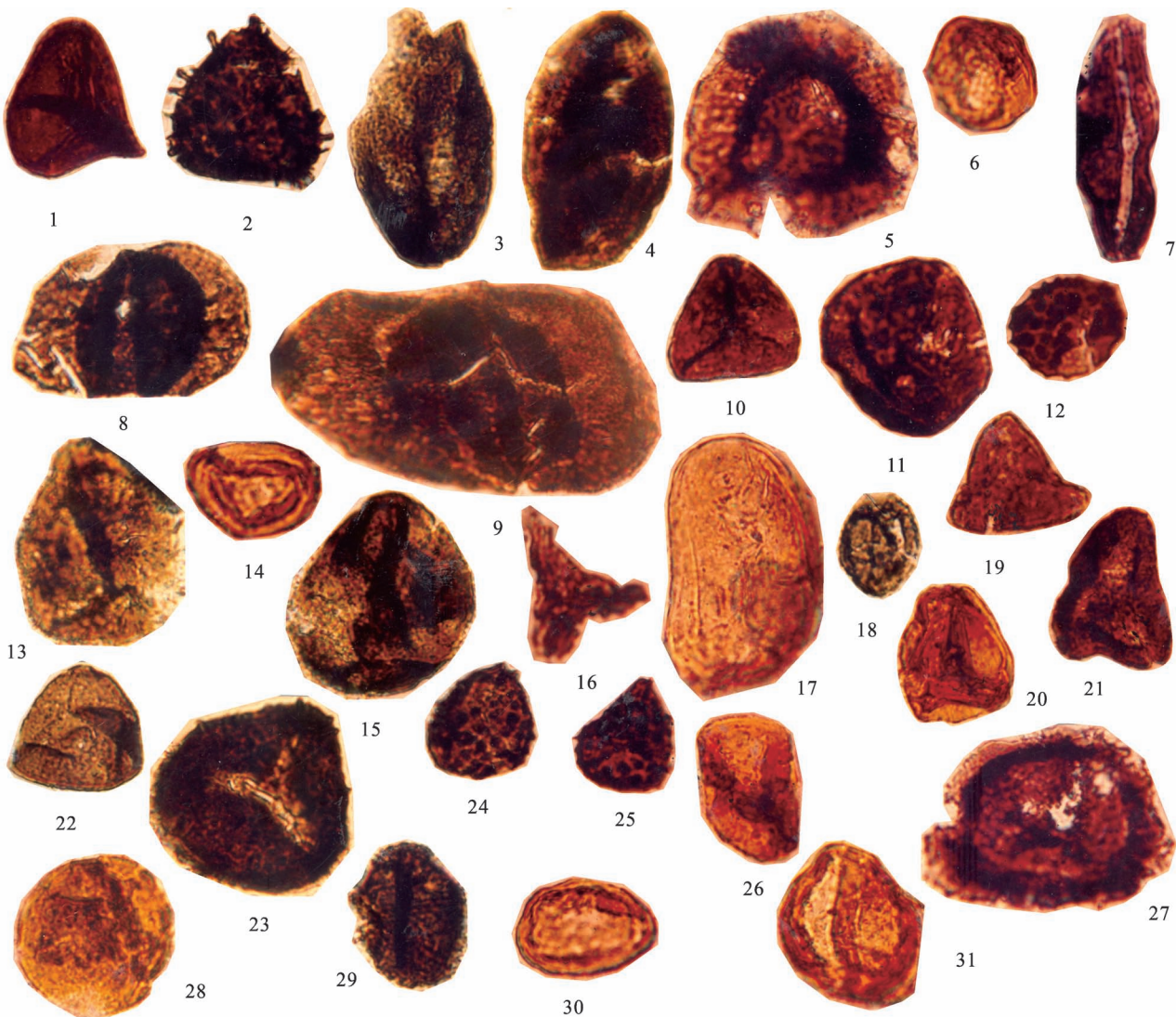


图 4 西藏改则县物玛乡剖面 A、B 和 C 的中生代孢粉

Fig. 4 Mesozoic sporopollen fossils from section A, B and C in Wuma Township, Gaize County, Tibet

