

地热与热液型铀矿成因联系:研究现状及解决方法

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内容提要: 地热能作为一种清洁可再生能源已被广泛应用于发电、供暖等领域, 放射性产热是地球内热的主要来源。目前, 关于放射性产热的研究主要集中于高放射性花岗岩在干热岩资源探测和潜力评价中的应用, 而同样作为放射性产热源的热液型铀矿与地热研究不足, 二者的成因联系尚缺乏深入研究。为加强对地热与热液型铀矿成因联系的研究, 本文收集和整理了国内外大量文献和相关资料, 分析了目前的研究进展及存在的问题, 创新性地提出了解决二者成因联系的研究方法。可望丰富地热地质理论, 为地热与热液型铀矿床的成因联系提供新方法和新参考, 同时, 为深部地热、干热岩和固体矿产综合找矿, 即“热矿共采”提供一定的理论依据。

关键词: 地热; 热液型铀矿; 成因联系; 研究现状; 解决方法

地热能作为一种清洁可再生能源, 因具有热流密度大、参数稳定(流量和温度)及收集输送方便等优点(Wang Kai et al., 2018), 在发电、供暖、制冷等领域广泛应用(Zheng Keyan et al., 2015; 多吉等, 2017)。据统计, 全球地热发电装机容量已超过12000 MWe, 实现发电大于73000 GWh, 到2020年装机容量将达21443 MWe(Bertani, 2015)。我国地热资源分布广、储量丰富, 如川西康定(张云辉, 2019)、河北马头营(张保建等, 2020)、青海共和盆地(唐显春等, 2020)和胶东半岛(史猛等, 2019)等。大力开发利用地热资源对于调整能源供应结构、缓解能源供应压力、保障能源供应安全和促进地区经济发展具重要作用(王贵玲等, 2017, 2018)。

研究发现, 放射性元素衰变释放的能量是地球内热和地壳热流的主要来源, 而放射性元素中铀元素对生热率的贡献最大(汪集旸和孙占学, 2001; Arevalo et al., 2009)。目前, 关于放射性产热的研究主要集中于高放射性花岗岩体在干热岩资源探测和潜力评价中的应用(甘浩男等, 2015; 王安东等, 2015; 杨立中, 2016; 杨立中等, 2016)。热液型铀矿作为重要的放射性热源, 与地热在时空上具有密

切联系(徐增亮, 1982; 李学礼, 1992; Hillis et al., 2004; 王微等, 2011; 王俊虎等, 2011)。热液型铀矿是经历热液活动形成的, 是在特定地质构造环境中古水热系统活动的产物(李学礼和杨忠耀, 1984), 大气降水在这一类铀矿的形成中通常起非常重要的作用(El-Feky, 2011; Helmy et al., 2014), 它们的形成不仅需要丰富的铀和迁移富集的介质, 而且还应具有良好的热源, 以促使铀的活化迁移成矿。水热型地热资源大部分是因大气降水深循环加热而形成的(罗璐等, 2019; 袁利娟等, 2020; 李泓泉等, 2020)。热液型铀矿与地热之间的相关性很可能预示二者在成因上的密切联系, 但目前关于二者的成因联系尚缺乏系统的研究。通过对地热与热液型铀矿成因联系的研究将为我国华南地区热液型铀矿与地热的成因联系研究提供新方法和新参考, 可望丰富地热地质理论。同时, 为深部地热、干热岩和固体矿产综合找矿, 即“热矿共采”提供一定理论依据。

1 地热与铀矿关系研究现状

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1.1 地热与铀矿空间关系

大凡盛产铀矿的地方,特别是与热液活动有关的铀矿,一般伴随大量的热泉、温泉或高温围岩蚀变等地热异常分布(李庆阳等,2010),中国许多热液型铀矿区现今仍是热高场区,仍有许多热泉分布(陈友良,2008),例如,华南地区热液型铀矿分布广泛、类型多(花岗岩型、火山岩型和碳硅泥岩型)(Hu Ruizhong et al., 2008),温泉的分布与铀矿床(田)、铀成矿带分布一致(图1)。许多学者已经发现地热、温泉与铀矿空间分布一致性的规律,即温泉出露在铀矿田内或附近,或者与铀矿田(带)受同一构造体系控制(张万良等,2015)。比如,江西、湖南、广东等省是我国东南沿海隆起型地下热水的主要分布区,是地下热水分布较集中的一个带,水温较高,这三省同时也是热液型铀矿床的主要产区(李学礼,

1992)。在江西,著名的相山铀矿田发育有汤溪温泉、临川温泉、德兴温泉和马鞍坪温泉等多处温泉(张卫民,1992; 张万良等,2015);寻乌县横跨地区121号铀矿化点与8号温泉出露位置较近(黄明光等,2010);丰州铀矿床与温泉共伴生(周立坚,2016);沿河草坑铀矿田的控矿断裂会昌断裂分布有谢坊温泉和罗塘温泉等多处温泉(张万良和李富梅,2014)。在湖南,“中国核工业第一功勋铀矿田”金银寨铀矿田内温泉、热异常点非常发育(祝青云和覃正良,2008),铀矿田内1号主井井内至今仍不断涌出热水;热水镇地热勘查区出露的多处温泉与鹿井—城口特大—大型铀矿床成矿带空间分布一致(周立坚,2016)。在广东,著名的下庄铀矿田地热资源非常丰富,发育有象司前温泉、茅山温泉和坝仔温泉等多处温泉,同时,还实施了一系列的地热钻

