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东营凹陷平南地区火山岩中发现含石盐子晶流体包裹体及其对岩浆脱气的启示

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内容提要:本文报道在东营凹陷平南地区滨古斜15井火山岩岩心中发现含石盐子晶流体包裹体。流体包裹体赋存于石英闪长玢岩的石英矿物之中, 经显微观察和-185°C低温原位拉曼光谱分析, 确定子晶为石盐矿物。显微测温分析结果显示, 原生包裹体的均一温度在359~496°C, 流体的盐度为43.26%~54.51%。研究样品中5种类型流体包裹体共生, 说明流体发生了不混溶作用。岩浆上升过程中温度和压力的降低导致流体发生强烈的不混溶作用, 并促进岩浆脱气。由于富CO₂流体以岩浆脱气的方式逃离岩浆, 而高盐度流体作为脱气之后的残余流体保存在包裹体中, 所以包裹体中以高盐度流体(NaCl-H₂O体系)为主, 而另一端元组分(富CO₂流体)较少。含石盐子晶流体包裹体的发现证实平南地区岩浆中存在高盐度流体, 研究结果表明该区的岩浆脱气过程与高盐度流体不混溶作用有关。这个发现对进一步了解岩浆脱气机制和中国东部无机成因CO₂成藏有重要的启示意义。

关键词:石盐子晶; 流体包裹体; 岩浆脱气; 不混溶作用; 东营凹陷; 平南地区

地球内部脱气作用形成的无机成因气藏在全球均有分布, 中国东部已发现大量与岩浆活动有关的无机成因CO₂气藏(戴金星等, 1992, 1995, 1997; 戴春森等, 1992, 1994, 1995; 陈永见等, 1999; 涂光炽, 1996; 王先彬等, 1997; 郑乐平等, 1995, 1998, 上官志冠等, 1996, 1997; Staudacher, 1997; 陶明信等, 1996), 关于CO₂的成因和气体地球化学特征已做了大量的研究(戴金星等, 1992, 1995, 1997; 樊棋诚等, 1996; 陶士振等, 1998; 张铭杰等, 1998, 1999; 许多等, 1999), 但一直未见有关CO₂气藏的火山岩含石盐子晶流体包裹体报道。平南地区位于东营凹陷的高青-平南断裂带, 沿该断裂带的平方王、平南、高青及花沟等地区分布着许多CO₂气藏(图1), 气体地球化学与同位素研究结果证实该区CO₂的来源主要与火山岩有关(戴金星等, 1995, 1997; 侯贵廷等, 1996; 孙明良等, 1996; 赫英等, 1996, 1998; 曾溅辉等, 2004)。流体包裹体因为能够真实地纪录地下流体活动的历史而备受地质学家们青睐(Roedder, 1979, 1984, 1992)。为更加深入地认识平南地区CO₂

气藏与油藏的成藏机制, 笔者对该区的流体包裹体样品进行了详细地研究, 并在平南油田滨古斜15井的火山岩岩心中发现了含石盐子晶流体包裹体。本文对平南地区火山岩中含石盐子晶流体包裹体进行了仔细分析和讨论。

1 研究区地质概况

平南地区位于济阳坳陷东营凹陷西斜坡上的北东-南西方向上延生的断块型古生界潜山带。其北为平方王潜山披覆构造带, 西靠里则镇洼陷, 东临小营油田, 南与樊家油田相接, 面积约50km², 属于平方王-平南潜山披覆构造带的一部分。该区处于北东-北西断裂带的交汇处, 构造运动频繁, 构造应力集中, 新老断层都很发育, 断层以北东走向为主(曾溅辉等, 2004; 郑德顺等, 2005)。平南地区东南紧靠高青-平南大断层(图1), 一方面由于断层长期剧烈活动, 在断层上升盘形成与断层走向基本一致且裂缝发育古生界断块潜山构造, 另一方面因为燕山期和喜山期的岩浆入侵为该区CO₂气藏带

注: 本文为国家973项目(编号G199043302)、中国石油大学研究生院创新基金项目(编号B2004-05)及胜利油田科技攻关项目子课题“平南油田油气成藏模式研究”资助成果(编号M040118)资助的成果。