标本保存于中国地质科学院地质研究所,化石描述格式:图片中编号,化石属种名称,样品编号,产出剖面(层位)

The specimens are deposited in the Institute of Geology, Chinese Academy of Geological Sciences(CAGS). The format of fossil description is: picture number, genus and species, sample number, occurrence section (horizon): 1, *Cyathidites* sp., 080908-13, A(3); 2, *Neoraistrickia taylorii* Playford et Dettmann, 080908-13, A(3); 3, *Aratrisporites* sp., 080908-13, A(3); 4, *Aratrisporites* sp., 080908-13, A(3); 5, *Plicatipollenites* sp., 080908-13, A(3); 6, *Classoplis* cf. *annulatus* (Verbitzka) Li, 080908-13, A(3); 7, *Cycadopites* sp., 080908-13, A(3); 8, *Alisporites* sp., 080908-13, A(3); 9, *Abietinaepollenites* sp., 080908-13, A(3); 10, *Cyclogranisporites* sp., 080903-8, B(19); 11, *Neoraistrickia* sp., 080903-8, B(19); 12, *Leavigatosporites* sp., 080903-9, B(19); 13, *Aratrisporites* sp., 080903-9, B(19); 14, *Granulatisporites* sp., 080903-18, C(7); 15, *Dictyophyllidites* sp., 080903-18, C(7); 16, *Gleichenioidites* sp., 080908-13, C(1); 17, *Granulatisporites* cf. *arenaster* Phillips et Felix, 080908-13, C(1); 18, *Apiculatisporis* sp., 080908-13, C(1); 19, *Anaplanisporites* sp., 080908-13, C(1); 20, *Anaplanisporites* sp., 080908-13, C(1); 21, *Foveosporites* sp., 080903-18, C(7); 22, *Kraeuselisporites* sp., 080908-13, C(1); 23, *Asseretospora* sp., 080908-13, C(1); 24, *Aratrisporites paenulatus* Playford et Dettmann, 080908-13, C(1); 25, *Classoplis* cf. *annulatus* (Verbitzka) Li, 080903-18, C(7); 26, *Chasmatisporites* sp., 080903-18, C(7); 27, *Deltoidospora* cf. *perpusilla* (Bolch.) Pocock, 080909-5, A(4); 28, *Foveosporites* sp., 080909-3, A(4); 29, *Anaplanisporites* sp., 080909-3, A(4); 30, *Aratrisporites scabratus* Klaus, 080909-6, A(4); 31, *Classoplis* cf. *annulatus* (Verbitzka) Li, 080909-6, A(4)