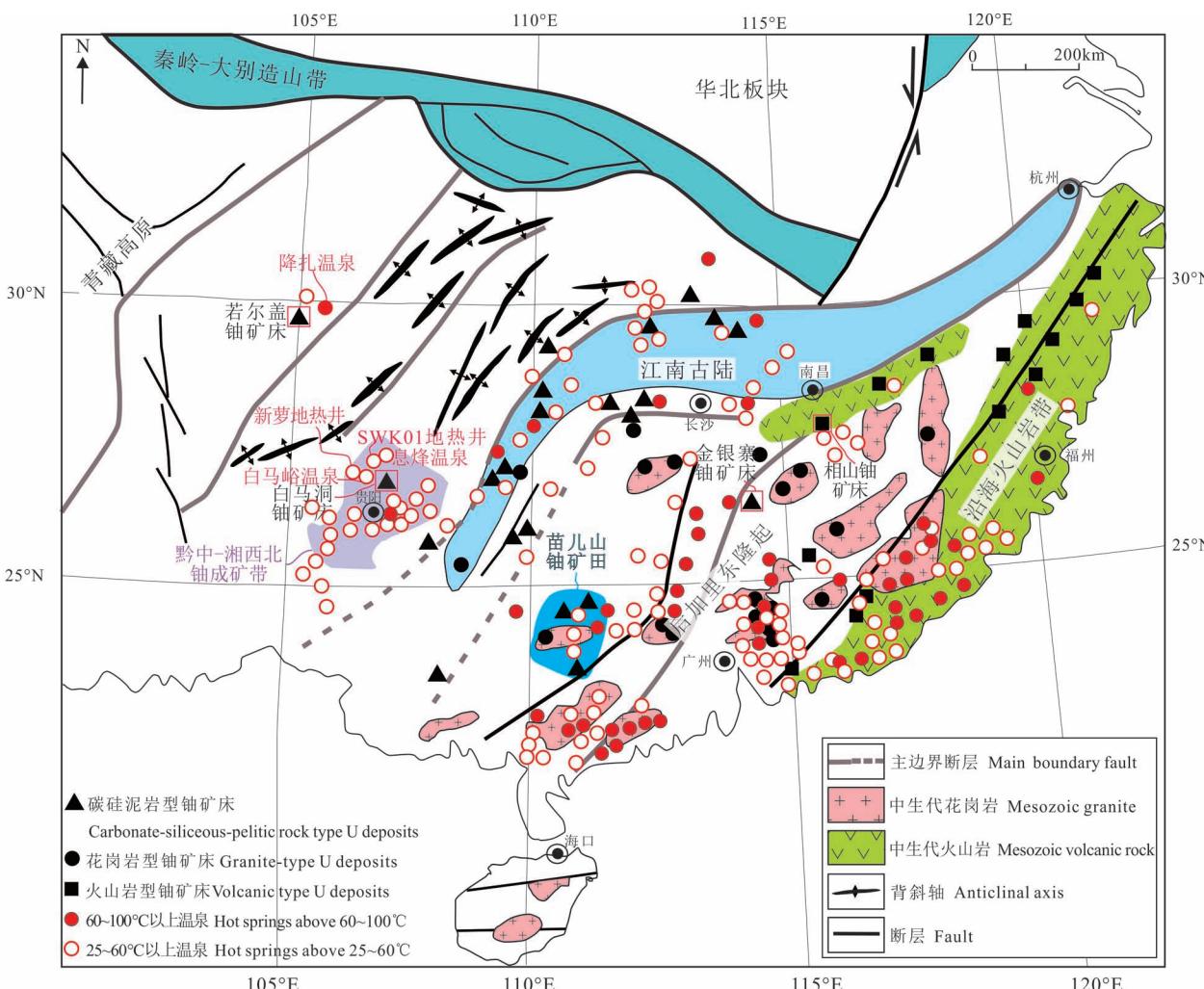


图 1 华南大陆地区温泉与铀矿床(田)分布图(据刘明光, 1998; Hu Ruizhong et al., 2008; 王贵玲, 2018)

Fig. 1 The distribution of hot springs and uranium deposits in south China continent
(After Liu Mingguang, 1998#; Hu Ruizhong et al., 2008; Wang Guiling, 2018#)

孔,并且有自流现象(彭辉才和李俊,2015);仁化县城口铀矿田与城口镇大小 6 处温泉相伴生(周立坚,2016),棉花坑铀矿床深部含矿带内存在大量地下热水(祁家明等,2019a)。此外,在四川,若尔盖铀矿田内分布有降扎温泉等多处温泉(陈友良,2008; 祝宏勋等,2008)。还有学者研究表明铀矿田都有优选定位于热隆内部及边缘的趋势,如诸广山的三大铀矿田(城口矿田、长江矿田、百顺矿田)和会昌矿田(杜乐天,2001a)。然而,热隆内部及边缘同时也是温泉、热水分布的重要位置,如江西安远热隆,热隆边部铀成矿潜力巨大,产有白面石铀矿床等多金属矿产,同时有众多的温泉、地热异常点出露(张万良等,2005)。国际上,澳大利亚 Mt. Painter Inlier 地区的 Paralana 热泉东部约 15 m 处发现中等规模的铀矿床,它们同受 Paralana 构造控制,该铀矿床的成矿元素铀来源于 Mount Painter Domain 地区,而该区地热资源非常丰富(Hillis et al., 2004);同时,澳大利亚南部 Cooper 盆地和 Flinders 山脉高放射性花岗岩分布与干热岩地热资源分布一致,花岗岩放射性产热是该区地热资源热的主要来源(Hillis et al., 2004; Goldstein et al., 2009)。