收稿日期: 2006-03-21; 改回日期: 2006-06-28; 责任编辑: 郝梓国, 刘志强。

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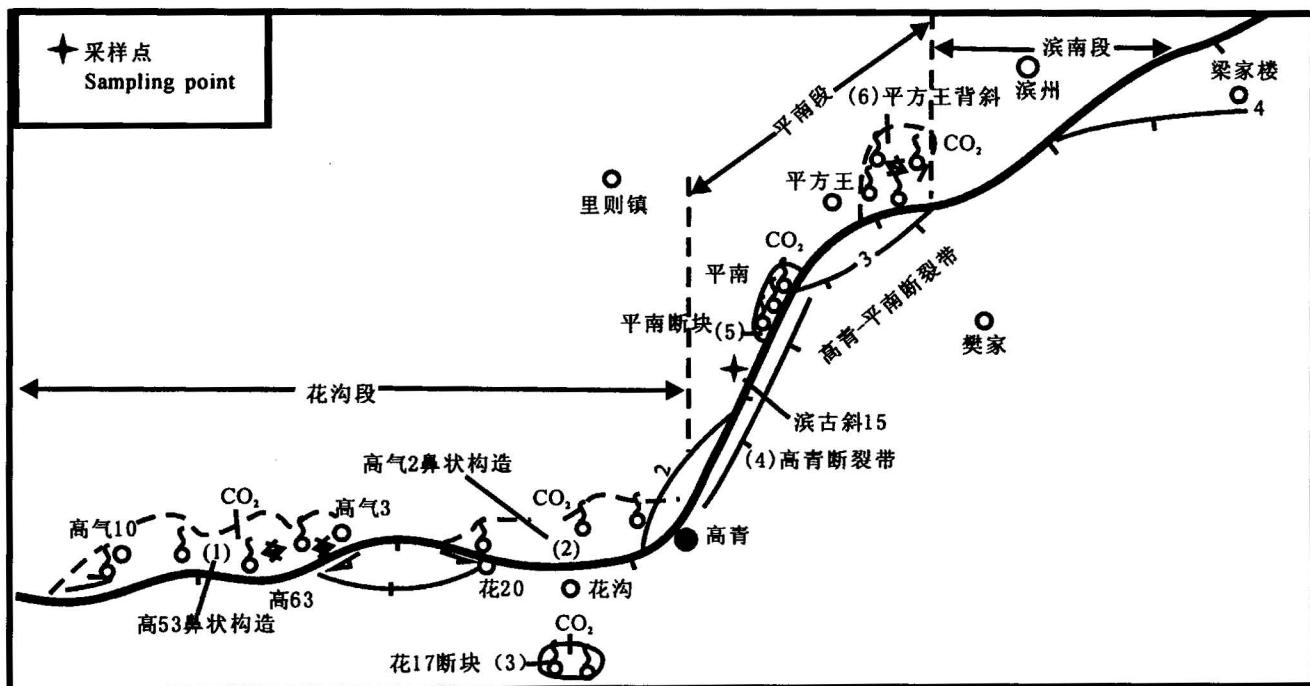


图 1 平南地区地质略图及采样点位置

Fig. 1 Geological outline of Pingnan area and sampling location

来了气源。勘探和研究结果显示该区存在大量无机成因 CO₂ 气藏。

2 样品特征

本次研究样品取自滨古斜 15 井的钻井岩心, 取样深度为 2418m, 岩性为浅灰色石英闪长玢岩, 为穿插在奥陶系地层中的侵入岩体, 据前人研究表明岩浆入侵时代为中生代早白垩世早期^①。岩屑录井资料显示该层闪长玢岩分布深度在 2310~2470m, 但是没有钻穿。样品的矿物成分有中酸性斜长石角闪石、黑云母、石英, 基质成分为闪长岩, 岩石为斑状结构, 部分碳酸盐化。将其中三块样品磨制成两面抛光的岩石薄片, 厚度约 0.1~0.2mm, 便于观察和测试流体包裹体。

3 实验条件

本文所有实验分析在中国石油大学地球化学与岩石圈动力学开放实验室完成。包裹体显微测温所用仪器为英国 Linkam 公司生产的 THMS-600 冷热台, 经过标准样品校正, 100℃以下误差为 0.1℃(包括冷冻过程), 100℃以上误差为 0.5℃。拉曼光谱分析仪器为法国 J. Y. 公司生产的 LabRam-010 型激光拉曼光谱仪, 该仪器采用显微共焦技术, CCD 信

号检测系统, 仪器光学分辨率约 1.5cm⁻¹。实验所用激光波长为 514.5nm, 束斑直径约 1μm, 共焦孔 1000μm, 狹缝 400μm, 信号积分时间为 60s, 积分次数为 3 次, 实验环境温度为 19℃。拉曼光谱分析前用单晶硅标样对光谱仪进行校正。

