

江南造山带西段四堡群的沉积地质特征 和构造属性探讨

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内容提要: 江南四堡群是江南造山带的重要组成部分, 是认识江南造山带弧沟体系的重要“窗口”。四堡群沉积地质研究表明: ① 四堡群在黔东南地区总体构成一向上变浅的沉积序列: 尧等组为深海-半深海相沉积; 河村组浅海-滨岸相沉积。四堡群在桂北地区总体上也为一向上变浅的沉积序列: 文通组为深海-半深海相沉积; 鱼西组为浊积岩相沉积; ② 桂北地区四堡群碎屑岩系的29组沙纹层理(鲍马序列c段)的玫瑰花图解表明, 古流向总体指向南, 证明物源在其北侧; ③ 四堡群砂岩骨架颗粒统计结果表明, 其物源构造属性为岩浆弧; ④ 基于以上研究, 推测四堡群沉积于弧前盆地。

关键词: 四堡群; 江南造山带; 沉积地质特征; 弧前盆地

新元古代早期, 西北部的上扬子陆块和东南部的华夏陆块拼接为一个构造整体, 形成一条呈向北西弧形突出的北东东-南西西向延伸的前南华纪基底出露带, 被称为“江南造山带”。四堡群位于江南造山带西段, 主要出露于桂北地区及黔东南地区(图1), 为江南造山带的重要组成部分。桂北地区四堡群主体是一套浅变质浊积岩、泥岩岩系, 间夹枕状玄武岩、玄武质科马提岩和火山碎屑岩, 黔东南地区四堡群主要由浅变质砂岩、千枚岩、板岩、泥岩、火山岩屑砂岩组成。

阎明等(1995)通过对四堡群文通组玄武岩, 特别是对科马提质玄武岩建造的地球化学特征进行了研究, 认为四堡群玄武岩建造形成于岩浆弧区。Zhou Meifu 等(2000)通过对出露于池洞地区的两层科马提质玄武岩地球化学特征的分析, 认为这些火山岩建造应是大陆裂谷的产物, 为深部地幔柱上升所控制并受地壳混染的岩浆分异而成。葛文春等(2001)研究了宝坛地区镁铁、超镁铁质侵入岩的构造属性, 认为它们形成于825 Ma前后的地幔柱上涌。周继彬等(2007)的研究进一步支持了葛文春等(2001)的认识。

Wang Xiaolei 等(2007)利用 LA-ICP-MS 方法研

究了出露于桂北地区的四堡群碎屑岩系的碎屑锆石年龄谱系特征, 确定最大底界年龄为 (866.7 ± 3.7) Ma, 并基于碎屑岩系中大量发育火山岩和时间上与碰撞时间相合而推测原型盆地为先陆盆地。Wang Wei 等(2012)进一步研究了桂北四堡群碎屑锆石的年龄谱系, 限定了最大沉积下限年龄为 (835.3 ± 3.6) Ma, 并通过地球化学特征分析, 认为四堡群碎屑岩系的物源为钙碱性火山岩, 构造环境为弧后前陆盆地。而且, Wang Wei 等(2012)的研究还揭示, 鱼西组、九小组的构造环境更倾向于活动陆缘, 而文通组则更倾向于大陆岩岩浆。

前人对该区的研究主要侧重于岩浆岩的研究, 而对沉积岩系研究却极少。四堡群碎屑岩系的物源区类型及原型盆地分析对江南造山带西段弧体系的建立有着重要的作用。本文运用沉积地质学理论对四堡群沉积地质进行研究, 对四堡群原型盆地类型及该区的大地构造格局进行探究。

1 地层序列与时空位置

四堡群主要出露于桂北地区及黔东南地区, 两地区地层关系并未确定, 下面笔者将分别对两地四堡群进行介绍。

注: 本文为中国地质调查局项目(编号: 1212011120115)资助的成果。

收稿日期: 2015-12-12; 改回日期: 2016-07-26; 责任编辑: 刘志强。Doi: 10.16509/j.georeview.2016.05.002

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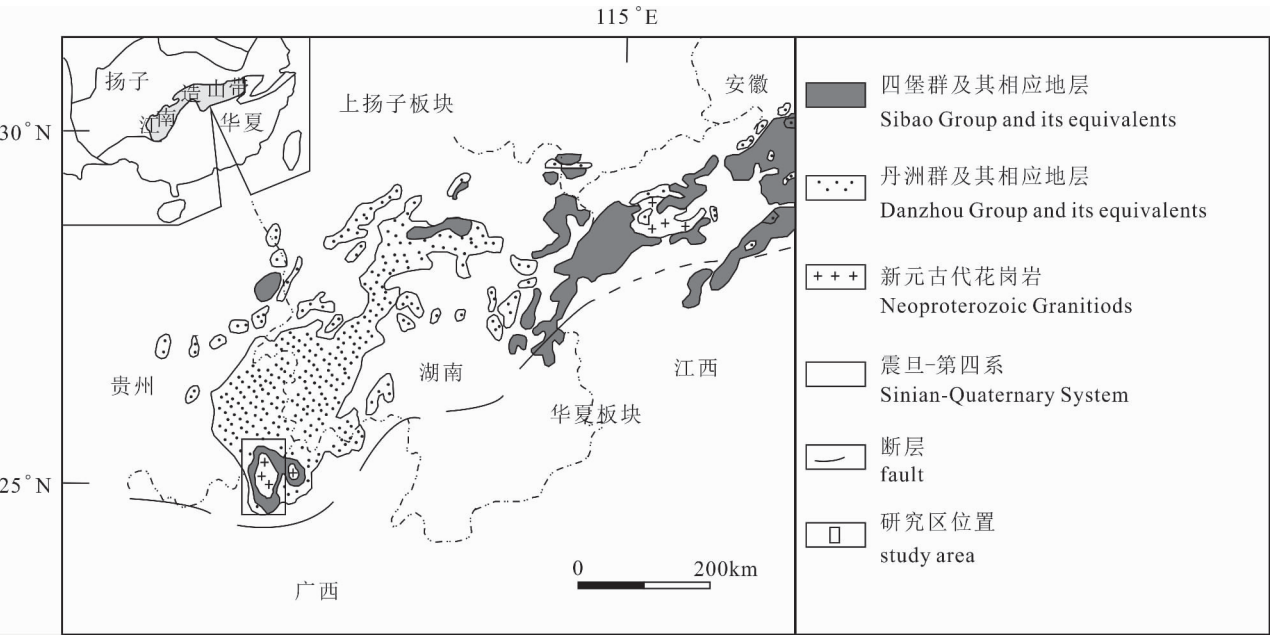