1.2 地热与铀矿成因关系

有学者研究发现,除温泉与铀矿空间上密切相关外,与铀矿化密切相关的热液蚀变特征与地热流体围岩蚀变特征也相似。比如绿泥石化等还原性蚀变,是地热流体与围岩发生水—岩反应的常见蚀变类型(Vidal et al., 2018),同时也是形成铀矿非常有利的还原性蚀变(李学礼,1992; 周立坚,2016)。Brugger 等(2005)研究发现 Mt. Painter Domain 当代地下热水的地球化学特征与区域成矿流体的地球化学特征关系密切,阐明了二者形成流体的相关性(Walker, 1999; 周立坚, 2016)。也有学者注意到华南的一些铀矿床中,U 与 Hg、Sb、As、Tl 密切共生,而在现代地热田中也如此(涂光帜,1990)。比如在云南腾冲地热田,沥青铀矿、铀石与黄铁矿、赤铁矿密切共生,并具异常含量的 Hg、Sb、As(佟伟,1989)。同时,华南铀成矿年龄(70 ~ 140 Ma 和 22 ~ 65 Ma, Hu Ruizhong et al., 2008)与华南地区温泉水的年龄同为新生代(孙占学等,1992)。还有学者注意到成热与成铀对岩石的物理特性要求较为一致,如水云母化、钠长石化、黄铁绢英岩化等岩石因有效孔隙度增大、渗透能力增高、抗压强度降低,从而更易破碎,更利于地热流体和铀成矿流体的运移和储存,对成热和成矿均有利(李学礼, 1992)。除

上述联系外,二者形成的构造条件、出露标高亦有很大的相似性和一致性(徐增亮, 1982; 杜乐天, 2001b; 周立坚, 2016)。

根据这些密切相关性,学者们提出隆起带古地下热水排泄区铀成矿模式、热水成矿模式、古脉状承压泄水区(减压区)铀成矿模式和“铀—水—热”三源结合研究铀矿成因的思路,认为热液型铀矿不仅在形成时与地热异常有关,在形成后还会对其附近的大地热流产生一定影响,形成一定的热异常(李学礼等, 1984, 1999; 李学礼, 1992)。还有学者提出岩石圈伸展体系下大陆热水系统铀矿床成因模式,认为岩石圈伸展导致地温梯度升高,大地热流平均值加大,在形成地热的同时为铀成矿提供热驱动力,同时,因岩石圈伸展作用引发热水流体大规模循环,最终演化为富铀热液流体,为铀成矿提供保障(商朋强, 2007)。此外,学者们还将地热标志作为找矿突破标志之一应用到热液型铀矿床找矿勘查中,尤其在利用铀矿与地热场分布的密切关系指导深部铀矿勘查上发挥了重大作用(郝士胤, 1981; 祝青云和覃正良, 2008; 李庆阳等, 2010; 黄明光等, 2010),认为地下热水的循环深度可能为铀矿床的最大成矿深度(祁家明等, 2019a)。

2 存在问题及解决方法

2.1 存在问题

目前,对于地热与热液型铀矿关系的研究大部分局限于一些相似现象的描述,而关于产生这些相似性的根本原因并未开展详细研究,更未从地热和热液型铀矿床形成的重要媒介——热液流体入手解决这一根本问题。地热和热液型铀矿床的形成离不开热液流体,它们是热液流体在特定地质背景下地质作用的结果,流体的来源、特征及演化等关乎到地热田和矿床的成因机制,是矿床学和地热学研究的热点。因此,揭示并对比深部古地热流体和成矿流体来源、特征及演化是解决二者成因联系问题的关键。具体解决方法如下,并见图 2。

2.2 解决方法

2.2.1 热液矿物流体包裹体地球化学

流体包裹体是古地热流体、成矿流体遗留下来的唯一直接样品,是了解古地热流体和成矿流体物理化学特征的最直接、最有效方法,是解开成矿流体和古地热流体成矿成藏作用机理的密码,可提供地热成藏、热液矿床成因方面诸多有价值信息。由于其研究方法简单易行,适用矿物较为常见,已被广泛

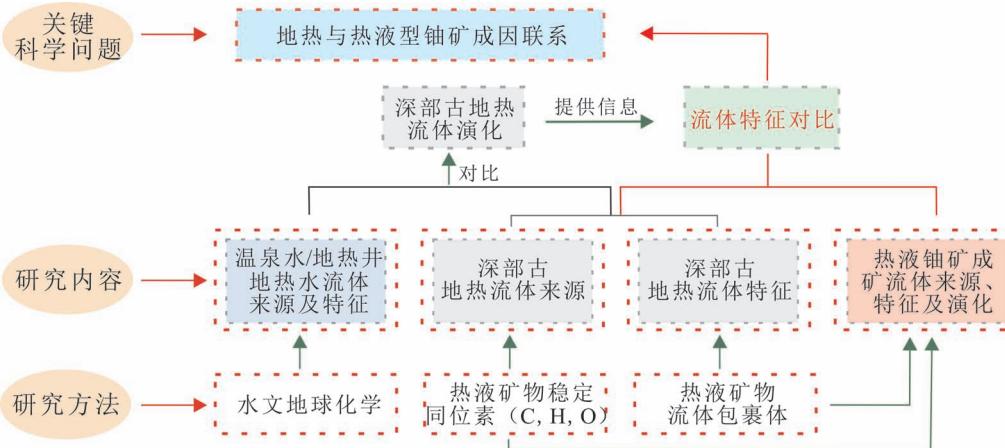


图 2 地热与热液型铀矿关系研究路线图

Fig. 2 Technological roadmap for studying the relations between geothermal energy and hydrothermal uranium deposits

