

豫西吉家洼金矿床成矿时代和成矿物质来源： 来自闪锌矿 Rb-Sr 同位素年龄和 Pb 同位素的证据

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内容提要: 豫西吉家洼金矿位于熊耳山金多金属矿集区西部, 是一构造蚀变岩-石英脉型金矿床。矿体产于近南北向断裂带中, 在倾向上呈“Y”字型分布, 矿化以细脉-网脉状、浸染状黄铁矿化蚀变岩型金矿为主, 其次为石英脉型金矿。热液成矿过程包括 4 个矿化阶段, 黄铁矿-石英阶段(I)、石英-黄铁矿阶段(II)、石英-多金属硫化物阶段(III)和石英-碳酸盐阶段(IV)。本次工作对吉家洼金矿床主要成矿阶段(III)的闪锌矿进行了 Rb-Sr 同位素测年研究, 以限定成矿时代。结果获得 Rb-Sr 等时线年龄为 118.2 ± 2.4 Ma, 表明矿床形成于早白垩世。该年龄与熊耳山地区中的庙岭、祁雨沟、公峪等金矿床的成矿年龄基本一致, 对于整个熊耳山地区的金矿成矿年龄具有一定的约束意义。铍-铅同位素研究结果显示, 吉家洼金矿的成矿物质为壳幔混源, 新太古界太华群基底及其重熔形成的花岗岩体可能为铅同位素的主要物源。我们认为, 吉家洼金矿床是熊耳山地区在早白垩世区域性强烈的构造-岩浆-流体活动事件的产物, 是整个中国东部金的大规模成矿作用的组成部分。

关键词: 闪锌矿 Rb-Sr 同位素测年; Pb 同位素; 成矿时代; 成矿物质来源; 吉家洼金矿; 熊耳山

精确测定金属矿床的成矿时代, 对于正确认识矿床成因和控矿因素、总结成矿规律并指导找矿勘探工作都具有极为重要的意义 (Zhai Yusheng et al., 2008; Liu Jianming et al., 1998a; Mao Jingwen et al., 2000, 2005)。目前, 金矿定年中最常用的方法有石英流体包裹体的 Rb-Sr 等时线法, 单颗粒热液锆石 U-Pb 法 (Wei Junhao et al., 2003), 含钾矿物的 $^{40}\text{Ar}/^{39}\text{Ar}$ 法 (Wang Yitian et al., 2001, 2002; Luo Zhenkuan et al., 2002; Zhang Yongmei et al., 2011; Zhai Lei et al., 2012), 硫化物 Rb-Sr 等时线法 (Brannon et al., 1992; Nakai et al., 1993; Christensen et al., 1995a, b; Pettke et al., 1996; Han Yigui et al., 2007)。熊耳山地区是华北地块南缘重要的金多金属矿集区。自 20 世纪 70 年代以来, 区内陆续发现大、中型金矿床 10 多个, 并发现雷门沟、鱼池岭大型钼矿和沙沟、铁炉坪、蒿坪沟等一系列中、小型银铅锌多金属矿床。前人对区内的金矿床做过大量的研究, 在金矿成矿时代 (Wang

Yitian et al., 2001; Mao Jingwen et al., 2005; Zhai Lei et al., 2012)、成矿物质来源 (Fan Hongrui et al., 1993, 1994; Mao Jingwen et al., 2002; Wang Tuanhua et al., 2009; Wang Weixing et al., 2010)、成矿流体来源 (Chen Yanjing et al., 1992, 2001; Lu Xinxiang et al., 2002; Zhang Xingkang et al., 2014)、矿床成因和成矿规律 (Fan Hongrui et al., 1998; Xiong Suofei et al., 2013; Ye Huishou et al., 2016) 等方面取得了重要成果。

吉家洼金矿位于河南省洛宁县境内, 是熊耳山金多金属矿集区一典型的蚀变岩-石英脉型金矿床, 以往对该金矿床的矿石特征、矿物特征、构造特征、金的赋存状态和成矿流体来源等方面做了研究 (Wu Changhang et al., 1998, 2008; Yan Zhengxin et al., 2012; Zhang Xingkang et al., 2014)。但对成矿时代的研究比较薄弱, 这在一定程度上制约了对该矿床及熊耳山地区成矿规律的认识。本文在系统总结矿床地质特征的基础上, 选取主要成矿阶段的闪锌

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矿开展了 Rb-Sr 同位素定年和 Pb 同位素的测定工作,厘定了吉家洼金矿床的成矿时代,探讨了成矿物质来源,为矿床成因和成矿规律研究提供了制约。

1 熊耳山金多金属矿集区地质特征

熊耳山金多金属矿集区位于华北地块南缘,南部紧靠北秦岭构造带,南以马超营断裂为界,西北部以洛宁断陷盆地为界,东北部以三门峡-宝丰断裂为界。区内主要出露新太古界太华群结晶基底和由中元古界熊耳群火山岩和少量的中元古界官道口群碳酸盐岩等组成的盖层。此外,在断陷盆地内还发育中新世代红色碎屑沉积岩(Guo Baojian et al., 1997, 2005; Li Yongfeng et al., 2006)(图 1)。褶皱构造有曹嘴沟复式向斜构造(基底褶皱)和龙脖-花山背斜(盖层褶皱)。区域上发育变质核杂岩构造,拆离断层沿新太古界太华群结晶基底与上覆的中元古界熊耳群火山岩的不整合面展布(Shi Quanzeng et al., 1993; Zhang Jingjiang et al., 2003; Sun Weizhi et al., 2013)。

矿集区断裂构造发育明显。断裂构造主要有 4 组:以近东西向、北东向为主,其次为近南北向和北西向。主要区域性断裂有洛宁山前断裂带和马超营断裂带。前者呈北东向,为熊耳山区北界,属上拆离盘犁式断裂;后者呈近东西向,为熊耳山区南界犁式逆冲断层系,具有长期活动的历史。北东向断裂广泛发育,主要有上宫-星星印断裂、陶村-红庄断裂、伊川-潭头断裂等,在平面上呈近等间距分布,是矿集区的主要导矿构造。

矿集区岩浆岩主要有 4 期,新太古代中基性火山岩和 TTG 岩系,经区域变质作用形成本区结晶基底太华群各类片麻岩和花岗片麻岩;中元古代(1800~1650Ma)熊耳群裂谷型中基性—中酸性火山岩;三叠纪碱性岩,如磨沟正长岩年龄为 $237.6 \pm 2.9 \sim 210.4 \pm 2.0$ Ma(Cao Jing et al., 2015);晚侏罗世—早白垩世,五丈山花岗岩基,其 SHIMP 锆石 U-Pb 年龄为 156.8 ± 1.2 Ma(Mao Jingwen et al., 2005);花山复式岩基,其 SHIMP 锆石 U-Pb 年龄为 $127.6 \pm 1.3 \sim 132.0 \pm 1.6$ Ma(Mao Jingwen et al., 2005; Li Yongfeng et al., 2005; Meng Fang et al., 2012; Xiao E et al., 2012);雷门沟花岗斑岩,其 SHIMP 锆石 U-Pb 年龄为 131.1 ± 0.6 Ma(Cao Jing et al., 2016);雷门沟石英斑岩 U-Pb 年龄为 127.2 ± 1.4 Ma(Chen Xiaodan et al., 2011),店房花岗斑岩 U-Pb 年龄为 142.6 ± 2.1 Ma(Tian Yongfei et