2 岩石地层单位厘定及其古生物学证据

2.1 剖面 A 的地层划分及其古生物学证据

剖面 A 的地层原划为上石炭统拉嘎组,其北侧的灰岩地层被归入中二叠统下拉组(四川省地质调

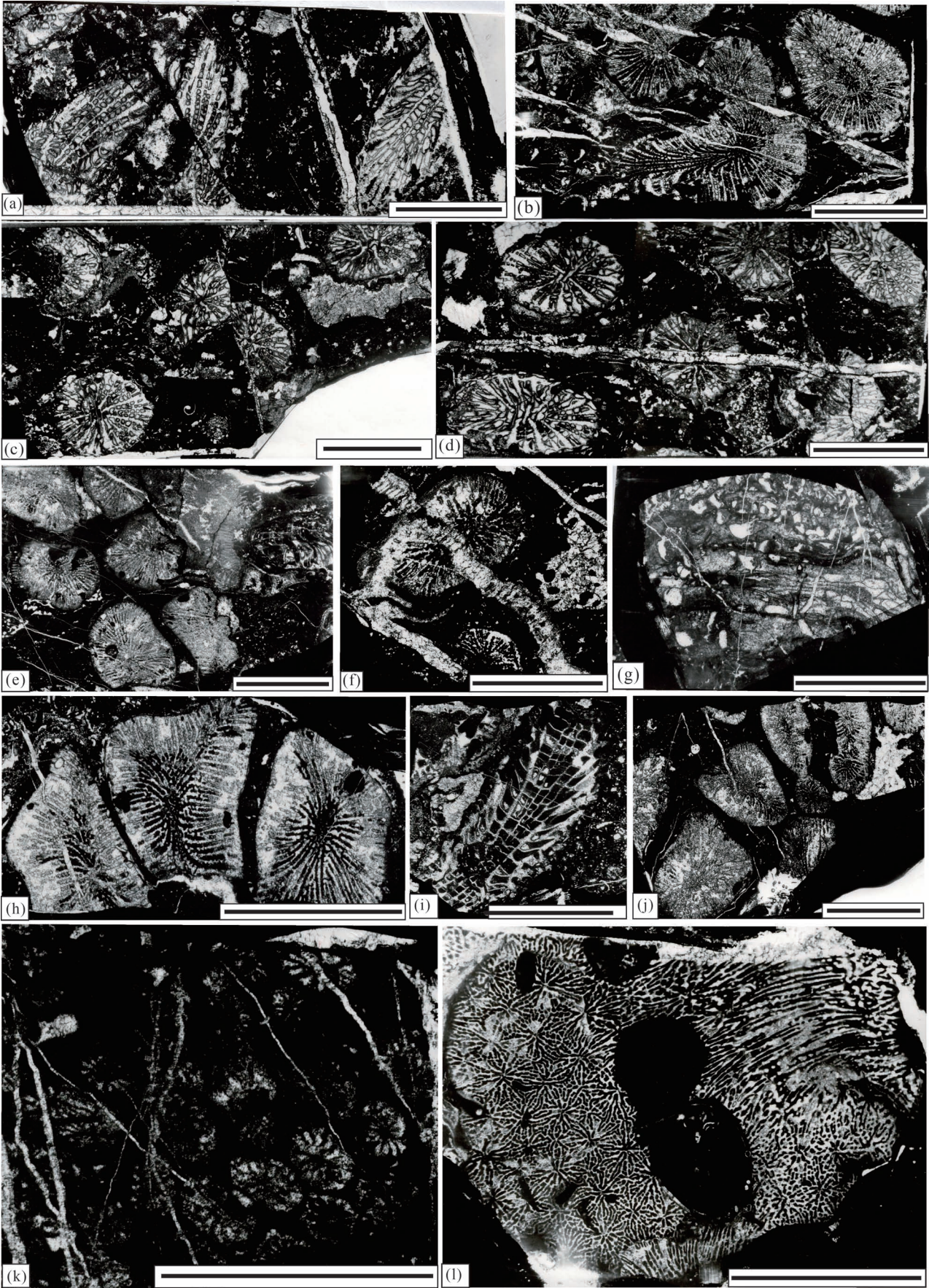


图 5 西藏物玛乡错果错东侧剖面 B 和西侧剖地质点 D 的中生代珊瑚化石

Fig. 5 Mesozoic Coral fossils from section B to the East of Cuoguocuo Lake and the geologic observation spot D to the west of Cuoguocuo Lake, Wuma Township, Tibet

标本保存于中国地质科学院地质研究所,化石描述格式:图片中编号,化石属种名称,样品编号,产出剖面(层位);比例尺为 10 mm;(a),(g),(i)为纵切面,(b)–(f),(h),(j)–(l)为横切面

The corals are kept in the Institute of Geology, CAGS. The format of fossil description is: picture number, genus and species, sample number, occurrence section (horizon). The scale bar is 10 mm. (a),(g),(i)are longitudinal sections, (b)–(f),(h),(j)–(l)are transverse sections; (a) *Thecosmilium* sp., 080903–7, B(16); (b) *Thecosmilium* sp., 080903–6, B(15); (c) *Thecosmilium* sp., 080903–7, B(16); (d) *Thecosmilium* sp., 080903–7, B(16); (e) *Dermosmilium* cf. *laxa* Étallon, 080903–6, B(15); (f) *Thecosmilium* sp., 080903–6, B(15); (g) *Thecosmilium* sp., 080903–7, B(16); (h) *Dermosmilium* cf. *laxa* Étallon, 080903–6, B(15); (i) *Dermosmilium* cf. *laxa* Étallon, 080903–6, B(15); (j) *Dermosmilium* cf. *laxa* Étallon, 080903–6, B(15); (k) *Stylosmilium* cf. *chauviti* Alloiteau, 080901–13, D; (l) *Thamnasteria* cf. *mettensis* Milne-Edwards et Haime, 080901–13, D