图 1 江南造山带西段的四堡群及其相当时代地层分布 (据周金城,2009,修改)
Fig. 1 Distribution of the Sibao Group and its corresponding formation in the west of the Jiangnan orogenic belt(modified from Zhou Jincheng,2009&)

1.1 桂北地区四堡群

桂北地区四堡群主体是一套浅变质浊积岩、泥岩岩系,间夹有超基性—基性—中酸性火山岩及层状、似层状基性、超基性岩体。自下而上划分九小组、文通组、鱼西组。

九小组主要由轻变质中—细粒长石质石英砂岩、轻变质粗—中粒砂岩、变质粉砂岩、变质泥质粉砂岩、石英绢云母千枚岩、千枚岩、板岩及层状、似层状基性—超基性岩组成。该组未见底,厚度大于 655 m;文通组主要由深灰色、灰绿色变质粉砂岩、板岩、基性熔岩、凝灰岩、科马提岩组成。该层厚约 2514 m;鱼西组主要由变质细砂岩、变质泥质粉砂岩、变质粉砂岩、板岩、绢云母千枚岩夹变质细砂岩、中酸性火山喷出岩组成。该组未见顶,厚度大于 671 m。

1.2 黔东南四堡群

黔东南四堡群主要由浅变质砂岩、千枚岩、板岩、泥岩、火山岩屑砂岩组成。自下而上可划分为尧等组、河村组。

尧等组主要由灰绿色绿泥石绢云母石英片岩、千枚岩、变余砂岩、火山岩屑砂岩组成。片岩、千枚岩为深绿色、灰绿色,后期构造强烈;河村组岩性主要为浅灰绿色变余砂岩、粉砂岩与绢云母石英千枚岩互层,局部可见火山岩屑砂岩。

王孝磊、周金城等(2006)采取四堡群和冷家溪群基底地层中的五个沉积岩样品,进行锆石微区定年工作。获得结果显示,这些样品最年轻年龄均靠近 860 Ma。将获得年龄的数据点进行统计,平均年龄为(866.7 ± 3.7) Ma,从而认为该基底沉积地层的沉积作用应发生在 860 Ma 之后。高林志等(2010)在黔桂交界的十万大山东段获得侵入四堡群的摩天岭花岗岩锆石 SHRIMP U-Pb 年龄(826.8 ± 5.9) Ma;李献华(1999)在摩天岭、元宝山地区获得侵入花岗岩 SHRIMP U-Pb 年龄(819.9 ± 9) Ma、(826 ± 10) Ma,高林志等(2012)在上覆地层下江群甲路组斑脱岩中获得 SHRIMP 锆石 U-Pb 年龄 814 Ma,进而认为四堡群属新元古代中期。