al., 2016)等。

矿集区出露众多的金、钼、银铅锌等矿床,主要分布于燕山期花山复式岩基外围、太华群片麻岩与熊耳群火山岩之间的拆离断层附近、近东西向断裂带与北东向断裂带的交汇部位。其中金矿床包括隐爆角砾岩型金矿(祁雨沟金矿和店房金矿),构造蚀变岩型金矿(上宫金矿、庙岭金矿),蚀变岩-石英脉复合型金矿(康山金矿、吉家洼金矿),金矿床的成矿年龄主要集中在早白垩世,如庙岭金矿成矿年龄为 117.0 ± 1.6 Ma(Zhai Lei et al., 2012)、祁雨沟金矿成矿年龄为 $126 \pm 11 \sim 115.1 \pm 1.5$ Ma(Wang Yitian et al., 2001; Han Yigui et al., 2007),公峪金矿成矿年龄为 122.6 ± 0.6 Ma(Li Li et al., 2002)、上宫金矿成矿早、中、晚阶段成矿年龄分别为 242 ± 10 Ma、 165 ± 7 Ma 和 113 ± 6 Ma(Chen Yanjing et al., 2004)。钼矿床包括早白垩世斑岩型钼矿(雷门沟钼矿、石窑沟钼矿)(Li Yongfeng et al., 2006; Gao Yalong et al., 2010),三叠世石英脉型钼矿(前范岭钼矿、纸房钼矿)(Gao Yang et al., 2010),晚三叠世碳酸岩脉型钼矿(黄水庵钼矿)(Cao Jing et al., 2014)。此外,区内还分布有早白垩世脉状银多金属矿床(铁炉坪银铅矿、沙沟银铅锌矿、嵩坪沟银铅锌矿)等(Ye Huishou et al., 2006; Gao Jianjing et al., 2011)(图 1)。

2 矿床地质特征

吉家洼金矿床位于河南省洛宁县底张乡,地理坐标:东经 $111^{\circ}27'46'' \sim 111^{\circ}29'46''$ 、北纬 $34^{\circ}10'08'' \sim 34^{\circ}12'34''$;现已查明金矿矿石量达 185.7×10^4 t 以上,资源/储量 9939kg 以上,矿床平均品位 5.35g/t,属中型金矿床(Yan Zhengxin, 2012)。

矿区地处华北陆块南缘熊耳山地区,花山-龙脖背斜核部南侧,瓦庙河次级倾伏向形的东翼。矿区出露地层主要有新太古界太华群和中元古界熊耳群。新太古界太华群的岩性主要有黑云斜长片麻岩、斜长角闪片麻岩、斜长角闪岩及浅粒岩等。中元古界熊耳群的岩性主要为杏仁状安山岩,斑状安山岩、玻基安山岩及英安岩等(图 2)。构造主要发育近南北向、北东向 2 组断裂带,是区内主要控矿构造。近南北向断裂带主要由 F_1 、 F_2 、 F_3 、 F_4 、 F_6 、 F_7 和 F_8 等断层组成,近平行展布于东西宽约 400m 的狭长地带,断裂走向近南北,倾向 $80^{\circ} \sim 110^{\circ}$,倾角 $65^{\circ} \sim 85^{\circ}$,局部近直立,长度 125~480m,宽度 0.2~2.0m,具分枝复合的特点。北东向断裂带主要有

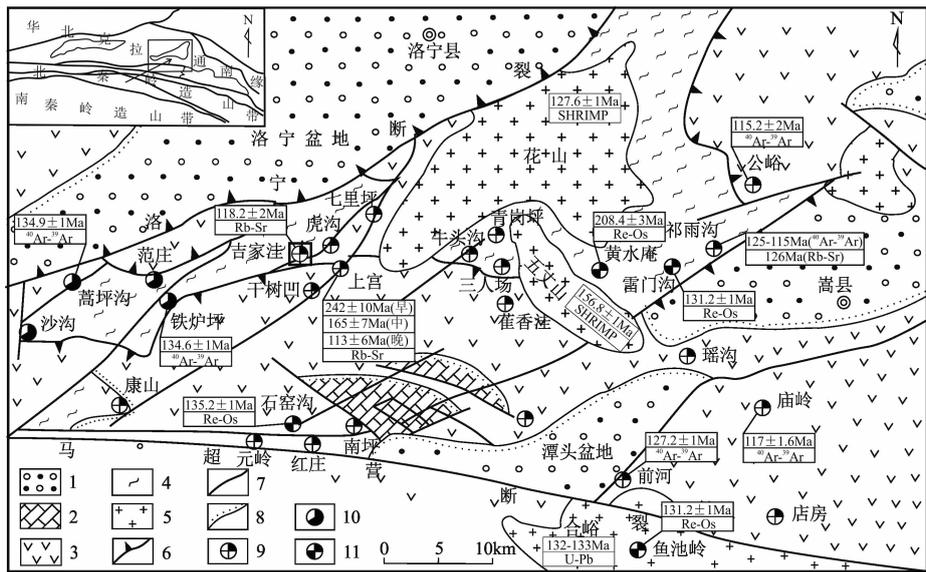


图 1 熊耳山地区地质简图及主要金属矿床分布(据 Guo Baojian et al., 2005 修改)
 Fig. 1 Geological sketch map and the distribution of ore deposits in Xiong'er mountain area
 (modified after Guo Baojian et al., 2005)

- 1—中生代沉积岩; 2—中元古界官道口群白云质大理岩; 3—中元古界熊耳群火山岩; 4—新太古界太华群片麻岩;
 5—中生代花岗岩; 6—拆离断层; 7—断层; 8—不整合界线; 9—金矿床; 10—银铅锌矿床; 11—钼矿床

- 1—Meso-Cenozoic sediments; 2—dolomite-marble of Middle Proterozoic Guandaokou Group; 3—volcanic rocks of the Mesoproterozoic Xiong'er Group; 4—Neoproterozoic gneiss of the Taihua Group; 5—Mesozoic granitoid; 6—detachment fault;
 7—fault; 8—unconformity; 9—Au deposit; 10—Ag-Pb-Zn deposit; 11—Mo deposit

F_9 、 F_{10} 、 F_{12} 、 F_{13} 等 4 条, 断裂走向 $25^\circ \sim 60^\circ$, 倾角 $55^\circ \sim 80^\circ$, 断裂面平直光滑, 带内主要充填蚀变碎裂岩、糜棱岩及少量的断层泥。侵入岩主要为辉绿岩脉。

金矿体主要赋存于近南北向断裂带中, 其中 I、II、III 号矿体为矿区主要工业矿体, 分别赋存于 F_1 、 F_2 、 F_3 等含矿断裂中。金矿体呈薄脉状、透镜状或不规则状, 在倾向上呈“Y”字型产出(图 3), 倾角 $65^\circ \sim 85^\circ$, 局部近直立, 部分矿段出现反倾现象, 在深部归并成一条矿体, 矿体规模明显增大, 厚度增加, 但品位有所下降, 局部有夹石存在。