应用到石油地质学、矿床地质学、构造地质学及地球内部流体迁移等研究领域(Bourdet et al., 2008; 孙贺和肖益林, 2009; Chi Guoxiang et al., 2017)。

研究表明井下取样获得地热流体在技术上很难实现(Truesdell and White, 1973),且深部地热流体在上升过程中由于压力骤降会发生闪蒸气损(脱气)作用,同时其在上升到地表的过程中还可能与浅层地下水混合,因此,温泉水/地热井井口获取的地热水地球化学特征仅代表现今的地热流体特征,不能反映深部古地热流体和真实热储特征(Janssen et al., 1997; Kirschner et al., 1999; Budai et al., 2002; Bird and Spieler, 2004)。幸运的是,热液矿物很好地记录了地热系统古地热流体特征、储层演化及水—岩反应等诸多重要信息,所以,想要获得完整的深部古地热流体特征,包括真实温度、压力和气—液组成等,必须对地热系统热液矿物中的流体包裹体开展研究(Belkin et al., 1985)。流体包裹体作为地热学研究的一种非常有效手段(Browne et al., 1976; Elders et al., 1978; Hedenquist, 1980),已被广泛运用到日本 Holi 和 Kirishima 地热田(Sasada, 1986; Sawaki et al., 1997)、中国台湾 Chingshui 地热田(Lu Yichia et al., 2018)、新西兰 Waiotapu 地热田(Hedenquist, 1980)、墨西哥 Los Humeros 地热田(Prol-Ledesma et al., 1989)、菲律宾 Tongonan、Okoy 和 Daklan 地热田等(Leach, 1982; Scott, 1990)。通过流体包裹体显微岩相学、流体包裹体显微测温学(明确深部古地热流体的温度、盐度和压力等)、流体包裹体激光拉曼光谱学(明确深部古

地热流体的气—液组成)和流体包裹体 LA-ICP-MS(明确深部古地热流体的气—液相元素地球化学特征)等可揭示古地热流体特征及演化、真实反映热储条件等,还可圈定地下热水聚集带,为地热勘探开发提供非常重要的指示(Grant et al., 1982; Leach, 1982; Roedder, 1984; Scott, 1990; Boles et al., 2004; Luetkemeyer et al., 2016)。

同时,流体包裹体作为矿床学的基本研究方法,已在厘定花岗岩型铀矿床、火山岩型铀矿床、不整合面型铀矿床、碳硅泥岩型铀矿床等热液铀矿床铀成矿流体特征和演化等中得到广泛应用(Cuney, 1978; Aniel and Leroy, 1985; Pecher et al., 1985; Chu Haixia et al., 2015; 胡旭刚, 2015; Chi Guoxiang et al., 2017; Wang Kewen et al., 2018; 祁家明等, 2019b)。

2.2.2 热液矿物稳定同位素地球化学

同位素地球化学是示踪流体来源的有效手段,比如, $\delta^{13}\text{C}$ 相较与 $\delta^{18}\text{O}$ 相对稳定,主要受水—岩反应和来源控制,因此其成为示踪热液流体来源的重要手段之一(郑永飞等, 2000)。

如前文所述,温泉水/地热井井口获取的地热水地球化学特征仅代表当下地热流体特征,而众多学者研究表明脉石矿物为研究地热系统中地热流体的演化提供了非常重要的记录,尤其是碳酸盐岩类矿物(如方解石等)的同位素特征对指示古地热流体来源、地球化学特征具有重要作用(Wallin and Peterman, 1999; Iwatsuki et al., 2002; Wang Peiling et al., 2010; Li Rongxi et al., 2013; Luetkemeyer et al., 2016; Lu Yichia et al., 2018)。目前,方解石稳定同位素(C, H, O)作为示踪古地热流体来源的有效手段已被广泛应用到中国台湾 Chingshui 地热田(Lu Yichia et al., 2018)、日本 Tono 地区地热系统(Iwatsuki et al., 2002)和美国加利福尼亚州

parkfield 地区地热系统等 (Luetkemeyer et al., 2016)。此外,通过方解石稳定同位素的变化特征还可指示古地热流体作用过程等,有助于揭示古地热流体演化历史 (Wang Peiling et al., 2010)。

上升地壳区,受一系列构造运动的影响(如燕山运动及喜马拉雅运动;余开富等,1995;梅冥相等,2005;何斌等,2005;Ling Kunyue et al., 2015;Long Yongzhen et al., 2017),研究区所在的扬子准