查院,2005)。纪占胜等(2011)根据珊瑚化石将剖面 A 北侧的下拉组修订为上侏罗统吐卡日组。吐卡日组是吉林省地质调查院(2005)创建的岩石地层单位。张志平等(2019)将剖面 A 北侧的中二叠统下拉组、剖面 A 原划为上石炭统拉嘎组都修改为吐卡日组,时代置于晚侏罗世—早白垩世,但将热不雄以西的上石炭统拉嘎组单独划分出来,修改为下白垩统多尼组(图 2b),但未见古生物学证据。

本次研究在剖面 A 中发现了孢粉化石,在第 3 层样品 080908–13 中发现有孢粉 *Acanthotriletes* sp., *Anaplanisprites* sp., *Apiculatisporis* sp., *Asseretospora* sp., *Cyathidites* sp., *Foveosporites* sp., *Gleichenioidites* sp., *Granulatisporites* cf. *arenaster* Phillips et Felix, *Kraeuselisporites* sp., *Lophotriletes* sp., *Neoraistrickia taylrii* Playford et Dettmann, *Punctatisporites* sp., *Sphagnumsporites* sp., *Verrucosisorites* sp., 具环单缝孢有 *Aratrisporites* cf. *flrabilis* Playford et Dettmann, *Aratrisporites paenulatus* Playfor et Detmann, 裸子植物花粉有 *Abietinaepollenites* sp., *Alisporites* sp., *Plicatipollenites* sp., 单沟粉有 *Cycadopites* sp., 以及对裂藻 *Schizosporites* sp.。在第 4 层采样 080909–5、080909–6,样品 080909–5 见孢子:*Deltoidospora* cf. *perpusilla* (Bolch) Pocock, *Sphagnumsporites* sp., *Aratrisporites* sp., *Punctatisporites* sp., *Lophotriletes* sp., *Apiculatisporis* sp., *Anaplanisprites* sp., *Neoraistrickia taylrii* Playford et , *Neoraistrickia* sp., 及对裂藻 *Schizosporites* sp.。样品 080909–6 见孢子:*Acanthotriletes* sp., *Aratrisporites* cf. *scobratus* Klaus, *Cyclogranisporites* sp., *Lophofriletes* sp., 花粉有 *Classopollis* cf. *annulatus* (Verbitzka) Li, *Cycadopites* sp.。 *Sphagnumsporites*、*Asseretospora*、*Aratrisporites*、*Deltoidospora*、*Classopollis* 是中生代常见

的孢粉属种,其中 *Aratrisporites* 在我国多数地区发现于早三叠世中、晚期以上的地层(曲立范等,1990)。如在陕甘宁盆地,新疆准噶尔盆地南缘在早、中三叠世地层有少量出现,但在晚三叠世地层中,该属在数量上与种类上多较多于前者地层。在西南、湖北蒲圻群下部、东部山东聊城组此属在孢粉组合中都占显著位置(曲立范,1980;曲立范等,1990;宋之琛等,2000)。该属虽为三叠纪地层中特有分子,但在某些地区侏罗纪地层中也有发现,如新疆准噶尔盆地早侏罗世八道湾组(张望平,1990)。*Classopollis* 广泛分布于世界各地中生代地层,时代从三叠纪至白垩纪,是早侏罗世繁盛的新兴分子(张望平,1990;余静贤,1990)。在国外分布于俄罗斯西伯利亚南部、加拿大西部、澳洲西部、德国、荷兰、法国等地区晚侏罗世地层中(徐钰林等,1980)。

剖面 A 中未发现有古生代孢粉化石,孢粉化石时代从三叠纪一直延续到白垩纪,表明剖面 A 地层时代为中生代。根据野外观察,剖面 A 的碎屑岩厚度大于 150 m,下伏地层为上侏罗统吐卡日组,结合剖面岩性、孢粉化石以及下伏地层时代,笔者等认为剖面 A 的碎屑岩地层应该像热不雄以西的拉嘎组一样,修订为下白垩统多尼组。