2 沉积地质特征

2.1 沉积相类型及特征

2.1.1 桂北地区

桂北地区四堡群文通组为深海一次深海相;鱼西组主体为浊积岩相,鲍马序列发育且典型(图 2)。

文通组为深海一次深海相,岩石主要为粉砂岩、泥质粉砂岩、泥岩,呈灰黑色,中—厚层状为主,可见水平层理。并可见大量的海相火山喷出岩,具有典型的枕状构造(图 3)。

鱼西组为浊积岩相,岩石类型主要有变岩屑中

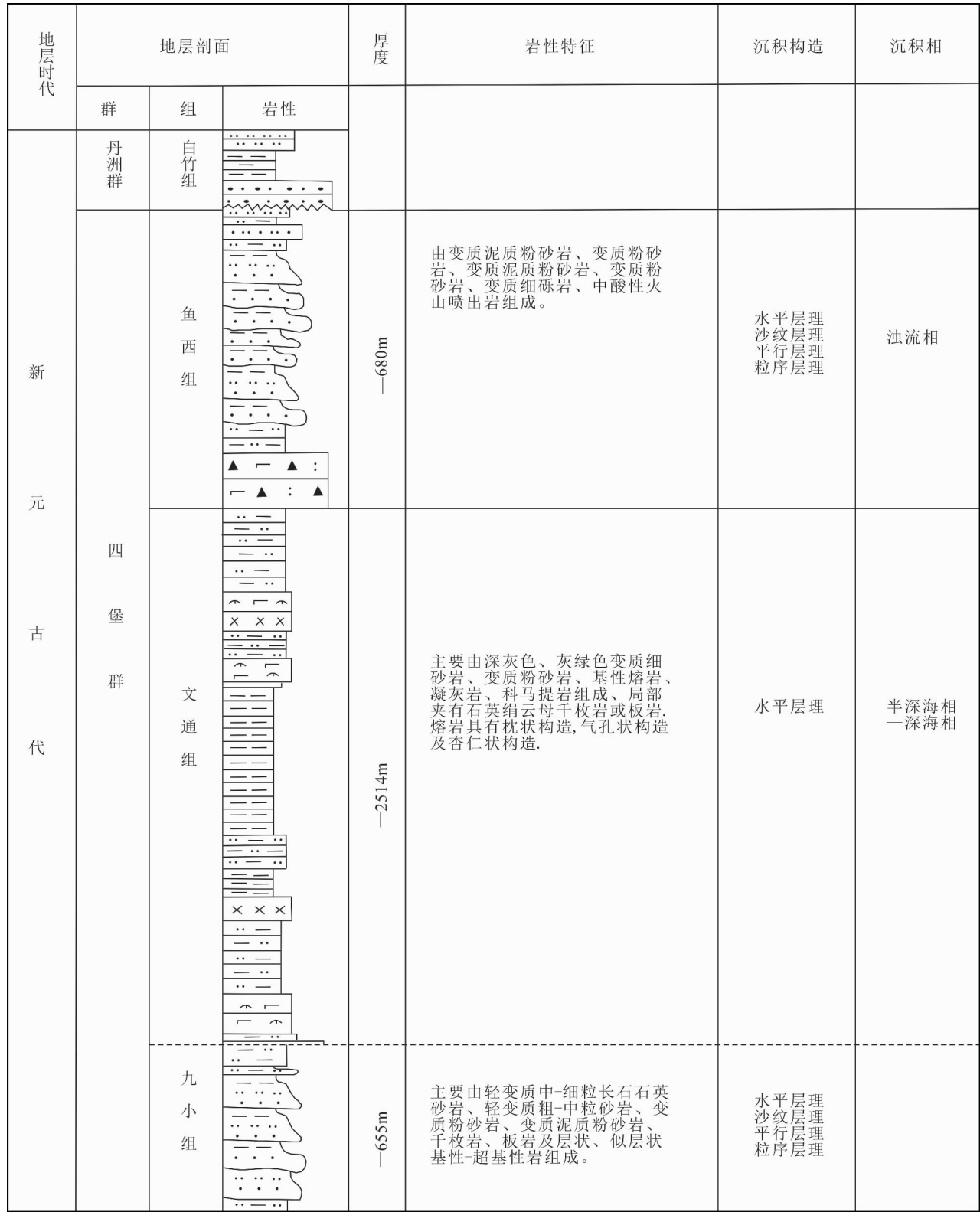


图 2 桂北四堡群综合柱状图

Fig. 2 Integrated histogram in the Sibao Group in north of Guilin

砂岩、细砂岩、粉砂岩,泥岩。呈现中砂岩与细砂岩互层。其沉积具有明显的旋回性,每一个旋回中,岩石颗粒向上依次变细,岩性由中砂岩到细砂岩,再到粉砂岩,代表深海浊流沉积特征。常见的沉积构造

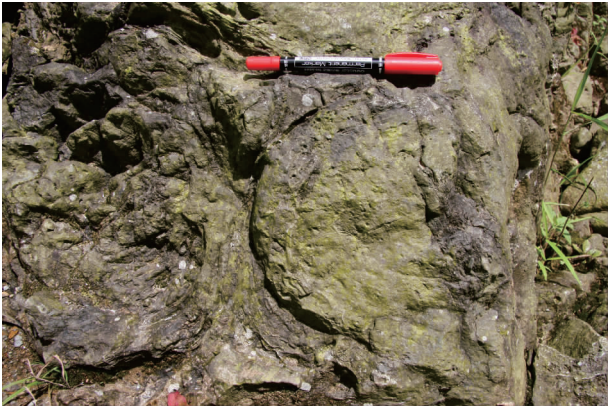


图 3 桂北四堡群火山岩(文通组海相火山岩;具枕状构造;文得村南;镜头朝南)
Fig. 3 Volcanic developed in the Sibao Group in north of Guilin(Marine volcanic rocks in Wentong formation; pillow structure; the south of Wende village; camera lens face south)

有粒序层理(图 4b)、沙纹层理(图 4d)、水平层理,底部发育冲刷构造(图 4c),同时发育有滑塌构造。岩层鲍马序列表现多为 abcde(图 4a)、ade 和 abe 组

合,局部也发育 cde 组合。
2.1.2 黔东南地区
黔东南地区四堡群总体构成一向上变浅的沉积序列:尧等组为深海-半深海浊积岩相;河村组下段为浅海相,上段则为滨岸相(图 5)。
滨岸相也是四堡群常见的沉积相,发育于河村组,主要岩石类型有含砾砂岩、砂岩、粉砂岩、泥岩和火山岩屑砂岩,呈现砂岩与泥岩互层,沉积有明显的