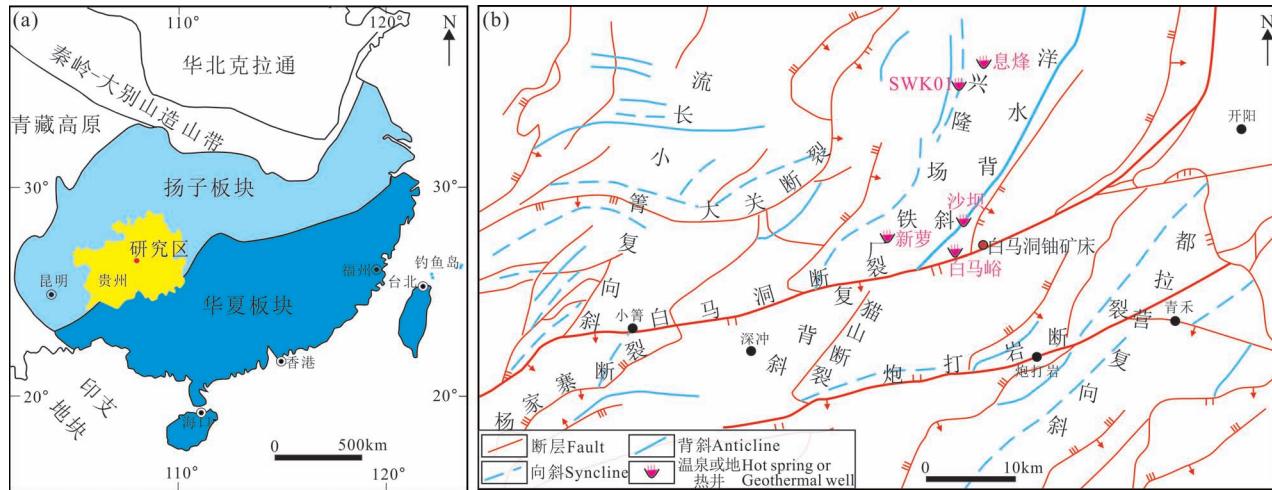


图 3 贵州息烽大地构造位置图(a, 据 Zhou Jiaxi et al., 2014)和构造纲要图(b, 据李燕燕, 2019)

Fig. 3 Regional map showing the tectonic setting and location (a, modified from Zhou Jiaxi et al., 2014) and structural outline map (b, modified from Li Yanyan et al., 2019) of the Xifeng area, Guizhou

同时,方解石等热液矿物稳定同位素(C、H、O)作为厘定铀成矿流体来源的基本研究方法已在花岗岩型铀矿床、火山岩型铀矿床、不整合面型铀矿床、碳硅泥岩型铀矿床等热液型铀矿床中得到广泛应用 (Min Maozhong, 1995; Hu Ruizhong et al., 2008; Liang Rong et al., 2017)。

2.2.3 水文地球化学

研究表明地热流体中某些元素含量的异常可为地热流体的成因等提供重要指示意义。如高 Li、B 含量及 Sr 的出现,暗示地热流体与深部岩浆热源有直接关系 (多吉, 2003);高 Rn 和 He 含量指示浅层局部存在放射性热源;高 U 含量指示这一流体可以通过与有机质或硫化物的还原反应发生铀矿化,从而说明地热流体与成矿流体性质类似,可能为同一流体系统 (Brugger et al., 2005)。因此通过对地热流体开展水文地球化学研究,有助于揭示其与热液铀矿床成因联系。

3 典型地热与热液型铀矿成因联系研究:贵州息烽地热田和白马洞铀矿

3.1 区域概况

贵州位于扬子板块西缘(图 3a),处于东亚中生代造山带与阿尔卑斯—特提斯新生代造山带之间的

地台的黔中隆起发生了不同程度的变化,为铀矿化和地热的形成提供了有利的地球动力学背景和足够的构造空间。全省铀矿和地热资源非常丰富 (毛健全, 2001; 王国坤等, 2018),铀矿床、铀异常点普遍发育,沿区域性断裂带展布,同时,沿断裂构造带伴有众多温泉出露,据统计贵州省温泉及地热井共计 212 处,地热资源共覆盖 72 个区县 (张世丛和陈履安, 1998; 毛健全, 2006; 王贵玲, 2018)。

息烽地热田位于贵阳市北部,地热田内主要分布有白马峪温泉、新萝地热井、息烽温泉等多处温泉 (图 3b),构造位置位于北东向德江—纳雍断裂和东西向白马洞断裂 (F1) 之间,其中白马洞深大断裂沟通深部热源,主要褶皱有兴隆场—铁厂复式背斜、洋水短轴背斜 (图 3b; 韩至钧和金占省, 1996; 李志成和申娴达, 2014)。根据贵州省铀成矿带、铀矿资源远景区以及温泉、地热异常点的出露情况 (张金带等, 2012; 王贵玲, 2018),发现息烽地热田及贵州其它温泉、地热异常点与热液铀成矿带及铀资源远景区在空间上具一致性 (图 1)。研究区热储盖层为寒武系金顶山组 (\mathbb{E}_{1j})、明心寺组 (\mathbb{E}_{1m}) 和牛蹄塘组 (\mathbb{E}_{1n}) 的页岩、砂岩和黏土岩,热储层为震旦系灯影组 (Z_{2dn}) 和陡山沱组 (Z_{1d}) 白云岩 (蔡兴林等, 2018) (图 4)。研究区远离板块边缘,无年轻岩浆体

分布,属板内地热系统,为低地温梯度背景区(Zhang Ligang, 1988)。基于氢、氧同位素组成研究表明,息烽地热田内温泉和地热流体来源为大气降水,认为大气降水以古溶蚀带和层间破碎带等为通道下渗,在下渗过程中深循环加热,热储温度达100℃以上(宋小庆等,2014;吉勤克补子,2015)。受构造控制一部分热流体在热储层中聚集储存,另一部分则沿次生断裂构造带上升出露地表形成温泉,浅部冷水的渗入是温泉温度偏低的重要原因(毛健全和陈阳,1987)。