2.2 剖面 B 地层划分及其古生物学证据

剖面 B 的第 1–20 层原来归入上石炭统拉嘎组,然而本次研究在剖面 B 中发现中生代孢粉化石,样品 080903–10 中,见 *Punctatisporites* sp., *Cyclogranisporites* sp. 与 *Cycadopites* 花粉。样品 080903–9 中,见孢粉 *Gleichenioidites* sp., *Cyclogranisporites* sp., *Leavigatosporites* sp., *Aratrisporites* sp., *Classopollis* sp.。样品 080903–8 中,孢粉有 *Sphagnumsporites* sp., *Cyclogranisporites* sp., *Neoraistrickia* sp., *Cycadopites* sp.。并且在第

15、16 层还发现珊瑚化石, 080903 - 6 中发现 *Thecosmilia* sp. , *Stylina* sp. , *Epistreptophyllum* sp. , *Dermosmilia* cf. *laxa* Étallon, *Thamnasteria mettensis* Milne-Edwards et Haime, *Thamnasteria* sp. , 080903 - 7 中发现有 *Aratrisporites* sp. , *Cyclogranisporites* sp. , *Neoraistrickia* sp. , *Leavigatosporites* sp. 。珊瑚化石 *Thecosmilia*、*Thamnasteria* 等主要出现在晚侏罗世牛津期—基默里奇期(王泽九等, 2014), 在西藏改则县物玛乡抢古村西北的萨波直不勒组、班戈县保吉乡各昌茶玉剖面中均有发现(纪占胜等, 2011; 孙倩等, 2018a)。珊瑚化石给出了比孢粉更为进一步的时代意见, 指示其时代为晚侏罗世。

根据珊瑚和孢粉的时代意见, 剖面 B 的第 1 ~ 20 层不能归入上石炭统拉嘎组, 而是建议将其与晚侏罗世的岩石地层单位进行对比。根据最新的研究结果(表 1), 研究区晚侏罗世碎屑岩相的岩石地层单位是萨波直不勒组。萨波直不勒组是整合于吐卡日组之下的一套以碎屑岩为主夹少量碳酸盐岩地层(纪占胜等, 2011; 孙倩等, 2017, 2018a, 2018b), 其中的灰岩夹层中产出晚侏罗世的珊瑚化石。剖面 B 的第 1 ~ 20 层的岩性特征就是碎屑岩夹灰岩, 灰岩夹层中含有晚侏罗世的珊瑚化石, 符合萨波直不勒组上部的岩性特征。因此我们将拉嘎组修订为萨波直不勒组, 整合在萨波直不勒组之上的、原来归入中二叠统下拉组的第 21 层及以上的灰岩地层修订为吐卡日组。

2.3 剖面 C 地层划分及其古生物学证据

剖面 C 的地层原来也归入上石炭统拉嘎组, 但是本次从中发现了孢粉化石: 样品 080903 - 24 中, 见 *Verrucosporites* sp. , *Classopollis* cf. *annulatus*。样品 080903 - 21 见孢子 *Dictyophyllidites* sp. , *Ciootiumspora* sp. 。样品 080903 - 18 见孢粉有 *Deltoidospora* sp. , *Granulatisporites* sp. , *Dictyophyllidites* sp. , *Foveosporites* sp. , *Acanthotriletes* sp. , *Asseretospra* sp. , *Classopollis* cf. *annulatus* , *Chasmatosporites* sp. 。样品 080903 - 14, 仅见两粒花粉 *Classopollis* sp. 与 *Cycadopites* sp. 。这些孢粉化石指示地层时代为中生代, 因此不能再归入拉嘎组。本次研究对剖面 C 以南的原划为中二叠统下拉组进行了以采集化石为目的观测, 并且在 D 点处采集到了晚侏罗世珊瑚化石 *Stylosmilia* cf. *chauuti* Alloiteau, *Thamnasteria* cf. *mettensis* Milne-Edwards et Haime, 因此该处下拉组应修订为上侏罗统吐卡日组。而剖面 C 的这套发现了中生代孢粉化石的