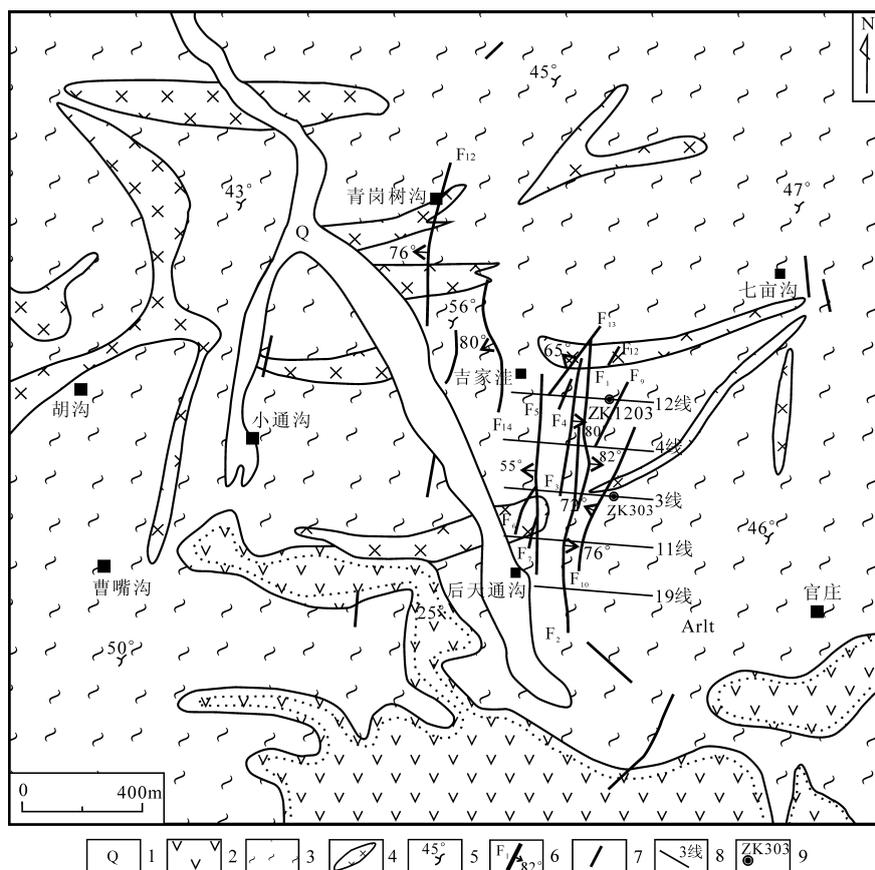
矿区金矿化类型有构造蚀变岩型和石英脉型。构造蚀变岩型金矿主要以黄铁绢英岩化蚀变岩、黄铁矿化蚀变岩、黄铁矿化碎裂岩为主。矿石矿物主要为黄铁矿, 其次为黄铜矿、方铅矿、闪锌矿、黝铜矿、自然金、碲金矿等; 脉石矿物主要为石英、绢云母、钾长石, 次为方解石、铁白云石、绿泥石、高岭石等。矿石构造为浸染状构造、碎裂构造、角砾状构造、块状构造等, 矿石结构多为自形一半自形粒状结构、他形粒状结构、包含结构, 为本区主要的金矿化类型。石英脉型金矿主要以多金属硫化物石英脉为主, 矿石矿物主要有闪锌矿、方铅矿, 其次为黄铜矿、

黄铁矿、自然金、银金矿、碲金矿等; 脉石矿物主要为石英, 其次为方解石、铁白云石、绢云母、萤石等。矿石构造为脉状构造, 梳状构造、晶洞构造等, 矿石结构主要为自形一半自形粒状结构、他形粒状结构, 为本区重要的金矿化类型。两种金矿化类型在垂向上呈渐变分带的关系, 大致以 600m 中段为界限, 上部为石英脉型金矿化为主, 下部为构造蚀变岩型金矿化为主。

矿区热液蚀变主要有钾长石化、硅化、绢云母化、黄铁绢英岩化、绿泥石化、碳酸盐化、萤石化等。其中硅化、绢云母化、黄铁绢英岩化与金矿成矿关系密切。蚀变岩多沿构造带及其两侧围岩呈线状或带状分布, 厚度变化较大, 从几厘米到数米不等, 与围岩呈渐变过渡关系。围岩蚀变略见分带, 从内到外蚀变强度逐渐减弱, 金品位逐渐降低。

根据矿脉穿插关系以及矿石矿物共生组合、结构构造, 可将热液成矿过程划分为以下 4 个成矿阶段。

第 I 阶段: 黄铁矿-石英阶段。常形成黄铁矿化石英脉, 含钾长石石英脉, 常呈脉状、透镜状充填于构造带中, 厚度一般十几厘米。石英多为乳白色, 油脂光泽, 他形粒状结构, 粒径较大, 多为 1~5mm, 最

图2 吉家洼金矿地质简图^①Fig. 2 Geological sketch map of the Jijiawa Au deposit^①

1—第四系;2—中元古界熊耳群火山岩;3—新太古界太华群片麻岩;4—辉绿岩脉;5—片麻理产状;6—含矿断裂带(矿脉)及编号;

7—断层;8—勘查线及其编号;9—钻孔位置及其编号

1—Quaternary sediments;2—Xiong'er Group in the middle Proterozoic;3—Taihua Group in the Archeozoic;

4—diabase dike;5—attitude of gneissosity;6—fault zone with ore (ore vein) and their numbers;7—fault;

8—exploration lines and their numbers;9—positions of borehole and their numbers

大粒径可达10mm。黄铁矿呈自形—半自形粒状结构,以立方体晶形居多,颗粒较大,呈星点状、浸染状分布于脉石英中。围岩蚀变主要为钾长石化、黄铁矿化。矿物组合为石英±黄铁矿±钾长石(图4A、B、C)。

第Ⅱ阶段:石英-黄铁矿阶段。常形成黄铁绢英岩化蚀变岩、黄铁矿化碎裂岩、石英黄铁矿脉等。石英以烟灰色为主,他形粒状结构,粒径多为0.01~0.3mm,沿裂隙充填于第Ⅰ阶段石英脉中,形成复合石英脉体或在构造裂隙形成黄铁绢英岩化蚀变岩、黄铁矿化碎裂岩。黄铁矿呈他形粒状,粒度极小,呈浸染状、细脉状产于构造蚀变岩或石英脉中,与金矿化最为密切。围岩蚀变主要有绢云母化、硅化、黄铁矿化、黄铁绢英岩化。矿物组合为石英±黄铁矿±绢云母±金。为主成矿阶段(图4C、D、E)。

第Ⅲ阶段:石英-多金属硫化物阶段。常形成多

金属硫化物石英脉,多金属硫化物碎裂岩,厚度多为2~20cm。石英为无色、透明,半自形—他形粒状结构,粒径介于0.03~0.3mm之间,常沿构造裂隙充填于第Ⅰ阶段和第Ⅱ阶段的石英脉复合体或构造蚀变岩中。黄铁矿呈半自形—他形粒状,多呈立方体,五角十二面体,呈浸染状、星点状产于石英脉中或两侧。黄铜矿呈他形粒状,呈浸染状分布于石英脉中部;方铅矿呈半自形—他形粒状分布,呈梳状分布于石英脉两侧;闪锌矿半自形粒状,粒径较大0.1~2cm,常呈脉状分布在石英脉一侧。蚀变主要为硅化、绢云母化等。矿物组合为石英±方铅矿±闪锌矿±黄铜矿±黄铁矿±绢云母±金。为重要成矿阶段(图4F、G)。

第Ⅳ阶段:石英-碳酸盐阶段。常发育石英方解石脉(团块)、萤石石英脉和碳酸盐化碎裂岩。石英方解石以胶结物形式充填于构造角砾岩之间或呈细

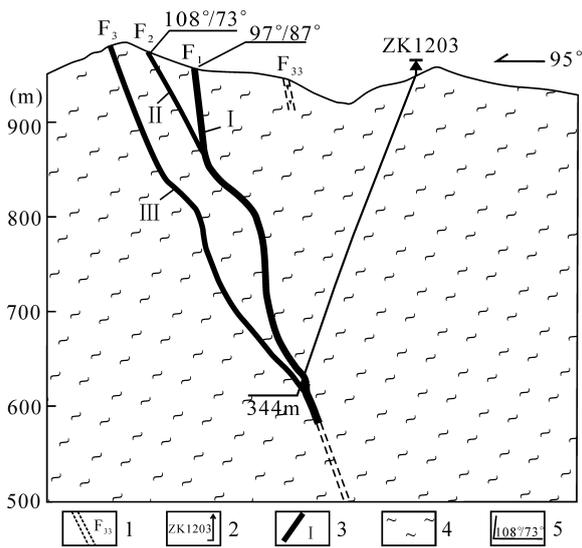


图 3 吉家洼金矿第 12 勘探线剖面图
(据 Yan Zhengxin, 2012 修改)