4 实验结果

4.1 包裹体形态及类型

在显微镜下观察发现, 在石英矿物中存在大量流体包裹体, 既有原生流体包裹体, 也有次生流体包裹体。原生流体包裹体无规律分布, 一般单独出现, 包裹体呈多种形态出现, 每个包裹体中的相态也不尽相同, 部分包裹体含有石盐子晶, 石盐的立方晶形清晰可见(图 2)。而后期次生流体包裹体都是沿裂隙成线状分布, 一般为气液两相或者单液相出现。本文主要对原生流体包裹体进行研究, 按相态可以把样品中的流体包裹体分为五种类型:

(I) 液-固(石盐子晶)两相包裹体(图 2a), 固液比变化较大, 显微镜下目测在 5%~50% 之间;

(II) 气-液-固(石盐子晶)三相包裹体(图 2b), 气泡占整个包裹体体积的比值在 3%~10% 之间, 而子晶所占体积比值较稳定, 在 10% 左右, 气相拉曼光谱分析未见 CO₂ 气体信号。

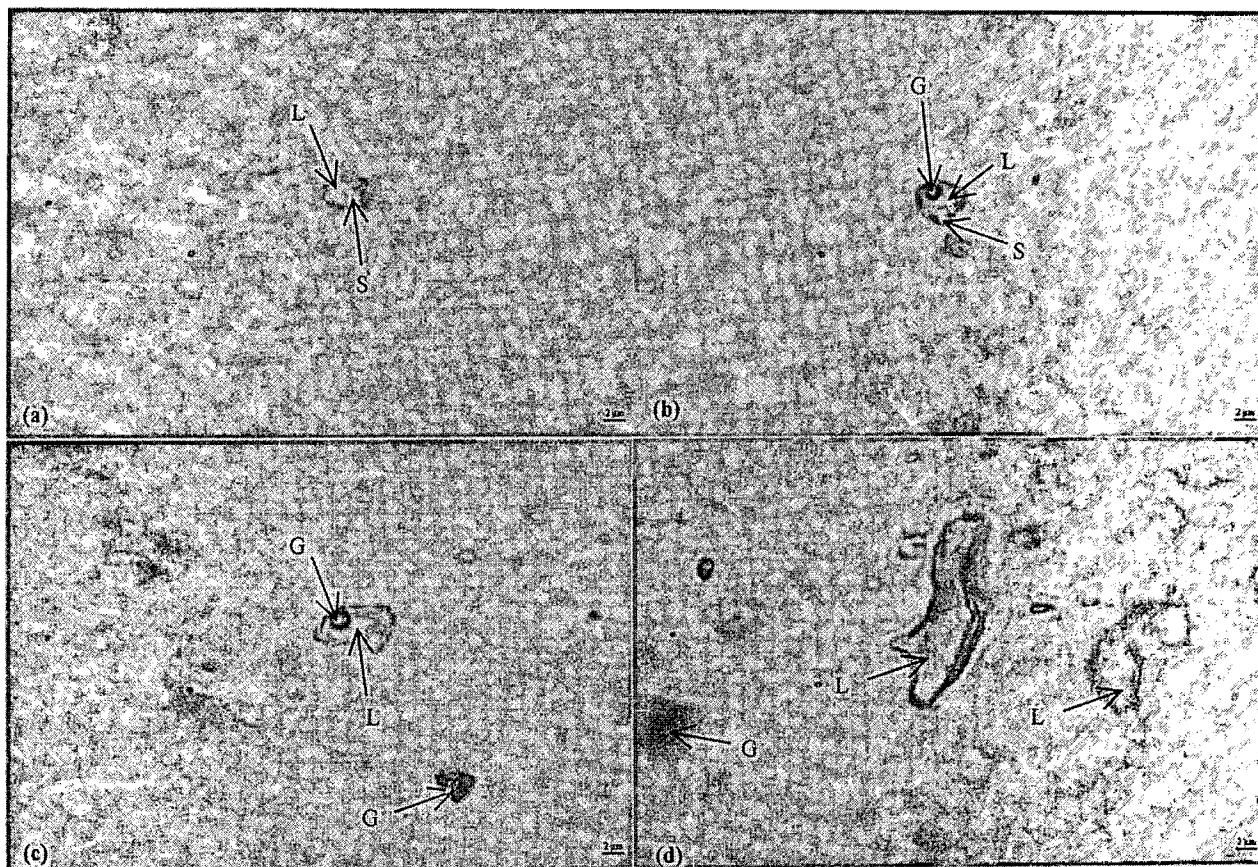


图 2 不同类型原生流体包裹体照片

Fig. 2 photos of primary fluid inclusion in different types

G—气相,L—液相,S—固相(石盐子晶)

G—gas phase, L—aqueous phase, S—solid phase.