碎屑岩地层整合在上侏罗统吐卡日组之下, 剖面 C 的碎屑岩地层建议修订为萨波直不勒组。

3 地质学意义讨论

(1) 本次研究提示我们要对班—怒带西段的地层混杂认识进行检验。地层古生物学的研究结果并不支持铁杂—日雍构造混杂岩的地层混杂认识。四川省地质调查院(2005)认为班—怒带西段存在古生界灰岩块、砂岩块和少量蛇绿岩残片卷入其中形成的混杂构造现象, 铁杂—日雍构造混杂岩带是其中的代表之一, 存在上石炭统拉嘎组和中二叠统下拉组外来岩块混杂在木嘎岗日岩群基质里的典型地质现象(图 2a)。然而, 铁杂—日雍构造混杂岩带内被视为混杂岩块的古生界地层相继被证明为中生代地层。通过珊瑚化石的研究, 中二叠统下拉组被相继证明为上侏罗统吐卡日组(纪占胜等, 2011; 马德胜等, 2015; 李月森等, 2015; 刘晓涛等, 2020)。张志平等(2019)将上石炭统拉嘎组修订为下白垩统多尼组, 本次的孢粉研究结果也支持这种修订。拉嘎组和下拉组的时代修订后, 铁杂—日雍构造混杂岩带不再存在原来认为的地层混杂现象。

地层古生物学研究是检验班—怒带西段地层混杂现象的有效手段。除了铁杂—日雍构造混杂岩带不再存在原来认为的地层混杂现象外, 铁杂—日雍构造混杂岩带的拉嘎组地层时代被修订后(图 2b), 俄雄—罗仁淌断裂南北两侧的地层就不再表现为古生界和中生界之间的巨大差异, 这意味着俄雄—罗仁淌断裂构造活动的强度和规模并没有原来认为的大。铁杂—日雍构造混杂岩内地层混杂和俄雄—罗仁淌断裂的新认识提示我们: 在地质填图实践层面, 班—怒带在物玛地区通过的位置和地质学标志特征可能需要重新考虑。在基础理论研究层面, “班—怒带西段普遍存在古生代地层混杂在其他地质体的地质现象”可能并不是一个毋庸置疑的客观事实, 而是一个有待于检验的主观认识。现阶段最值得检验的就是所谓的古生代外来岩块究竟是不是古生代的。从本文的案例中可以看出, 只需在相关地区开展古生代岩块的古生物学研究即可, 无论是灰岩地层的珊瑚, 还是碎屑岩地层中孢粉, 都是区别古生代地层和中生代地层的有效手段。

(2) 本次研究揭示了物玛地层分区的上侏罗统是萨波直不勒组和吐卡日组。冈底斯地层分区上侏罗统的岩石地层单位存在沙木罗组、达雄群、日松组、沙木罗组、萨波直不勒组和吐卡日组等多种认识

(表 1)。其中日松组的岩石地层单位就来自于本次研究的物玛分区,四川省地质调查院(2005)在 1:25 万物玛幅中将塔拉不错南东区域内碎屑岩地层划为中侏罗统多仁组 and 上侏罗统日松组。张志平等(2019)在上侏罗统日松组中发现早白垩世双壳类 *Neithea nipponica* Hayami 化石,并且获得火山岩锆石 U-Pb 同位素年龄 121.6 ± 2.0 Ma,将该地层重新厘定为下白垩统多尼组。日松组被修订为多尼组之后,物玛地层分区的上侏罗统地层究竟以何为代表就成了未知数。

西藏区调队(1987)^②在错果错西侧建立中侏罗统错果组。夏代祥等(1997)进行西藏地层清理工作中将错果组废弃,后来的地质学家认为其时代与接奴群相当(高振家等,2000;王思恩等,2000)。错果组下部为灰色砂岩、粉砂岩夹少量灰岩与硅质岩,厚达 660 m,其上部为灰岩、生物碎屑灰岩与鲕状灰岩等,厚 150~1000 m,含有双壳、腕足类、珊瑚、菊石等,剖面下未见底,上部与白垩系不整合接触(赵政璋等,2001a,b)。四川省地质调查院(2005)将错果错附近的灰岩地层归入中二叠统下拉组,临近的碎屑岩地层归入上石炭统拉嘎组,但没有实测剖面的证据,错果错地区地层的时代有待于查证,因此开展了本次考察工作。