Fig. 3 Cross-section of typical exploration line of the Jijiawa Au deposit (modified after Yan Zhengxin, 2012)
1—含矿断裂带及其编号; 2—钻孔位置及编号; 3—矿体及编号;
4—新太古界太华群片麻岩; 5—断层产状(倾向/倾角)
1—Fault zone with ore and their numbers; 2—positions of borehole and their numbers; 3—ore body and their numbers; 4—Taihua Group in the Archeozoic; 5—fault attitude (tendency/dip angle)

脉穿切交代早阶段的石英脉; 萤石石英脉多呈细脉穿切早阶段石英脉。石英为乳白色, 他形粒状结构, 粒径介于 0.05~0.4mm 之间。方解石为白色, 半自形—他形粒状结构, 晶粒粗大, 晶形较好, 常见方解石晶簇。萤石为紫色, 呈半自形—他形粒状分布在石英中, 粒径一般为 0.1~1mm。围岩蚀变为硅化、碳酸盐化、萤石化。矿物组合为方解石±石英±萤石(图 4H、I)。为热液活动的尾声。

3 样品特征与实验方法

本次实验的 9 件闪锌矿样品采自吉家洼金矿 550m 中段第 7 勘探线至第 11 勘探线之间的不同位置, 样品均为石英—多金属硫化物阶段(Ⅲ)的多金属硫化物石英脉, 石英脉宽约 2~8cm, 石英为无色、透明, 半自形—他形粒状结构。所采样品的主要金属矿物为闪锌矿、方铅矿、黄铜矿、黄铁矿。闪锌矿呈棕褐色, 反射色为灰色, 其他形粒状结构和固溶体分离结构; 方铅矿呈铅灰色, 反射色为灰白色, 呈半自形粒状和他形粒状结构, 常交代黄铁矿和磁黄铁矿; 黄铜矿呈铜黄色, 反光镜下亦呈铜黄色, 除在闪锌矿中呈乳滴状结构外, 亦可见独立的黄铜矿(图 5)。

为了挑选出纯度较高的闪锌矿, 选样前首先磨制光片, 尽量挑选出固溶体分离结构不发育的闪锌矿样品挑选单矿物。将样品粉碎到 40~80 目, 在双目镜下挑选出闪锌矿, 样品的纯度达 99% 以上。将选出的闪锌矿用蒸馏水清洗, 低温蒸干, 然后将纯净的单矿物样品在玛瑙研钵内研磨至 200 目左右待测。因为闪锌矿等金属矿物的 Rb、Sr 含量较低, 甚至低于 0.01×10^{-6} , 为了确保 Rb、Sr 同位素定年的可行性, 我们首先在南京现代分析中心同位素分析室对上面的 9 个闪锌矿样品进行了微量元素 Rb、Sr 含量的草测, 在此基础上, 挑选适合定年的样品在南京现代分析中心同位素分析室进行 Rb、Sr 含量和同位素组成测定。具体分析方法如下: 原粉末样品用混合酸溶解, 取清液上离子交换柱分离, 采用高压密闭熔样和阳离子交换技术分离和提纯, 然后用英国产的 VG354 质谱仪测定, 测定方法见文献(Wang Yinxi et al., 2007a, 2007b)。用于测定的美国 NBS987 同位素标样为: $^{87}\text{Sr}/^{86}\text{Sr} = 0.710236 \pm 7$, Sr 的全流程空白为 $5 \times 10^{-9} \sim 7 \times 10^{-9} \text{g}$, $^{87}\text{Sr}/^{86}\text{Sr}$ 同位素比值用 $^{86}\text{Sr}/^{88}\text{Sr} = 0.1194$ 进行标准化。 $^{87}\text{Sr}/^{86}\text{Sr}$ 的分析误差为 $\pm 1\%$, $\lambda_{\text{Rb}} = 1.42 \times 10^{-11} \text{a}^{-1}$ 。等时线年龄用 ISOPLOT (Ludwing, 1998) 程序计算。

我们选取了 4 件黄铁矿样品, 其中含钾长石石英脉(I)中的黄铁矿 1 件, 黄铁绢英岩(II)中的黄铁矿 1 件, 烟灰色石英脉(II)中的黄铁矿 2 件, 测试黄铁矿的微量铅含量。将样品粉碎到 40~80 目, 在双目镜下挑选出闪锌矿, 样品的纯度达 99% 以上。铅同位素分析由核工业北京地质研究院分析测试研究中心完成。样品先用三酸分解, 然后用树脂交换法分离出铅, 蒸干后进行同位素测定, 所用仪器为 MAT-261 型质谱仪, 分析精度对 $1\mu\text{g}$ 铅含量其 $^{204}\text{Pb}/^{206}\text{Pb}$ 低于 0.05%, $^{208}\text{Pb}/^{206}\text{Pb}$ 一般不大于 0.005%。对国际标样 NBS981 的测试结果为: $^{208}\text{Pb}/^{206}\text{Pb} = 2.162189$, 误差 0.0027%; $^{207}\text{Pb}/^{206}\text{Pb} = 0.913626$, 误差 0.0059%; $^{204}\text{Pb}/^{206}\text{Pb} = 0.059201$, 误差 0.0015%。

4 测试结果

闪锌矿的 Rb、Sr 含量和同位素组成测试结果见表 1。得到的 $^{87}\text{Rb}/^{86}\text{Sr}$ — $^{87}\text{Rb}/^{86}\text{Sr}$ 图表现出很好的线性关系。计算获得闪锌矿等时线年龄 $t = 118.2 \pm 2.4 \text{Ma}$, 初始铷同位素组成 $I_{\text{Sr}} = 0.712523$, $\text{MSWD} = 0.25$ (图 6)。



图4 吉家洼金矿4个成矿阶段矿化特征

Fig. 4 Characteristics of four ore-forming mineralization of the Jijiawa Au deposit

(a)—乳白色石英脉(I),围岩发生钾长石化和黄铁矿化;(b)—黄铁矿化含钾长石石英脉(I);(c)—石英黄铁矿脉(II)被碳酸盐(IV)胶结,并见少量含钾长石石英脉(I);(d)—黄铁矿化角砾岩(II),围岩发生硅化和黄铁绢英岩化;(e)—黄铁矿化角砾岩(II),角砾为乳白色石英,胶结物为硅质、绢云母、黄铁矿;(f)—多金属硫化物石英脉(III)被石英方解石脉(IV)穿插,围岩发生硅化和绢云母化;(g)—多金属硫化物石英脉(III),见黄铜矿、方铅矿、闪锌矿分布于石英脉中部和两侧;(h)—石英碳酸盐团块(IV);(i)—萤石石英脉(IV);Kfs—钾长石;Qz—石英;Py—黄铁矿;Gn—方铅矿;Sp—闪锌矿;Ccp—黄铜矿;Carb—碳酸盐;Cal—方解石;Fl—萤石

(a)—Milky white quartz vein of stage I, K-feldspathization and pyritization in the wall rock; (b)—pyritization K-feldspar quartz vein of stage I; (c)—quartz-pyrite vein of stage II cemented by carbonate of stage IV, and meet with little K-feldspar quartz vein of stage I; (d)—pyritization breccia of stage II, silicification and beresitization in the wall rock; (e)—pyritization breccia of stage II; (f)—Ccp-Gn-Sp quartz vein of stage III cut by quartz-calcite vein of stage IV, silicification and sericitization in the wall rock; (g)—Ccp-Gn-Sp quartz vein of stage III; (h)—quartz carb of stage IV; (i)—fluorite-quartz vein of stage IV; Kfs—K-feldspar; Qz—quartz; Py—pyrite; Gn—galena; Sp—sphalerite; Ccp—chalcopryrite; Carb—carbonate; Cal—calcite; Fl—fluorite