(Ⅲ) 气-液两相包裹体, 气液比在 3%~10%, 包括两种流体, 即早期岩浆活动石英矿物结晶条件下捕获的流体和后期构造裂缝愈合捕获的流体(图 2c), 其均一温度具有明显的差别。室温下对原生流体包裹体进行拉曼光谱分析显示, 包裹体气相成分存在微弱的 CO₂ 信号, 液相未检测出 CO₂ 的拉曼信号, 原生包裹体为 H₂O-CO₂(极少)-NaCl 体系; 而次生包裹体为低盐度盐水体系。

(Ⅳ) 单液相包裹体(图 2d), 拉曼光谱分析表明为 NaCl 盐水体系。

(Ⅴ) 单气相包裹体(图 2c,d), 拉曼光谱分析显示含 CO₂ 气体。

其中含石盐子晶包裹体体积较小, 一般 2~5 μm 左右, 其它三类包裹体体积变化较大, 多数在 1~15 μm 之间。I 型、Ⅱ型、Ⅲ型和Ⅳ型包裹体在样品中最常见、数量最多, 而单气相包裹体数量较少、

不常见。

4.2 拉曼光谱分析

将样品薄片从载玻片上取下, 用丙酮洗净胶质后, 再用去离子水将样品冲洗干净, 然后将样品放在冷热台中。通过温度控制和原位拉曼光谱监测, 获得了在 19°C 和 -185°C 低温条件下包裹体中盐水相的拉曼光谱(图 3)。室温下拉曼分析结果表明, 所有原生包裹体的液相在室温(19°C)下均具有很高的盐度, 盐度约为 25W% NaCl(陈勇等, 2002), 水的伸缩振动拉曼峰表现为尖锐而不对称, 其峰强度的最大值偏移到 3441 cm⁻¹。在 -185°C 下获得了包裹体中冰和 NaCl 水合物的拉曼特征峰(图 3)。冰的拉曼特征峰位于 3091 cm⁻¹ 和 3216 cm⁻¹, 而位于 3401 cm⁻¹、3418 cm⁻¹、3434 cm⁻¹ 及 3531 cm⁻¹ 的拉曼峰是 NaCl 水合物的典型拉曼特征峰(Dubessy et al., 1982; Ni et al., 2006), 这充分证实了包裹体中的子晶为石盐

晶体。

4.3 显微测温分析结果及讨论

将样品置于冷热台上对流体包裹体进行均一过程分析,包裹体的均一温度分布直方图如图4所示,由图明显可以看出两期流体的存在。均一过程显微测温分析结果表明,所有石盐子晶在359~460℃之间完全熔化,查相关数据表(卢焕章等,2004)得到含石盐子晶包裹体的盐度分布范围在43.26%~54.51%,这个结果证实了原生包裹体中流体具有非常高的盐度。原生包裹体的均一温度在359~496℃,都是均一液相,500℃以上部分包裹体出现爆裂现象。次生流体包裹体的均一温度分布在100~130℃。

水的临界温度为374.3℃,由图4可知大部分原生包裹体是在水的超临界状态下捕获的。超临界流体与一般的气体、液体在物化性质上有很大的不同,它对盐类的溶解度会大大增高,特别是对一些气水热液难溶或不溶物的矿质(如某些金属卤化物、络合物)却有很强的溶解性,但当这种流体温压降至临界点时,溶解度会突然变小,被溶矿质大量析出沉淀(徐有生等,1995;施立达等,2001)。由此可知,超临