通过本次剖面 B 和剖面 C 的研究,笔者等在 2 个剖面的上石炭统拉嘎组的地层中都发现了中生代的孢粉化石,在剖面 B 拉嘎组的灰岩夹层中还发现了晚侏罗世的珊瑚化石,在剖面 C 南侧原划为中二叠统下拉组的灰岩(图 2a 中的 D 点处)中采集到了晚侏罗世的珊瑚化石。根据目前的上侏罗统地层划分方案(纪占胜等,2011;孟亚洲等,2017,2018;孙倩等,2017,2018a,b),笔者等将错果错附近的原划为上石炭统拉嘎组地层修订为上侏罗统萨波直不勒组,将中二叠统下拉组修订为上侏罗统吐卡日组。上述修订揭示了物玛地层分区的上侏罗统以萨波直不勒组和吐卡日组为代表。

4 结论和建议

(1)根据孢粉化石、珊瑚化石时代和区域地层对比研究,本文赞同张志平等(2019)将物玛地区班公湖—怒江地层分区的上石炭统拉嘎组修订为下白垩统多尼组的意见。错果错附近的上石炭统拉嘎组的碎屑岩为主的地层应修订为上侏罗统萨波直不勒组、中二叠统下拉组的灰岩地层应修订为上侏罗统吐卡日组。地层修订后,物玛地层分区的上侏罗统

以萨波直不勒组和吐卡日组为代表,与其北侧的班公湖—怒江地层分区的上侏罗统萨波直不勒组和吐卡日组可以对比。

(2)铁杂—日雍构造混杂岩带的下拉组和拉嘎组时代被修订为吐卡日组和多尼组后,铁杂—日雍构造混杂岩带中不存在古生代地层以外来岩块的形式混杂在中生代地层中的现象,俄雄—罗仁淌断裂南北两侧不再表现为石炭系地层与侏罗系地层的断层接触关系。铁杂—日雍构造混杂岩带和俄雄—罗仁淌断裂作为班—怒带在本区通过的证据需要重新商榷。

(3)建议开展班—怒带西段的古生代混杂岩块的古生物学检验工作。本文以班公湖—怒江缝合带西段铁杂—日雍构造混杂岩带为例,证实混杂岩带开展古生物学研究,厘定不同地层单元年代对认识构造混杂岩构造地质学意义至关重要。通过古生物学研究厘定地层时代后,当地的区域构造认识就会发生显著的改变,前人认为班—怒带西段蛇绿构造混杂岩带内地层组成和结构复杂,但从铁杂—日雍构造混杂岩带的具体填图实践看,具有大地构造地质学意义的地层混杂就是古生代地层混杂在中生代地层中。因此对于班—怒带西段而言,重点应该加强对所谓的古生代外来岩块的古生物学研究。古生代岩块时代检验有可能使班—怒带在地质图上的位置发生变化,使我们更为深入地理解班—怒带内地质体的混杂样式、构造活动强度和大地构造属性。

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注 释 / Notes

- ① 四川省地质调查院. 2006. 1:25 万物玛幅地质调查报告.
- ② 西藏区调队. 1987. 1:100 万日土幅区域地质调查报告.

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Mesozoic fossils were found in the original “Late Paleozoic” Mélange blocks at Wuma area in western part of the Bangonghu–Nujiang suture zone

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Objectives: Bangonghu–Nujiang suture zone is one of the most important tectonic units in the north of the Qinghai–Tibet Plateau. It was previously considered to be a common phenomenon that the Paleozoic strata were mixed up in the Mesozoic strata as exotic blocks in the west part of the Bangonghu–Nujiang suture zone. This paper aims to introduce the results of biostratigraphic study on the so-called Paleozoic exotic blocks at Wuma in the Tieza–Riyong Mélange Belt where is thought to develop typical phenomenon of strata-mixing in the west part of Bangonghu–Nujiang suture zone and illustrate that it is necessary to strengthen biostratigraphic research on the exotic blocks in the mélange belt.