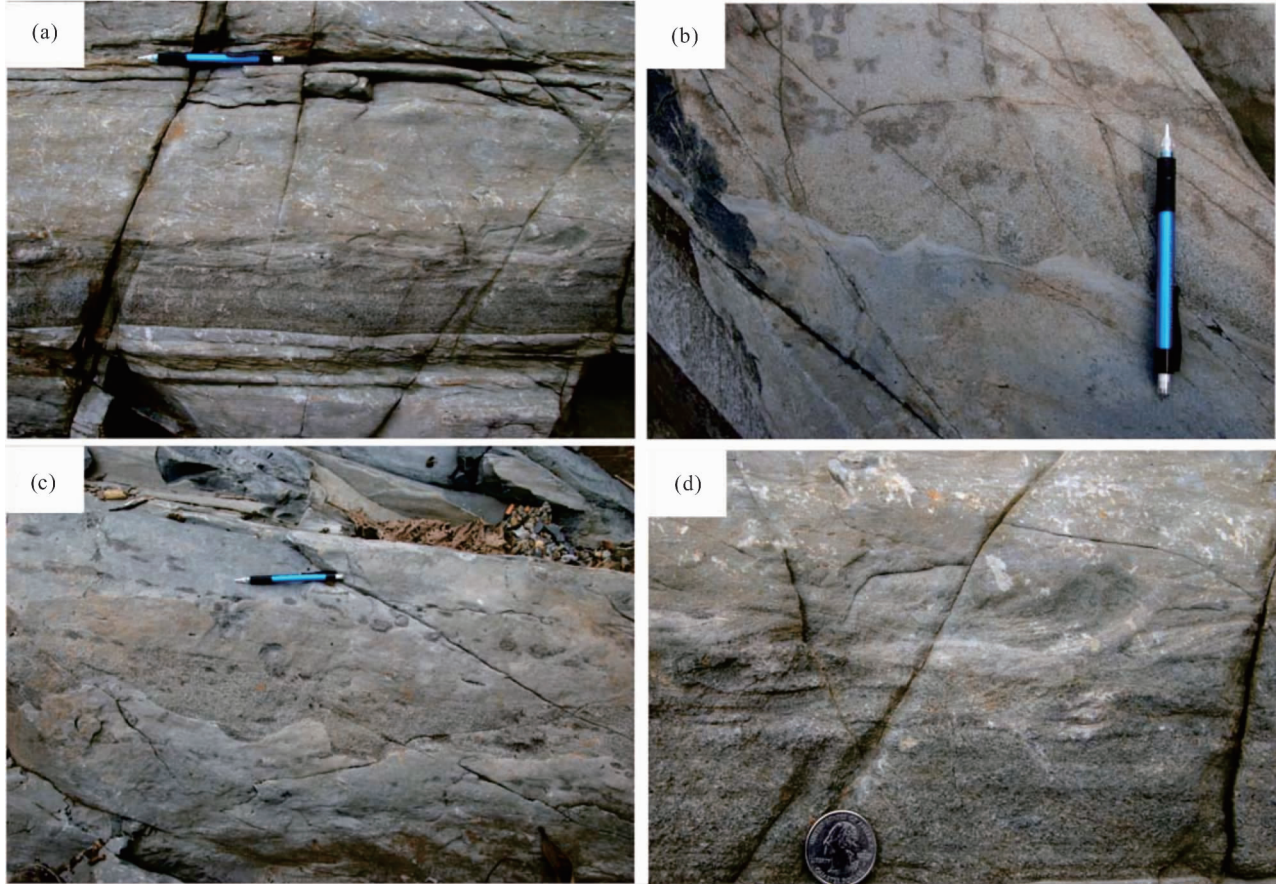


图 4 桂北四堡群中发育的沉积构造

Fig. 4 Typical sedimentary structures developed in the Sibao Group in north of Guilin
(a) 鲍马序列 abcde 段,文通组,牛鼻河北岸,镜头朝北;(b) 鲍马序列 a 段粒序层理,文通组,牛鼻河北岸,镜头朝北;(c) 鲍马序列 a 段底部侵蚀面,文通组,牛鼻河北岸,镜头朝北西;(d) 鲍马序列 c 段沙纹层理(左侧),文通组,牛鼻河北岸,镜头朝北西
(a) The complete Baoma sequence, Wentong formation, the north of Niubi River, Camera lens face north; (b) Graded bedding of Baoma sequence, Wentong formation, the north of Niubi River, Camera lens face north; (c) Erosion surface of Baoma sequence, Wentong formation, the north of Niubi River, Camera lens face north; (d) Current ripple cross lamination of Baoma sequence, Wentong formation, the north of Niubi River, Camera lens face north

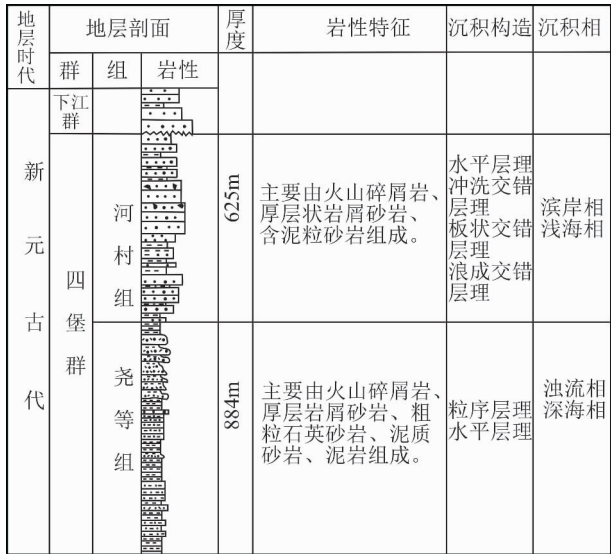


图 5 黔东南四堡群综合柱状图
Fig. 5 Integrated histogram in the Sibao Group in southeast of Guizhou