吉家洼金矿床获得的4件铅同位素样品的分析结果见表2。Pb同位素分析表明,吉家洼金矿黄铁矿的铅同位素组成变化范围较大: $^{206}\text{Pb}/^{204}\text{Pb} = 17.042 \sim 18.647$,平均值为17.844; $^{207}\text{Pb}/^{204}\text{Pb} = 15.460 \sim 15.653$,平均值为15.545; $^{208}\text{Pb}/^{204}\text{Pb} = 37.746 \sim 39.665$,平均值为38.630。所有的铅同位素投影点都落在地幔铅和造山带铅演化线之间(图7),具有壳幔混合源的特点。

5 讨论

5.1 成矿时代

矿床成矿时代的精确测定一直是矿床学研究的一个难点。近年来,利用矿石矿物(如闪锌矿等)或与成矿有关的脉石矿物(如萤石等)的Rb-Sr同位素体系来直接获得成矿年龄已涌现大量成功实例(Nakai et al., 1990, 1993; Brannon et al., 1992;

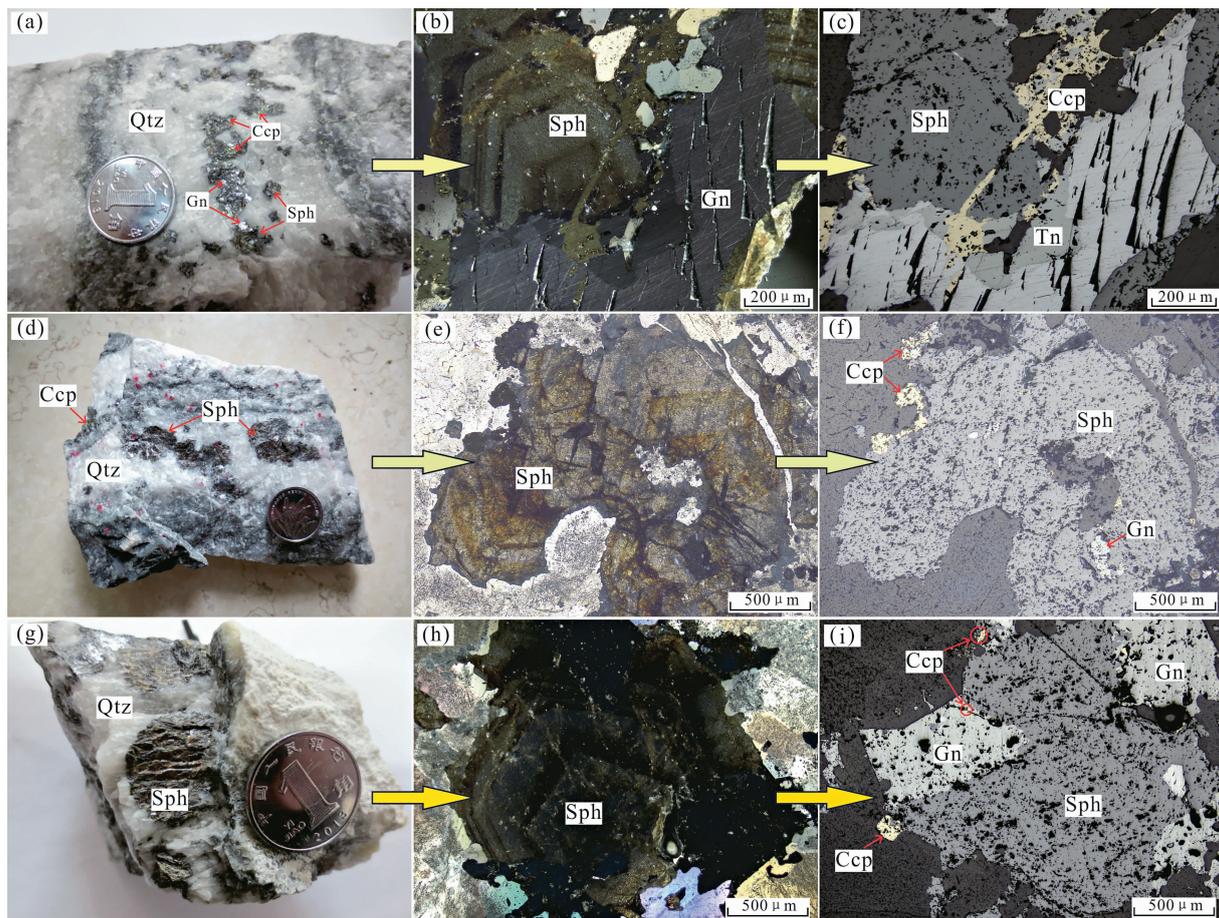


图 5 吉家洼金矿床石英-多金属硫化物阶段闪锌矿特征

Fig. 5 Characters of the ore mineral samples of quartz-polymetallic sulfide phase in the Jijiawa Au deposit

(a) 多金属硫化物石英脉,方铅矿、闪锌矿和黄铜矿共生产出;(b) 共生的闪锌矿和方铅矿,闪锌矿具环带(单偏光);(c) 闪锌矿、方铅矿、黄铜矿共生(反射光);(d) 黄铜矿闪锌矿化石英脉,闪锌矿和黄铜矿共生产出;(e) 闪锌矿呈棕黄色,具环带(单偏光);(f) 闪锌矿、方铅矿、黄铜矿共生,闪锌矿与黄铜矿形成固溶体分离结构(反射光);(g) 闪锌矿化石英脉,闪锌矿呈团块状;(h) 闪锌矿呈棕褐色,具环带(单偏光);(i) 闪锌矿、方铅矿、黄铜矿共生,闪锌矿、方铅矿与黄铜矿形成固溶体分离结构(反射光);Sph—闪锌矿;Gn—方铅矿;Ccp—黄铜矿;Qtz—石英

(a)—Polymetallic sulfide quartz vein, galena and sphalerite associated with chalcopyrite; (b)—sphalerite associated with galena, sphalerite with zone (under transmitted light); (c)—galena and sphalerite associated with chalcopyrite (under reflective light); (d)—quartz vein with chalcopyrite and sphalerite, sphalerite associated with chalcopyrite; (e)—yellowish brown sphalerite, with zone (under transmitted light); (f)—galena and sphalerite associated with chalcopyrite, chalcopyrite and sphalerite assuming interstitial separation structure (under reflective light); (g)—quartz vein with sphalerite, sphalerite crumbly; (h)—dark sepia sphalerite, with zone (under transmitted light); (i)—galena and sphalerite associated with chalcopyrite, chalcopyrite, galena and sphalerite assuming interstitial separation structure (under reflective light); Sph—sphalerite; Gn—galena; Ccp—chalcopyrite; Qtz—quartz

Tretbar et al., 2000; Yang Jinhui et al., 2000a, 2000b, 2001; Wang Yanbin et al., 2004; Hu Qianqing et al., 2012; Zheng Wei et al., 2013; Li Tiegang et al., 2014)。热液矿物 Rb-Sr 等时线年龄测年的基本前提是同源、同时、封闭性、一致的 ($^{87}\text{Sr}/^{86}\text{Sr}$), 以及具有不同的 ($^{87}\text{Rb}/^{86}\text{Sr}$); (Li Wenbo et al., 2002)。实验过程中, 将闪锌矿粉碎至 200 目以下后, 然后进行超声波清洗, 基本可排除次生及原生包裹体的干扰 (Liu Jianming et al., 1998b)。热液矿床形成的时间一般为数百万年, 而一组热液共生矿物的生成时限往往则只有数十万年, 因此对