黔中—湘西北铀成矿带上的白马洞铀矿床位于息烽地热田内,同时息烽地热田内的白马峪温泉、新萝地热井等又出露在白马洞铀矿田内(图3b),因此是研究地热与热液型铀矿成因联系的理想窗口。白马洞铀矿床地理位置位于贵阳市开阳县白马洞乡,大地构造位置位于黔中深断裂、木黄—贵阳—普安断裂和垭都—紫云断裂相交并向其收敛的有利三角构造地带。研究区主要出露寒武系地层,包括牛蹄塘组(ϵ_{1n})、明心寺组(ϵ_{1m})、金顶山组(ϵ_{1j})、清虚洞组(ϵ_{1q})、高台组(ϵ_{2g})、石冷水组(ϵ_{1s})和娄山关群(ϵ_{2-3ls})(图4)(Pi Daohui et al., 2013; Long Yongzhen et al., 2017)。研究区断裂构造发育,其中东西向白马洞断裂是黔中地区的控热控矿断裂,区域自西向东发育有北东向流长一小箐复向斜、兴隆场—铁厂复背斜、都拉营复向斜(莫帮洪等,2016)。

3.2 研究进展

3.2.1 时空及成因联系

空间上,息烽地热田内白马峪温泉、新萝地热井、两岔河地热井等与白马洞铀矿床相毗邻,整个地热田与黔中铀成矿带分布一致(图1);时间上,白马洞铀矿床铀成矿年龄(41.7~51.5 Ma, 44.7~51.5 Ma, 27.8~28.9 Ma, 37.3~37.5 Ma 和 20~28.9 Ma, 唐俊儒和朱杰辰, 1986; 陈露明, 1990; 黄凯平等, 2011)与息烽地热田温泉水形成年龄(9466 ± 473 a, 吉勤克补子, 2015)均为新生代,尽管息烽地热田热储层(蔡兴林等, 2018)和白马洞铀矿含矿层不同(李燕燕, 2014; Li Yanyan et al., 2019)(图4)。但基于二者的时空关

系,暗示它们成因上的关联性。

何文君等(2014)的研究认为息烽地热田内新萝地热井地热水中的氡来源于深部岩石裂隙中的氡气,而该氡气的来源与白马洞铀矿床铀矿层有一定联系。研究发现,息烽地热田内温泉和地热的形成与白马洞铀矿形成的地质背景和机理较为一致,均受白马洞深大断裂控制,且大气降水是二者形成的重要流体来源(宋小庆等, 2014; 李燕燕, 2019)。同时,息烽地热田热储温度(大于100℃, 宋小庆等, 2014)和白马洞铀矿床成矿流体温度(120~180℃, 李燕燕, 2019)相似,均为中—低温,说明二者形成流体可能具一致性。此外,更值得注意的是研究区围岩蚀变强烈,其中白马洞铀矿床热液蚀变

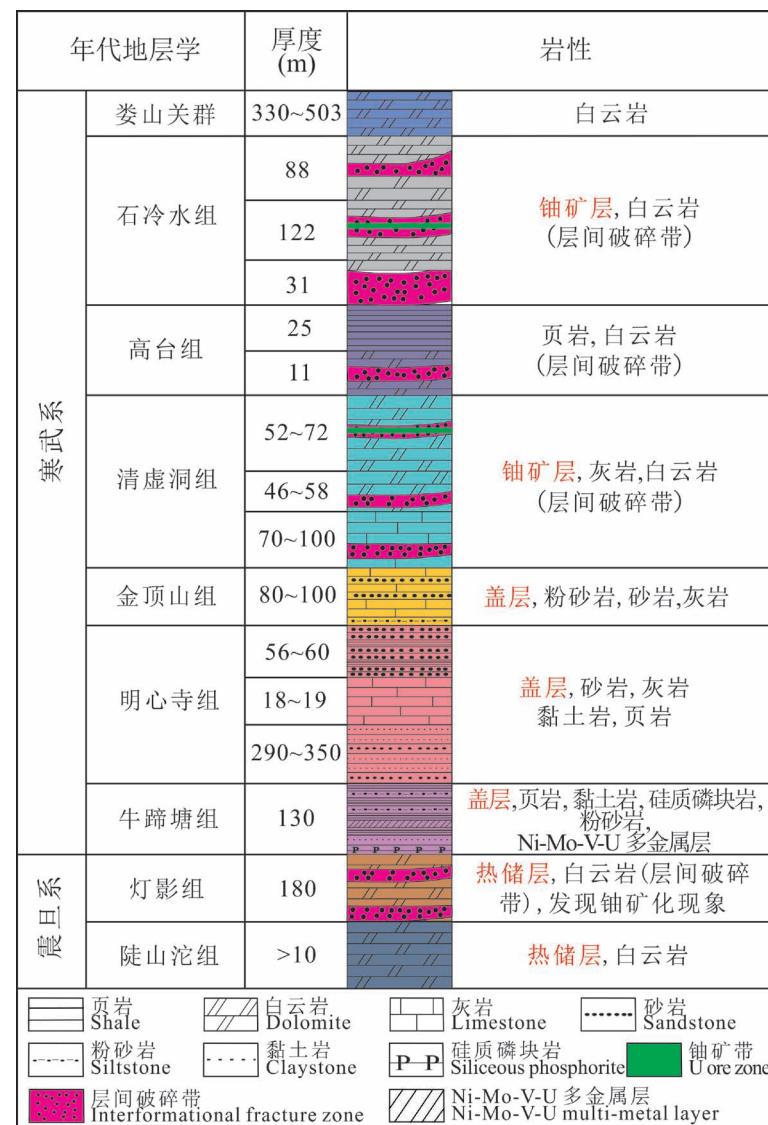


图4 贵州息烽地区地层柱状图(据张小强等, 2017修改)

Fig. 4 The stratigraphic column of the Xifeng area, Guizhou
(Modified from Zhang Xiaoqiang et al., 2017)

种类多,包括一系列的还原蚀变和氧化蚀变,而其中对铀矿化非常有利的还原蚀变也是地热流体与围岩发生水—岩反应常见的蚀变类型(李学礼, 1992; Li Yanyan et al., 2019),进一步说明二者成因上可能具有联系。