旋回性,呈现砂岩、粉砂岩、泥岩韵律层,岩石颗粒自下向上由粗变细,有时底部可见含砾砂岩(图 6a),发育大型交错层理(图 6b)、冲洗层理(图 6c)、浪成交错层理(图 6d)、波痕(图 6e)、水平层理,为典型滨岸带沉积。

2.2 沉积物扩散方向

沉积物的扩散方向,特别是浊积岩沉积构造所

反映的沉积物扩散方向与盆地斜坡倾向一致,因此沉积物扩散方向对于认识原型盆地结构和属性具有重要意义。桂北地区四堡群碎屑岩系,基本是由一套成熟度较低的浊积岩组成,其浊流的古流向如何在本地区研究甚少,可以查看的资料几乎没有记载。因此,笔者将对这一问题进行深入的研究。为了确定四堡群的沉积物扩散方向即古水流方向,实测了维洞口—马槽山剖面、四堡—文得剖面中发育的 26 组鲍马序列 c 段沙纹层理位态数据,3 组槽模位态数据。

对实测的数据进行水平校正后,获得沙纹层理形成时的前积层位态(表 1)。根据校正后的前积层倾向做出玫瑰花图(图 7),结果显示:在三岔地区古流向指向南东,在维洞口地区古流向指向南西。由此可知,沉积物的物源来自北方。

2.3 砂岩骨架颗粒特征

Dickinson (1985)等通过对世界上近百个已确定区域构造环境的现代海相和陆相砂岩组分的统计对比和判别分析,建立了大家熟悉的定量判别标准和 Dickinson 三角形模式图,为判断物源区和沉积盆地构造环境提供了一个重要的途经,其中物源区的划分和判别图解的样式在国内外都得到了广泛的应用,并成功地解释了许多物源区的构造背景。其通过对选定层位砂岩样品中的石英、长石和岩屑的含量分别进行统计,用 Dickinson 骨架碎屑颗粒三角图

表 1 四堡群浊积岩系沙纹层理前积层、槽模产状野外实测与水平校正数据表

Table 1 Attitudes of small-size cross-bedding, slot mold in turbidites of the Sibao Group

测点位置: N25°00'43.5"、E108°51'59.7"					
层面 S ₀	前积层 (S _C)		层面 S ₀	前积层 (S _c)	
	野外测量值	水平校正值		野外测量值	水平校正值
99° ∠ 73°	71° ∠ 67°	253.41° ∠ 41.90°	99° ∠ 73°	81° ∠ 76°	237.83° ∠ 40.25°
	81° ∠ 69°	254.74° ∠ 45.90°		82° ∠ 74°	251.05° ∠ 34.79°
	80° ∠ 65°	246.63° ∠ 39.80°		69° ∠ 76°	263.29° ∠ 37.32°
	83° ∠ 74°	252.60° ∠ 36.58°		79° ∠ 74°	249.48° ∠ 41.02°
	78° ∠ 73°	257.71° ∠ 36.33°		73° ∠ 76°	255.22° ∠ 30.41°
	86° ∠ 77°	246.33° ∠ 44.70°		83° ∠ 76°	257.73° ∠ 30.97°
	75° ∠ 69°	243.12° ∠ 46.31°		89° ∠ 71°	265.95° ∠ 35.87°
	72° ∠ 69°	261.82° ∠ 40.73°		79° ∠ 71°	253.41° ∠ 41.90°
	87° ∠ 68°	248.11° ∠ 35.73°		87° ∠ 79°	241.24° ∠ 46.10°
	88° ∠ 78°	251.17° ∠ 37.01°		87° ∠ 71°	260.33° ∠ 37.88°
	91° ∠ 72°	233.52° ∠ 42.87°		87° ∠ 72°	259.79° ∠ 36.94°
	81° ∠ 69°	243.74° ∠ 48.4°		71° ∠ 73°	238.54° ∠ 43.78°
	71° ∠ 70°	247.08° ∠ 38.44°		71° ∠ 76°	235.61° ∠ 41.53°
测点位置 2: N25°06'45.20"、E108°46'51.26"					
143° ∠ 40°	148° ∠ 42°	146.38° ∠ 81.91°		124 ∠ 49	128.76 ∠ 87.48
	143° ∠ 48°	143.00° ∠ 88.00°			