于不同矿物的 Rb-Sr 等时线定年而言可视为基本同时形成 (Liu Jianming et al., 1998b)。本次工作选择未见或少见裂隙, 且结晶较好的多金属硫化物石英脉型矿石为研究对象, 且闪锌矿纯度高, 采自同一矿体局部较小的范围内, 最大程度满足了 Rb-Sr 同位素测年的前提条件。硫化物 Rb-Sr 同位素定年是直接将硫化物全溶, 测其 Rb-Sr 同位素组成, 以确定其形成年龄, 即成矿年龄。

Li Wenbo et al. (2002) 提出了可利用 $1/\text{Sr}-^{87}\text{Sr}/^{86}\text{Sr}$ 图和 $1/\text{Rb}-^{87}\text{Rb}/^{86}\text{Sr}$ 图来判别闪锌矿生长期间 $^{87}\text{Sr}/^{86}\text{Sr}$ 初始值是否保持不变, 进而判别

表1 吉家洼金矿床闪锌矿 Rb-Sr 同位素组成

Table 1 Rb-Sr isotopic analyses of sphalerite from the Jijiawa Au deposit

样品号	样品名称	Rb($\times 10^{-6}$)	Sr($\times 10^{-6}$)	$^{87}\text{Rb}/^{86}\text{Sr}$	$^{87}\text{Sr}/^{86}\text{Sr}$	($^{87}\text{Sr}/^{86}\text{Sr}$) _i	1/Rb($\times 10^{-6}$)	1/Sr($\times 10^{-6}$)
JJW-C1	闪锌矿	0.1124	2.689	0.1237	0.712745 \pm 9	0.71254	8.896797	0.371885
JJW-C2	闪锌矿	0.1538	2.105	0.2142	0.712841 \pm 8	0.71248	6.501951	0.475059
JJW-C3	闪锌矿	0.2705	2.582	0.3085	0.713072 \pm 10	0.71255	3.696858	0.387297
JJW-C4	闪锌矿	0.4314	2.243	0.5679	0.713423 \pm 9	0.71247	2.318034	0.445831
JJW-C6	闪锌矿	0.9536	1.104	2.541	0.716882 \pm 8	0.71262	1.048658	0.905797
JJW-C8	闪锌矿	0.1057	3.647	0.0852	0.712685 \pm 10	0.71254	9.460738	0.274198
JJW-C9	闪锌矿	1.802	0.8134	6.534	0.723483 \pm 9	0.71252	0.554939	1.229407
JJW-C11	闪锌矿	0.2456	0.9788	0.7627	0.713823 \pm 9	0.71254	4.071661	1.021659
JJW-C12	闪锌矿	0.9283	1.621	1.693	0.715312 \pm 8	0.71247	1.077238	0.616903

测试单位:南京大学现代分析中心同位素分析室。

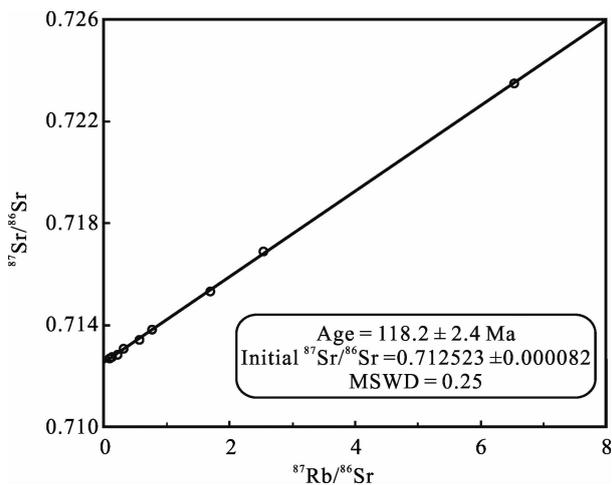


图6 吉家洼金矿床单矿物闪锌矿 Rb-Sr 等时线图

Fig. 6 Rb-Sr isochron of sphalerite from the Jijiawa Au deposit

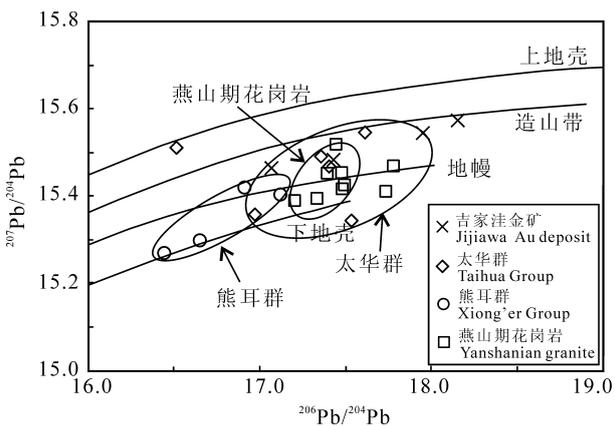


图7 吉家洼金矿床铅同位素构造模式图

(底图据 Zartman et al., 1988)

Fig. 7 Plumb tectonic model for the Jijiawa Au deposit

(base diagram from Zartman et al., 1988)

所测数据的合理性。本次测试结果显示在 $1/\text{Rb}-^{87}\text{Rb}/^{86}\text{Sr}$ 和 $1/\text{Sr}-^{87}\text{Sr}/^{86}\text{Sr}$ (图 8) 关系图解中显示, 不同闪锌矿单矿物样品的 Rb、Sr 含量不同, 而 $^{87}\text{Sr}/^{86}\text{Sr}$

表2 熊耳山地区地层、岩体及吉家洼金矿 矿石的铅同位素组成

Table 2 Lead isotopic composition of ores from the Jijiawa Au deposit and rocks in the Xiong'er shan area

位置	测试对象	$^{206}\text{Pb}/^{204}\text{Pb}$	$^{207}\text{Pb}/^{204}\text{Pb}$	$^{208}\text{Pb}/^{204}\text{Pb}$	来源
吉家洼金矿	含钾长石石英脉中的黄铁矿(I)	17.424	15.485	38.142	本文
	黄铁绢英岩中的黄铁矿(II)	17.960	15.550	38.868	
	烟灰色石英脉中的黄铁矿(II)	17.042	15.460	37.746	
	烟灰色石英脉中的黄铁矿(II)	18.149	15.575	38.728	
	平均(n=4)	17.844	15.545	38.630	
太华群	黑云斜长片麻岩	17.400	15.469	38.174	范宏瑞等, 1994
	黑云斜长片麻岩	15.406	15.188	37.526	
	角闪斜长片麻岩	16.511	15.512	36.266	
	黑云斜长片麻岩	17.353	15.492	42.558	
	金云母角闪岩	16.968	15.359	37.775	
	斜长角闪岩	17.609	15.547	37.654	
	角闪斜长片麻岩	17.530	15.345	38.569	
	平均(n=7)	16.968	15.416	38.360	
熊耳群	大斑安山岩	16.907	15.421	36.346	范宏瑞等, 1994
	安山岩	16.647	15.300	36.876	
	杏仁状安山岩	16.439	15.271	36.489	
	蚀变安山岩	17.116	15.405	37.345	
	平均(n=4)	16.777	15.349	36.764	
花山花岗岩体	角闪黑云二长花岗岩	17.476	15.418	37.832	范宏瑞等, 1994
	黑云二长花岗岩	17.329	15.396	37.488	
	钾长石	17.440	15.520	37.975	
	钾长石	17.199	15.391	37.447	
	钾长石	17.473	15.455	37.886	
	二长花岗岩钾长石	17.771	15.471	37.863	陈旺等, 1995
	二长花岗岩钾长石	17.745	15.412	37.983	王志光等, 1997
	二长花岗岩全岩	17.495	15.426	37.850	
	二长花岗岩钾长石	17.371	15.455	37.769	
平均(n=9)	17.458	15.429	37.754		

