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The Lower Mantle Minerals in Ophiolite-hosted Diamond

YANG Jingsui^{1*}, WIRTH Richard², XIONG Fahui¹, TIAN Yazhou¹, HUANG Zhu¹,
ROBINSON Paul T.¹ and DILEK Yildirim³

1 *CARMA, Institute of Geology, Chinese Academy of Geological Sciences, Beijing 100037, China*

2 *Helmholtz Centre Potsdam, GFZ German Research Centre for Geosciences, Germany*

3 *Department of Geology, Miami University, Oxford, Ohio, USA*

Ophiolites are fragments of ancient ocean lithosphere emplaced on continental margins, in island arcs or in accretionary prisms, and have long been studied to better understand the evolution of ocean basins and collision of tectonic plates, the processes of mountain building and the occurrence of valuable ore bodies, such as podiform chromitites (Dilek and Furnes, 2011). Because these bodies contain both upper mantle and crustal rocks of extensional origin, their magmas have traditionally been thought to form at shallow partial melting levels above spreading axes in the open ocean or above subduction zones. The podiform chromitites are currently interpreted as magmatic deposits formed during melt-rock reaction in MOR or suprasubduction zone (SSZ) environments (Zhou et al., 1994; Arai, 1997). However, discoveries of in-situ diamonds and other ultrahigh pressure (UHP) minerals, highly reduced phases and native elements, along with many crustal minerals in ophiolites and chromitites of southern Tibet and northern Ural, Russia, strongly suggest that these bodies formed at mantle depths of 150-300 km or greater, perhaps near the mantle transition zone (Bai et al., 2003; Yang et al., 2007; 2014, 2015; Xu et al., 2009, 2015; Dobrzhinetskaya et al., 2009, 2014; Yamamoto et al., 2009; Trumbull et al., 2009; Robinson et al., 2015).

Recently, diamond, moissanite and other unusual minerals were also reported from the podiform chromitite in the Hegenshan ophiolite of Inner Mongolia and the Sartohai ophiolite of Xinjiang (Tian et al., 2015; Huang et al., 2015). Both are Late Paleozoic (360-400 Ma) in age and occur along the Central Asian Orogenic Belt (CAOB), which is a huge complex extending E-W for over 4000 km and marks the boundary between the Siberian Block to the north and North China Block to the south. Here we report the first discovery of Ca-silicate perovskite as mineral

inclusions in diamond from podiform chromitite in the Hegenshan ophiolite of Inner Mongolia and the Sartohai ophiolite of Xinjiang.

The mineral inclusions were identified in the foil #4070 by using FIB and TEM analyses at the GFZ, Germany. Ca-silicate perovskite in diamond from the Sartohai chromitite occurs as a 60 nm granular euhedral grain, associated with an equal size alloy. Both occur at the margin of a large graphite inclusion in diamond. EDS shows one is NiMnCo alloy (Ni 68.01, Mn 25.71, Co 6.27), the other is an oxide, with Si= 47.02at%, Ca= 42.19at% and Mn=10.78at%. The calculated mineral formula of the oxide is $\text{Ca}_{0.8}\text{Mn}_{0.2}\text{SiO}_3$, Mn-bearing Ca-silicate perovskite. The electron diffraction patterns can be indexed based on an orthorhombic symmetry of Ca-silicate perovskite and cubic structure of the alloy.

Ca-silicate perovskite (foil #4069) in diamond from Hegenshan chromitite occurs as a 50 nm grain within a large inclusion containing both Ni-Mn-Co alloy and Nd-Se-Cu-S phase. By EDS the chemical composition of the perovskite is Ca=48.3at%, Si=37.7at% and Mn=14.1at% plus oxygen. TEM diffraction data show that the Hegenshan inclusion has d-spacings and angles between adjacent lattice planes consistent with Ca-silicate perovskite with an orthorhombic structure (Li et al., 2006). Comparing to the research by others (Komabayashi et al., 2007), the formational P-T should be $T > 1600^\circ\text{C}$, $P > 20$ GPa.

CaSiO_3 perovskite (CaPv) is an important mineral in both transition zone and lower mantle. Its modal mineral proportions are supposed to be about 5 and 25vol% in pyrolitic mantle and subducted mid-oceanic ridge basalt (MORB) materials, respectively, under the lower mantle conditions (e.g., Mao et al., 1989; Hirose et al., 2005).

* Corresponding author. E-mail: yangjsui@cags.ac.cn

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