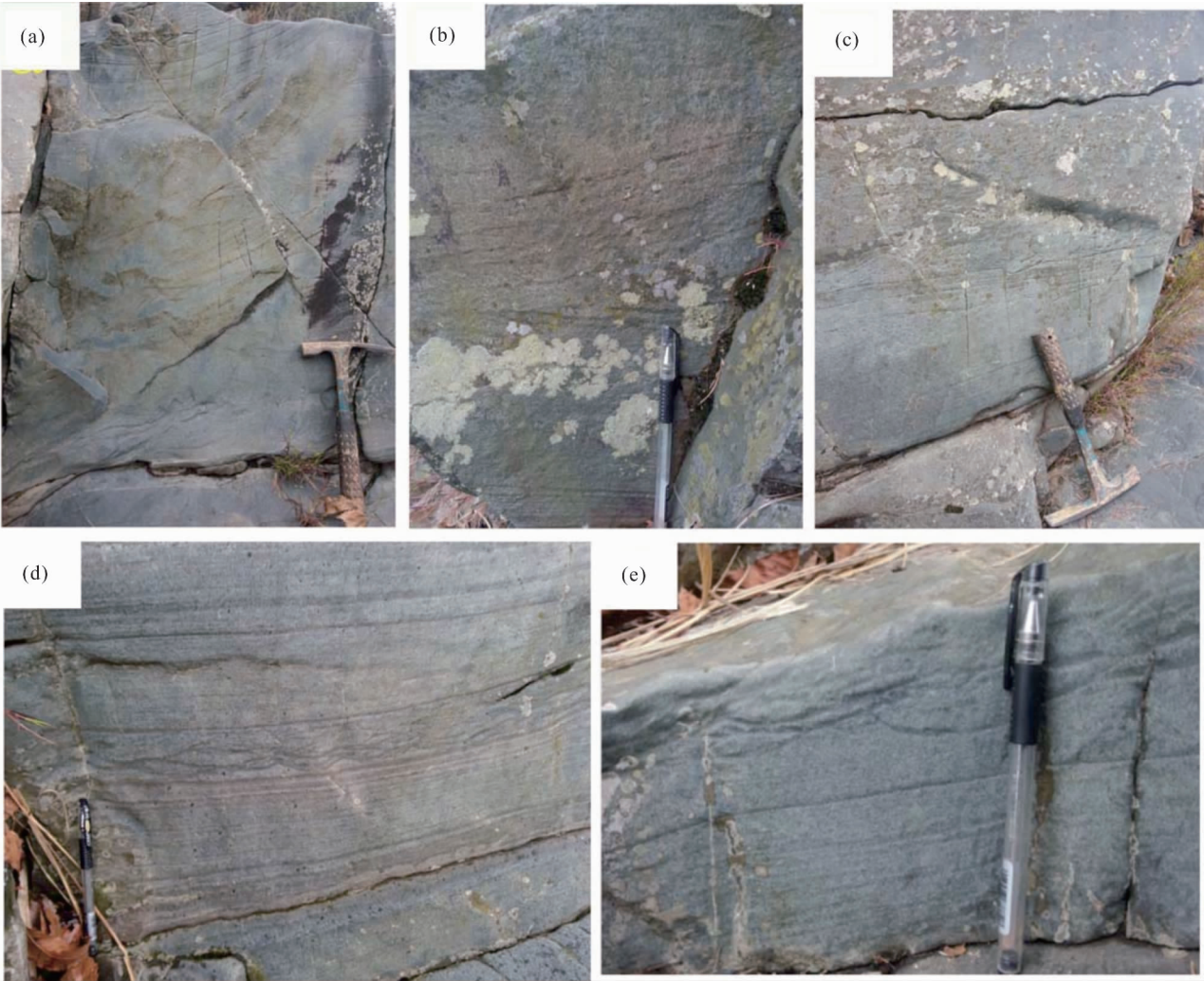


图 6 黔东南四堡群中发育的沉积构造

Fig. 6 Typical sedimentary structures developed in the Sibao Group in southeast of Guizhou

(a) 小水流波痕;从江县令里;镜头朝西北;(b) 冲洗层理与大型交错层理;从江县令里;镜头朝西北(c) 冲洗层理与大型交错层理;从江县令里;镜头朝西北;(d) 平行层理夹浪成交错层理;从江县令里;镜头朝西北;(e) 浪成交错层理与波痕;从江县令里;镜头朝西北
(a) Current-rippled lamination; Lingli village of Congjiang ; Camera lens face northwest; (b) Swash cross bedding and large scale cross-bedding; Lingli village of Congjiang ; Camera lens face northwest; (c) Swash cross bedding and large scale cross-bedding; Lingli village of Congjiang ; Camera lens face northwest; (d) Parallel bedding and wave-built cross bedding; Lingli village of Congjiang ; Camera lens face northwest; (e) wave-built cross bedding and ripple; Lingli village of Congjiang ; Camera lens face northwest

进行投图,根据点的分布情况确定物源类型。

(1)桂北地区四堡群砂岩骨架颗粒特征。本次研究共采集了桂北四堡群砂岩样品 75 个。其中,在三岔地区采集砂岩样品 30(SY-01-1 ~ SY-01-30),文得地区采集砂岩样品 30 块(S0731-1 ~ S0731-30),在维洞口地区采集砂岩样品 15(SY-03-1 ~ SY-03-15)块,其中大部分构造变形强烈,影响矿物的识别,本次选取其中的 32 块进行砂岩骨架颗粒组分分析。结果显示,石英含量较低,在 30% ~ 50%,且单晶石英较多,而长石则高达 40% 左右,岩屑含量较

高在 30% ~ 60%。

(2)黔东南地区四堡群砂岩骨架颗粒特征。本次研究在黔东南四堡群采集了四堡群砂岩样品 18 个,并对其进行砂岩骨架颗粒组分分析。结果显示,石英含量较低,在 20% ~ 40%,且单晶石英较多,几乎不含多晶石英,长石含量为 10% ~ 30%,岩屑含量则高达 40% ~ 65%,且多为火山岩屑,沉积岩屑极少见。