和 $^{87}\text{Rb}/^{86}\text{Sr}$ 值相对稳定, 说明闪锌矿生长期间 $^{87}\text{Sr}/^{86}\text{Sr}$ 初始值基本保持不变, 因此该等时线具有实际地质意义。

从图 5 中可看出, 等时线上所有的样点几乎全部落在等时线上, 说明矿石矿物形成过程中 Sr 同位

素是均一的,而且得到了很好的封闭,因此拟合的等时线年龄具有很高的精度。本次获得的闪锌矿 Rb-Sr 等时线年龄为 118.2 ± 2.4 Ma,与熊耳山地区中庙岭金矿钾长石 $^{40}\text{Ar}-^{39}\text{Ar}$ 等时线年龄 (117.0 Ma) (Zhai Lei et al., 2012)、祁雨沟金矿钾长石 $^{40}\text{Ar}-^{39}\text{Ar}$ 等时线年龄 (125.1~115.1 Ma) (Wang Yitian et al., 2001) 等金矿年龄,在误差范围内基本一致,也客观地说明了本次所测成矿年龄的准确性和可行性。因此,吉家洼金矿的成矿时间为 118.2 Ma,即早白垩世。

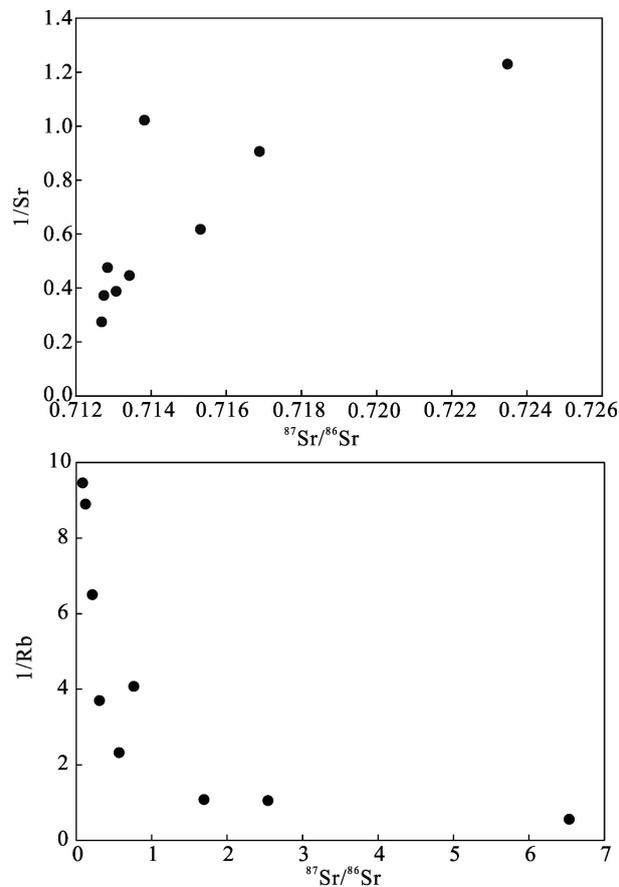


图 8 吉家洼金矿床闪锌矿 $1/\text{Sr}-^{87}\text{Sr}/^{86}\text{Sr}$ 和 $1/\text{Rb}-^{87}\text{Rb}/^{86}\text{Sr}$ 关系图

Fig. 8 Diagrams of $1/\text{Sr}-^{87}\text{Sr}/^{86}\text{Sr}$ and $1/\text{Rb}-^{87}\text{Rb}/^{86}\text{Sr}$ of sphalerites from the Jijiawa Au deposit

5.2 成矿物质来源

吉家洼金矿床闪锌矿锶同位素的测试结果(表 1)显示,闪锌矿的 Rb 含量较低,变化范围为 $0.1057 \times 10^{-6} \sim 1.802 \times 10^{-6}$,Sr 含量变化相对较大,从 $0.8134 \times 10^{-6} \sim 3.647 \times 10^{-6}$ 。同位素比值 $^{87}\text{Rb}/^{86}\text{Sr}$ 变化范围较大,从 0.0852~6.534,平均值为 1.426, $^{87}\text{Sr}/^{86}\text{Sr}$ 比值变化相对较小,从 0.7127~0.7235,平均值为 0.715。

($^{87}\text{Sr}/^{86}\text{Sr}$)₀ 是判断成岩成矿物质来源的重要指标,在矿床地质研究中常利用其来示踪成矿物质来源、岩浆流体、深源流体的壳幔混染作用(Hou Minglan et al., 2006)。为避免放射性 ^{87}Rb 衰变对锶同位素造成显著影响,将成矿时代换算到 118 Ma,采用软件 Geokit(Lu Yuanfa, 2004) 分别对硫化物进行了初始同位素计算(表 2)。从表 2 可以看出吉家洼金矿床闪锌矿等时线年龄给出的初始同位素比值非常接近,且比值相对较高,介于 0.71247~0.71262 之间,平均值为 0.7125,小于大陆地壳锶同位素 $^{87}\text{Sr}/^{86}\text{Sr}$ 平均值 0.719(Sun Shengli, 2001),而高于地幔 Sr 的初始值 0.707(Faure, 1986),显示吉家洼金矿的成矿物质来源为壳幔混合,以壳源为主。

对 Pb 同位素的研究显示,吉家洼金矿矿石中的 Pb 同位素投影点落在地幔铅和造山带铅演化线之间(图 7),具有壳幔混源的特征,这与 Sr 同位素指示的特征相一致。结合前人对太华群、熊耳群和花山花岗岩的铅同位素研究结果,不难发现,太华群变质岩的铅同位素组成变化很大,从下地壳演化线变化到造山带演化线,显然存在放射性成因铅;熊耳群火山岩的铅同位素主要分布在上地幔演化线两侧,具有幔源铅的特点;花山花岗岩体的投影点均分布在太华群变化范围之内,另有研究显示,矿区附近的花山花岗岩体的成岩物质来源于受幔源改造的太华群变质岩,(Fan Hongrui et al., 1994; Xiao E et al., 2012; Hu Xinlu et al., 2013),可以认为花山花岗岩体的铅来自新太古界太华群变质岩。对比吉家洼金矿矿石与区内其他地质体的铅同位素组成可发现,矿石铅同位素比值远高于熊耳群的铅同位素,且大部分落在太华群和燕山期花岗岩的变化范围之内(图 7),暗示新太古界太华群基底及其重熔形成的花岗岩体为吉家洼金矿铅同位素的主要物源。

综上所述,吉家洼金矿的成矿物质来自壳幔混源,新太古界太华群基底及其重熔形成的花岗岩体为吉家洼金矿铅同位素的主要来源。

5.3 地质意义

在地壳演化过程中,成矿过程总是受到一个特定区域内的重要构造-热事件或其他异常地质事件的影响和制约(Zhou Taofa et al., 2000; Mao Jingwen et al., 2004; Wang Yanbin et al., 2004)。熊耳山地区处于华北地台南缘和秦岭造山带的衔接过渡带上,区内金矿床已知成矿年龄有:庙岭金矿为 117.0 ± 1.6 Ma(Zhai Lei et al., 2012),祁雨沟金矿

为 $125.1 \pm 1.59 \sim 115.1 \pm 1.5 \text{Ma}$ (Wang Yitian et al., 2001) 和 $126 \pm 11 \text{Ma}$ (Han Yigui et al., 2007), 公峪金矿为 $115.2 \pm 2.58 \text{Ma}$ (Li Li et al., 2002), 前河金矿为 $127.0 \pm 1.6 \text{Ma}$ (Tang Kefei et al., 2013), 上宫金矿早、中、晚成矿年龄分别为 $242 \pm 10 \text{Ma}$ 、 $165 \pm 7 \text{Ma}$ 和 $113 \pm 6 \text{Ma}$ (Chen Yanjing et al., 2004)。本文获得的吉家洼金矿年龄为 118.2Ma , 与之基本一致, 属于区域上早白垩世同一地质成矿事件的产物, 对熊耳山地区金矿成矿时代具有一定的约束意义。