3.2.2 成因联系探讨

白马洞铀矿床铀来源于牛蹄塘组(Li Yanyan et al., 2019),牛蹄塘组是研究区出露的最为富铀的地层(Coveney et al., 1991; 李朝阳等, 2003; 陈兰, 2005; Yang Enlin et al., 2013),该地层铀含量($43.23 \mu\text{g/g}$)是地壳沉积岩平均铀含量的10倍,是地壳丰度值的20倍(倪师军等, 2009),同时,研究区其它地层铀元素含量也非常丰富,如热储层灯影组($U = 12.4 \mu\text{g/g}$, 杨瑞东等, 2014)。牛蹄塘组黑色岩系长期处于还原环境(陈兰, 2005),铀在还原环境下是以 U^{4+} 形式稳定存在的,只有在氧化环境下才以 U^{6+} 形式活化迁移,因此这些铀可以长期保存,为大规模的铀成矿作用和放射性产热提供物质准备。

研究区经历了多期热液活动。首先,早期的热液活动使得富 U 牛蹄塘组中的海相有机质在深部热液流体的驱动下沿断裂至清虚洞组和石冷水组层间破碎带圈闭沉淀(形成还原环境),同时,富硅热液流体携带 U 、 Hg 、 Mo 沿断裂向上运移,使得 U 初步富集(李燕燕, 2019; Li Yanyan et al., 2019)(图5a);随后,因研究区长期处在拉张构造背景下(Hu Ruizhong et al., 2008),加之F1断裂构造不断扩充及加强,角砾岩带的发育以及黔中地区多次隆升,研究区地层已经很破碎,渗透性已很高,极大地促进了大气降水的深部循环,古大气降水下渗深度较周缘地区为最大(魏肖等, 2018),循环深度约3 km(宋小庆等, 2014; 吉勤克补子, 2015),大气降水只需深循环 $2 \sim 3 \text{ km}$ 即可演化为成矿热液(严冰, 2012),认为白马洞铀成矿流体与息烽地热田地热流体可能为同一热液流体系统—古大气降水热液系统。氧化性的大气降水($\text{H}_2\text{O} + \text{O}_2$)沿断裂、裂隙及层间破碎带下渗,活化迁移流经的富 U 岩石(例如灯影组白云岩),使得 $\text{U}^{4+} \rightarrow \text{U}^{6+}$ 活化迁移($\text{U}^{4+} \rightarrow \text{UO}_2^{2+}$), U^{6+} 主要以 $[\text{UO}_2(\text{CO}_3)_3]^{4-}$ 、 $[\text{UO}_2(\text{CO}_3)_2]^{2-}$ 形式迁移。这种携带 U^{6+} 的氧化性流体遇到早期热液活动在清虚洞组和石冷水组圈闭创造的还原环境时, U^{6+} 被还原为 U^{4+} ,发生 U 的沉淀,同时 Fe^{2+} 被氧化为 Fe^{3+} ,赤铁矿和针铁矿沉淀,铀矿化发生即 $\text{U}^{6+} + \text{Fe}^{2+} \rightarrow \text{U}^{4+} + \text{Fe}^{3+}$ (图5b)。其中, Fe^{2+} 主要来源

于早期热液活动形成的黄铁矿(Fe_2S)(李燕燕, 2019; Li Yanyan et al., 2019)。

3.2.3 铀矿体放射性产热机理

一般情况下,温度函数 $T(x, y, z)$ 的热传递定解为:

$$\nabla^2 T(x, y, z) = \frac{\partial^2 T}{\partial x^2} + \frac{\partial^2 T}{\partial y^2} + \frac{\partial^2 T}{\partial z^2} \\ = - \frac{A(x, y, z)}{K} \quad (1)$$

$$T(x, y, z) | \Gamma_1 = T_0 \quad (2)$$

$$\left. \frac{\partial T(x, y, z)}{\partial n} \right| \Gamma_2 = \frac{q}{K} \quad (3)$$

其中, $A(x, y, z)$ 是生热率分布函数, K 为热导率, q 为大地热流, Γ_1 为第一类温度边界, Γ_2 为第二类边界条件的热流量分布函数。据此,李庆阳等(2010)设计了如图6a所示的铀矿体放射性产热模型,采用边界单元数值模拟,通过计算获得如图6b所示的结果。根据计算结果发现因热源的存在对地热空间分布产生了明显影响,在热源处引起热场畸变,即在热源所在空间附近地温梯度产生陡变,这是铀矿体产热的本质,铀矿体可等同于一个热源(李庆阳, 1992; 李庆阳等, 2010)。

此外,根据Rybarch(1976)修正的岩石放射性产热与岩石中的放射性元素含量的关系方程:

$$\frac{A}{\mu\text{W}/\text{m}^3} = 0.01 \frac{\rho}{\text{g}/\text{cm}^3} \times \\ \left(3.48 \frac{w_{\text{K}}}{\%} + 9.52 \frac{w_{\text{U}}}{\mu\text{g/g}} + 2.56 \frac{w_{\text{Th}}}{\mu\text{g/g}} \right) \quad (4)$$