为了进一步确定这套碎屑岩的物源构造性质,对上述两个地区统计的样品分别进行了 Q—F—L

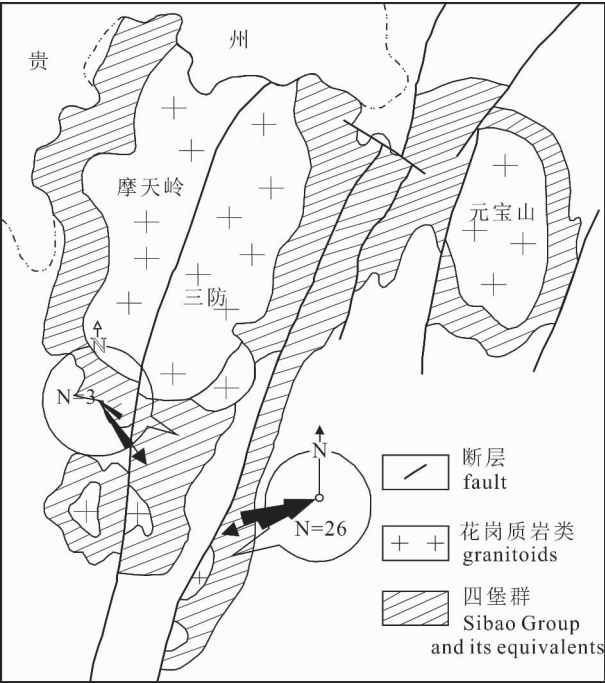


图 7 四堡群碎屑岩系沙纹层理、冲槽产状玫瑰花图解与推测的古流向
Fig. 7 Rose diagrams of small-size cross-beddings, slot mold developed in the Sibao Group and the induced paleocurrent directions

图解、Qm—F—L 图解和 Qp—Lvm—Lsm 图解。砂岩图解结果显示,其物源区构造属性均是岩浆弧(图 8, 图 9)。

综上所述,华南四堡群碎屑岩系的物源区构造属性为岩浆弧,从 Dickinson 三角图解中可以看出,

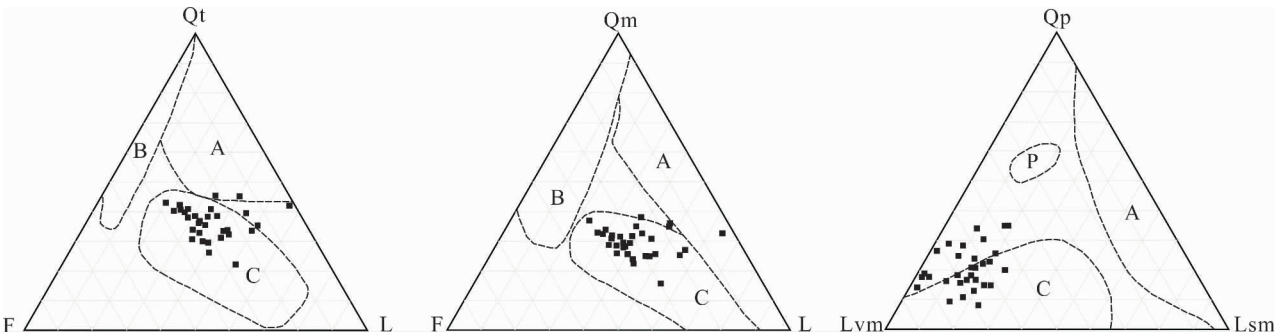


图 8 桂北四堡群 Qt—F—L、Qm—F—L 和 Qp—Lvm—Lsm 三角图解的物源板块构造类型(物源分区据 Dickinson,1984)
Fig. 8 Qt—F—L, Qm—F—L and Qp—Lvm—Lsm triangular diagrams showing dynamic environments of sandstone from the Sibao Group, in north of Guilin (after Dickinson,1984)
Qt—所有石英颗粒, Qp—多晶石英颗粒, F—所有长石颗粒; Lvm—所有火山岩和变质火山岩颗粒, Lsm—所有沉积岩或变质沉积岩颗粒;
A—再旋回造山带; B—稳定克拉通或隆起的基底; C—火山弧; P—增生楔
Qt—total quartzose grains; Qp—polycrystalline quartz; F—total feldspar grains; Lvm—total volcanic—metavolcanic rock fragments; Lsm—total sedimentary—metasedimentary rock fragments. . A—re-cyclic orogenic belt; B—stable craton or uplifted basement; C—magmatic arc; P—accretionary wedges

四堡群碎屑岩系中火山碎屑含量较多,表示距离火山弧源区很近,通过对黔东南四堡群和桂北四堡群中砂岩骨架统计结果进行对比分析,可发现黔东南四堡群的火山岩屑含量较多,这可能是当时黔东南地区比桂北地区岩浆弧更近。且古流向指向由北向南,表明当时地势北高南低。因此,可以推测四堡群北部有一个岩浆弧,由于岩浆弧的喷发,源源不断的向南提供碎屑沉积物,结合沉积相垂向变化及构造变形特征,综合表明四堡群碎屑岩系为弧前盆地沉积。

3 盆地原型及古地理格局属性

四堡群发育碎屑浊积岩相,浊积岩相可见鲍马序列、滑塌构造,其应属活动类型的沉积建造;通过对四堡群维洞口、三岔地区的古流向示向沉积构造的野外测量和室内分析,可知维洞口古流向指向南东,三岔地区古流向指向南西,可知四堡群古流向总体上指向南,物源来自于北方,四堡群位于盆地的北缘;砂岩骨架颗粒组分统计结果显示,四堡群物源构造属性为火山弧。综合分析,笔者认为四堡群的盆地原型为弧前盆地。