在矿区外围东侧出露了五丈山、花山花岗岩体, 以及零星发育一些小的花岗斑岩和岩脉。五丈山岩体 SHRIMP 锆石 U-Pb 年龄为 $156.8 \pm 1.0 \text{Ma}$ (Mao Jingwen et al., 2005), 花山岩体 SHRIMP 锆石 U-Pb 年龄为 $127.6 \pm 1.3 \text{Ma}$ (Meng Fang et al., 2012), 蒿坪岩体 SHRIMP 锆石 U-Pb 年龄为 $130.7 \pm 1.4 \text{Ma}$ (Li Yongfeng et al., 2005)、 $128.7 \pm 1.0 \text{Ma}$ 和 $129.3 \pm 2.4 \text{Ma}$ (Xiao E et al., 2012), 金山庙岩体 SHRIMP 锆石 U-Pb 年龄为 $129.3 \pm 1.5 \text{Ma}$ (Meng Fang et al., 2012) 和 $127.6 \pm 1.6 \text{Ma}$ (Xiao E et al., 2012), 雷门沟石英斑岩 U-Pb 年龄为 $127.2 \pm 1.4 \text{Ma}$ (Chen Xiaodan et al., 2011)。吉家洼金矿的成矿年龄略晚于岩体的形成。Zhang Xingkang et al. (2014) 对矿区的流体包裹体和成矿流体的 H、O 同位素组成进行了研究, 结果表明初始成矿流体来源于燕山期花岗岩浆热液, 随着成矿作用的持续进行, 有不同比例的大气降水的混合, 成矿温度和矿化强度逐渐减弱, 即成矿流体由岩浆热液向岩浆、大气降水混合热液方向演化。说明吉家洼金矿与燕山期花岗岩体有成因上的密切联系。

在华北板块与扬子板块于印支期碰撞并闭合之后, 华北陆块南缘熊耳山地区卷入到秦岭造山带的地质演化过程中, 成为了秦岭造山带的重要组成部分, 并记录了其自中生代的构造演化过程。在印支期, 秦岭地区全面的陆陆碰撞闭合。晚三叠世中期, 秦岭造山带发生构造体制转换(区域挤压构造体制向伸展构造体制转换), 推覆造山(Zhang Guowei et al., 1996, 2001), 小秦岭—熊耳山地区进入中生代早期的陆内区域伸展构造体制的演化期(Mao Jingwen et al., 2005); 在晚侏罗世(大致约 165Ma), 碰撞造山作用结束, 受太平洋构造域的影响, 整个中国大陆中东部的区域构造体制发生转换(Ren Jishun et al., 1991; Mao Jinwen et al., 2008a, 2008b), 包括小秦岭—熊耳山地区在内的华北克拉

通南缘转入中生代晚期的伸展构造环境(Mao Jingwen et al., 2005); 早白垩世, 沿造山带走向形成拆离断层体系, 导致软流圈物质上涌和地幔楔部分熔融引起上部地壳的伸展隆升, 岩石圈大规模快速减薄(Ren Jishun et al., 1997)。伴随着壳-幔相互作用、构造及岩浆活动、地热异常、流体活动的不断进行, 导致了华北地台南缘的燕山中晚期强烈的岩浆活动和大规模成矿作用, 形成了小秦岭、熊耳山等金钼多金属矿集区。由此可见, 吉家洼金矿的形成与区域上早白垩世岩石圈减薄环境下构造-岩浆-流体活动有密切的成因关系。

同样位于华北板块南缘的小秦岭变质核杂岩和崤山变质核杂岩中, 金矿床普遍发育。 $^{40}\text{Ar}-^{39}\text{Ar}$ 年代学研究表明, 小秦岭地区主要金矿化发生在 $130 \sim 125 \text{Ma}$ (Xu Qidong et al., 1998; Wang Yitian et al., 2001; Li Qiangzhi et al., 2002; Li Jianwei et al., 2012), 崤山地区金矿的成矿年龄为 130Ma (Zhu Jiawei et al., 1999)。此外, 我国胶东地区石英脉型金矿床的主成矿期成矿时代为 $123 \sim 119 \text{Ma}$ (Yang Jinhui et al., 2000a, 2001), 破碎带蚀变岩型金矿的成矿时代主要集中在 $100 \sim 117 \text{Ma}$ (Yang Jinhui et al., 2000b)。辽宁地区的五龙金矿主成矿期成矿时代为 $120 \sim 112 \text{Ma}$ (Wei Junhao et al., 2001)。本次所获得的闪锌矿 Rb-Sr 等时线年龄与以上数据基本一致。因此, 吉家洼金矿所在的熊耳山金矿集区可能是整个中国东部金矿大规模成矿作用中的组成部分。

6 结论

(1) 熊耳山地区吉家洼金矿床石英-多金属硫化物阶段的闪锌矿单矿物 Rb-Sr 等时线年龄为 118.2Ma , 指示其成矿时代为早白垩世。

(2) 铷-铅同位素研究结果显示, 吉家洼金矿的成矿物质来自壳幔混源, 新太古界太华群基底及其重熔形成的花岗岩体可能为铅同位素的主要物源。

(3) 该年龄与熊耳山地区中的庙岭、祁雨沟、公峪等金矿床的成矿年龄基本一致, 对于整个熊耳山地区的金矿成矿年龄具有一定的约束意义。吉家洼金矿床是熊耳山地区在早白垩世区域性强烈的构造-岩浆-流体活动事件的产物, 是整个中国东部金矿大规模成矿作用中的组成部分。

注 释

① 河南省地质矿产开发局第一地质矿产调查院, 1997. 河南省吉家

洼金矿床详查报告.

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Metallogenic Epoch and Source of Ore-Forming Material of the Jijiawa Au Deposit, Western Henan Province: Evidence from Rb-Sr Dating and Pb Isotopes

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Abstract

The Jijiawa gold deposit, located in the western part of the Xiong'er shan gold polymetallic ore district, Henan Province, is typical of tectonically altered rock-quartz vein-type gold deposit. Orebodies are hosted mainly within nearly NS-trending fault zone and are distributed in the form of "Y" shape. Mineralization is dominated by vein-veinlet and disseminated pyritized altered rock gold deposits, with minor being quartz vein-type gold deposits. The hydrothermal ore-forming process can be divided into four stages: pyrite-quartz vein stage (I), quartz-pyrite vein stage (II), quartz-polymetallic sulfides stage (III), and quartz-carbonate vein stage (IV). This study carried out Rb-Sr isotopic dating for the sphalerites from main mineralization stage of the Jijiawa Au deposit to restrain the ore-forming time. The yielded isochron age of 118.2 ± 2.4 Ma suggests that the deposit was formed in Early Cretaceous. The age is basically in agreement with that of Miaoling, Qiyugou, Gongyu and other gold deposits in the Xiong'er shan area, which is of significance in constraining the ore-forming age in the Xiong'er shan area. Sr and Pb isotope analysis shows that the Jijiawa Au deposit had a mixing origin of the materials from crust and mantle, with the Early Cretaceous granite likely contributing main metallogenic materials. In the context of regional geology, we suggest that the Jijiawa Au deposit was the product of the regional intensive tectono-magmatic-fluid activity in the Xiong'er shan area in Early Cretaceous, and it is also a component part of the large-scale gold mineralization in eastern China.

Key words: sphalerite Rb-Sr isotopic dating; Pb isotope; metallogenic epoch; source of ore-forming material; Jijiawa Au deposit; Xiong'er shan