式中 A 为岩石放射性产热率, ρ 为岩石密度, w_{K} 、 w_{U} 、 w_{Th} 分别为岩石中 K 、 U 和 Th 含量。显然,岩石生热率与岩石密度 ρ 以及 U 、 Th 、 K 含量四个独立变量有关,而各类岩石的密度相差不大,除角闪岩为 $3.031 \text{ g}/\text{cm}^3$ 外,其它各类岩石均在 $2.402 \sim 2.796 \text{ g}/\text{cm}^3$ 之间,因此,岩石中 U 、 Th 、 K 越富集,岩石的产热率越大,而根据公式(4),三种元素中 U 元素对生热贡献最大,因此 U 元素越富集,生热率越高,且各类岩石只要含有放射性元素,就会不断的释放热量(赵平等, 1995)。

白马洞铀矿床规模中等、品位较高,铀矿体赋存在清虚洞组和石冷水组的灰岩和白云岩中(李燕燕, 2014)。综上,铀矿化发生后,清虚洞组和石冷水组可作为浅部放射性热源为地表出露的温泉和热异常提供源源不断的热,牛蹄塘组作为深部放射性热源为深部的灯影组和陡山沱组热储提供热源(图



图 5 贵州息烽地热田白马洞地区地热与铀矿成因关系模式图

Fig. 5 Genetic model of relationship between geothermal energy and uranium deposit in the Baimadong area, Xifeng geothermal field, Guizhou

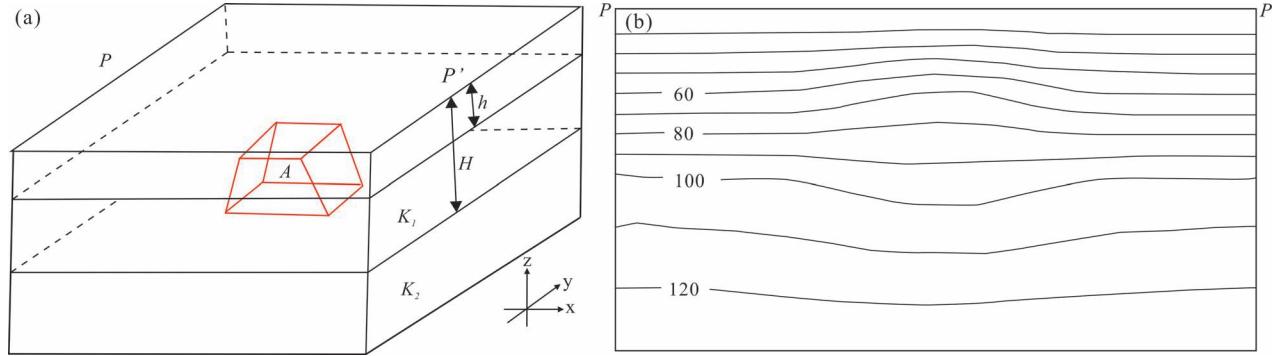


图 6 (a) 铀矿体产热计算模型;(b) 模型对应的计算结果(据李庆阳等, 2010)

Fig. 6 (a) Computational model for heat production of uranium ore body; (b) computational result of the model in Fig. 6(a) (from Li Qingyang et al., 2020&)

5b)。含铀丰富的铀矿体和岩石不停的释放热量, 生热率较大, 通过放射性产热对其附近的大地热流产生一定的影响, 形成一定的热异常, 进而形成丰富的地热资源。息烽地热田地热和白马洞铀矿床不仅在形成过程中密切联系, 在形成后也彼此成就。在地热和铀矿形成时, 古大气降水下渗深循环加热成为地热流体, 为铀的活化迁移提供热动力条件, 是启动铀矿化的关键因素之一, 同时, 因活化迁移搬运了 U^{6+} , 使得 U 富集成矿, 其还扮演着成矿流体的角色。在铀矿床形成后, 铀矿体起到放射性热源的作用, 为深部热储和地表温泉、热异常提供源源不断的热。

4 结论

本文通过总结前期研究工作和野外调查观测, 同时收集和整理前人研究资料, 详细阐述了地热与铀矿的时空密切相关性, 并阐明了解决其成因联系的研究方法。结合前人及已有研究成果, 认为地热流体与铀成矿流体可能具有相似性, 或属于同一热液流体系统。地热流体为铀活化迁移成矿提供充足的热动力条件, 是铀迁移富集最终成矿的运输媒介, 同时通过铀矿化铀元素的富集而提供源源不断的放射性热, 二者不仅在形成时密切相关, 在形成后也相互成就。但目前的研究主要集中在地热田内温泉、地热井地热水的水文地球化学研究, 而对热—铀有时空关系的地热田、热异常区内的热液矿物开展的

研究较少, 因此是接下来研究地热与热液型铀矿成因联系研究的重点。通过地热与热液型铀矿成因联系研究可实现“热—铀兼探”和“热—铀兼采”, 为深部地热、干热岩和固体矿产综合找矿, 即“热矿共采”提供一定的理论依据。

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Genetic relationship between geothermal energy and hydrothermal uranium deposits: research progress and method

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Abstract: As a type of clean and renewable resource, geothermal energy has been widely utilized in power generation and heating. Radiogenic heat is a primary source for geothermal energy in earth. The research on radiogenic geothermal energy has been focused on its application in detecting and evaluating potential for dry hot rock. However, as vital radioactive source, the research about the genetic relationship between hydrothermal uranium deposits and geothermal energy is poor. This paper collected and sorted out a large number of domestic and foreign literature as well as related materials to enhance the research about the genetic relationship between geothermal energy and hydrothermal uranium deposits. Based on the research progress and existing problems, method to solve this issue is put forward. Through this study, a new thought and method can be provided for studying the genetic relationship of geothermal energy and hydrothermal uranium deposits and certain theoretical basis could be supplied for “geothermal energy—ore jointly exploring”.

Keywords: geothermal energy; hydrothermal uranium deposits; genetic relationship; research progress; method

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