4 结论

(1)桂北四堡群总体构成一向上变浅的沉积序列:文通组主要为深海—次深海相,可见大量的海相火山喷出岩;鱼西组主体为浊积岩相,鲍马序列发育且典型。黔东南的四堡群总体也构成一向上变浅的沉积序列:尧等组为深海—半深海泥岩和浊积岩;河

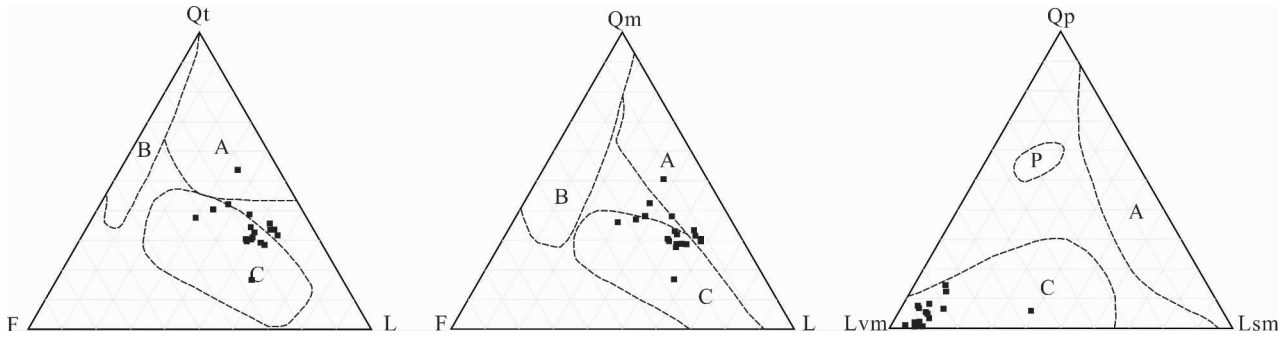


图 9 黔东南四堡群 Qt—F—L、Qm—F—L 和 Qp—Lvm—Lsm 三角图解的物源板块构造类型(物源分区据 Dickinson,1984)

Fig. 9 Qt—F—L, Qm—F—L and Qp—Lvm—Lsm triangular diagrams showing dynamic environments of sandstone from the Sibao Group, in Southeast of Guizhou (after Dickinson,1984)

Qt—所有石英颗粒, Qp—多晶石英颗粒, F—所有长石颗粒; Lvm—所有火山岩和变质火山岩颗粒, Lsm—所有沉积岩或变质沉积岩颗粒; A—再旋回造山带; B—稳定克拉通或隆起的基底; C—火山弧; P—增生楔

Qt—total quartzose grains; Qp—polycrystalline quartz; F—total feldspar grains; Lvm—total volcanic—metavolcanic rock fragments; Lsm—total sedimentary—metasedimentary rock fragments. A—re-cyclic orogenic belt; B—stable craton or uplifted basement; C—magmatic arc; P—accretionary wedges

村组下段为浅海碎屑岩,上段则为滨岸带沉积。

(2) 桂北地区四堡群沙纹层理层理(鲍马序列 c 段)、冲槽等古流向示向沉积构造位态统计分析结果表明:维洞口地区古流向朝南西,三岔地区古流向朝南东,四堡群古流向总体上朝南,位于盆地的北缘,物源来自北方。沉积相对比分析进一步印证了这一认识。

(3) 四堡群砂岩骨架颗粒组分统计分析表明,石英含量较低,在 30% ~ 50%,且单晶石英较多,而长石则高达 40% 左右,岩屑含量较高在 30% ~ 60%,火山岩屑较多,沉积岩屑少见,砂岩骨架颗粒成分三角图解显示,这些碎屑岩的物源区为火山弧。

(4) 综合分析四堡群形成的构造环境、古流向、砂岩骨架颗粒特征等,四堡群碎屑岩形成于活动陆缘区的弧前盆地内,推测当时处于汇聚型板块边缘内,属弧前盆地沉积。

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A Discussion on Sedimentary Characteristics and Structural Properties of the Sibao Group in the West Segment of the Jiangnan Orogenic Belt

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Abstract: The Sibao Group in South China is an important part of the Jiangnan Orogenic Belt, it's a "window" to recognize Jiang Orogenic Belt's trench—arc system. we use sedimentology to study its prototype basin and tectonic paleogeographic pattern. Based on these works, we have five following understanding: ① The Sibao Group, in Southeast of Guizhou, is an upward-shallowing sequence. Yaodeng Formation is deep sea—bathyal mudstone and turbidite; The upper of Hecun Formation is neritic clastic, the lower of Hecun Formation is shore facies. Sibao Group, in north of Guilin, are turbidite facies, and the Bouma sequence are developed. ② The rose diagram of inclined bedding in clastic of Sibao Group, in north of Guilin, show that paleocurrent direction Pointing to the south, the provenance is from the north. ③ Statistical results of skeleton particles show that the tectonic feature of the provenance of Sibao Group is Magmatic Arc. ④ Based on the above findings, we suggest that the tectonic environment of Sibao Group's clastic rocks is forearc basin.

Keywords: Sibao Group; Jiangnan orogenic belt; sedimentary characteristics; fore-arc basin

Acknowledgements: This study was supported by China Geological Survey, We are grateful to anonymous reviewers and editing committee for their thoughtful reviews and constructive comments.

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Manuscript received on: 2015-12-12; Accepted on: 2016-07-26; Edited by: LIU Zhiqiang.

Doi: 10.16509/j.georeview.2016.05.002