Methods: In the case of the Tieza–Riyong Mélange Belt, it was thought that the extraneous rocks, the Upper Carboniferous Laga Formation and the Middle Permian Xiala Formation, were mixed up in the Jurassic Muganggrangri Group, however, the Xiala Formation was revised to be upper Jurassic Tukari Formation based on the study on the age of corals. Thereafter, it is critical to determine the age of the Laga Formation true or not because it involves that if there are Paleozoic exotic blocks or not in the study area. In this study, sporopollen is utilized to check the age of the clastic rocks Laga Formation and coral is used to determine the age of the carbonate rocks Xiala Formation nearby the observed Laga Formation.

Results: First, fourteen species and unidentified species of Mesozoic sporopollen fossils have been found from 9 samples in the strata that were formerly classified as Upper Carboniferous Laga Formation in section A in the Tieza–Riyong Mélange Belt. The sporopollen are *Abietinaepollenites* sp., *Alisporites* sp., *Anaplanisporites* sp., *Apiculatisporis* sp., *Aratrisporites paenulatus* Playfor et Detmann, *Aratrisporites* sp., *Asseretospora* sp., *Classopllis* cf. *annulatus* (Verbitzka) Li, *Cyathidites* sp., *Gleicheniidites* sp., *Granulatisporites* cf. *arenaster* Phillips et Felix, *Kraeuselisporites* sp., *Neoraistrickia tayloxii* Playford et Dettmann, *Plicatipollenites* sp. The discovery of these fossils provides paleontological evidence for the revision of the Laga Formation to the Lower Cretaceous Duoni Formation. After this revision, there is no longer the phenomenon of the Laga Formation mixed with Mesozoic strata in the study area.

Secondly, in Wuma Substratigraphic Zone of the Gangdise–Tengchong Stratigraphic Zone which is to the south of the Tieza–Riyong Mélange Belt, four undetermined species of Mesozoic sporopollen fossils, i. e. *Aratrisporites* sp., *Cyclogranisporites* sp., *Leavigatosporites* sp., *Neoraistrickia* sp. were also found in the Laga Formation in section B and 6 species and unidentified species of Mesozoic sporopollen fossils, i. e. *Chasmatosporites* sp., *Ciootiumspora* sp., *Classopllis* cf. *annulatus* (Verbitzka) Li, *Dictyophyllidites* sp., *Foveosporites* sp., *Granulatisporites* sp. were also found in the Laga Formation in section C in this study. Beyond the discovery of the sporopollen fossils, six species and undetermined species of Late Jurassic coral fossils, i. e. *Dermosmilia* cf. *laxa* Étallon, *Epistreptophyllum* sp., *Stylina* sp., *Thamnasteria mettensis* Milne–Edwards et Haime, *Thamnasteria* sp., *Thecosmilia* sp., were found in the limestone interlayer in the Laga Formation in section B. And late Jurassic Coral fossils, *Stylosmilia* cf. *chauuti* Alloiteau and *Thamnasteria* cf. *mettensis* Milne–Edwards et Haimeare, were found

in the so-called Middle Permian Xiala Formation which is located to the south of section C. According to the above fossils and regional stratigraphic correlation, the so-called Laga Formation and Xiala Formation are respectively revised to be Upper Jurassic Sabozhibule Formation and Tukari Formation.

Conclusions: First, the Laga Formation is reasonable to be revised to the Lower Cretaceous Duoni Formation for the discovery of Mesozoic sporopollen fossils in the Tieza–Riyong Mélange Belt. After the revision of the age of the Laga Formation, there is no longer the phenomenon of the Laga Formation mixed with Mesozoic strata in the Tieza–Riyong Mélange Belt. It is necessary to pay more attention on the paleontological research of the unchecked so-called Paleozoic exotic blocks in the mélange belt in the western part of Bangonghu–Nujiang suture zone. Secondly, it reveals that the upper Jurassic is Sabozhibule Formation and Tukari Formation in Wuma Substratigraphic Zone of the Gangdise–Tengchong Stratigraphic Zone by the age revisions of the Laga Formation and Xiala Formation in section B and section C.

Keywords: Tibet; Bangonghu – Nujiang suture zone; Sabozhibule Formation; Tukari Formation; coral; sporopollen

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