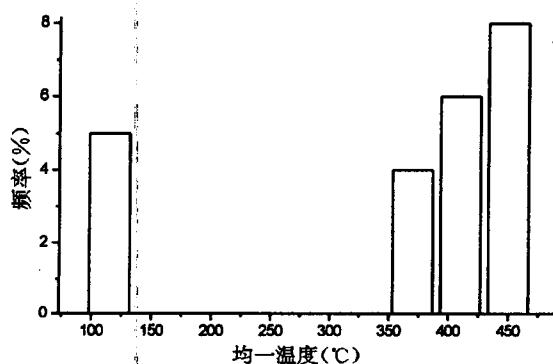


图4 均一温度分布直方图

Fig. 4 Histograms of homogenization temperatures of fluid inclusions

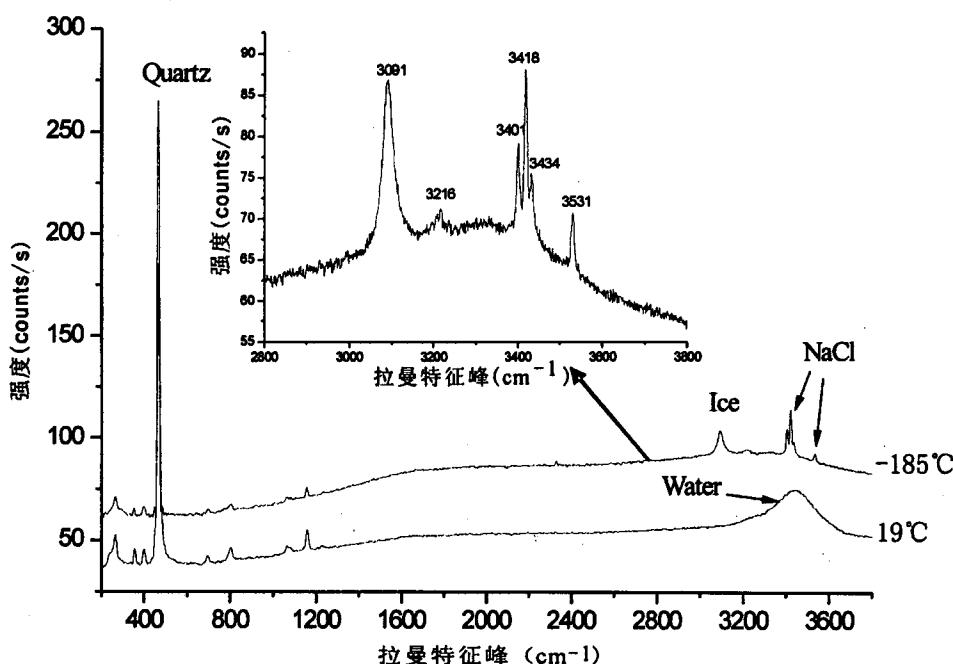


图3 19℃和-185℃下包裹体中液相的拉曼光谱

Fig. 3 Raman spectra of aqueous liquid phase in fluid inclusion at 19°C and -185°C

可知,超临界条件有利于NaCl溶解而易形成高盐度流体,在温度降低后可以结晶出石盐晶体,这解释了原生包裹体中出现石盐晶体的原因。

5 流体不混溶作用与岩浆脱气

值得注意的是,上述五类流体包裹体在同一颗石英矿物中出现(图1中包裹体均在同一石英颗粒中),表现为共生关系,这说明包裹体中流体是在不混溶条件下捕获的(卢焕章等,2004),暗示在岩浆上升过程中发生了明显的不混溶作用(Roedder, 1992)。造成流体不混溶的原因可能有两个方面,一方面是由于岩浆上升过程中压力降低导致CO₂的溶解度降低迅速脱气造成不混溶;另一方面是由于岩浆上升过程中温度降低引起的流体发生不混溶。

关于岩浆中CO₂的溶解行为已有大量的研究成果(Vivian et al., 1991; Dixon et al., 1995a, 1995b, 1997; Papale et al., 1997, 1999; Duan Zhenhao et al., 1995)。其中Duan Zhenhao等(1995)对CO₂-H₂O-NaCl体系的物理学化学性质及其相变进行了非常详细地研究,其研究结果指出,一般情况下CO₂-H₂O体系不会发生混溶作用,但是如果该体系有NaCl出现就会发生强烈的不混溶作用,CO₂-H₂O-NaCl体系发生不混溶作用的最小温

度区间在 400~500℃,而且在压力不变情况下,温度升高或降低都会扩大不混溶作用的范围。所以,不混溶现象可以存在于水的临界温度(374.3℃)之上,而且在高温条件下的相分离作用可以把中等盐度的水溶液变为高盐度流体。尽管 NaCl 的存在大大降低了水的活度,却增加了 CO₂ 的活度,结果导致 CO₂ 在溶液中的溶解度急剧下降(Duan Zhenhao et al., 1995)。而 Webster 等(1999)用实验证明了 Cl⁻在玄武岩脱气中发挥重大的作用。因此,高盐度可以促使 CO₂ 气体从岩浆中脱离。CO₂ 和 H₂O 是岩浆中最主要的挥发分,由于 NaCl 的出现,流体发生不混溶作用会产生两种端元流体。对 CO₂-H₂O-NaCl 体系就会形成富 CO₂ 贫水贫 NaCl 流体和富水贫 CO₂ 高盐度流体,即 CO₂-H₂O(贫水)体系和 NaCl-H₂O 体系。本次研究样品中以纯盐水包裹体(纯液相)和含石盐子晶流体包裹体(液固两相)为主,属于富水贫 CO₂ 高盐度流体(NaCl-H₂O 体系),几乎不含 CO₂,而另一个端元组分富 CO₂ 体系在样品中表现为纯气相包裹体,但数量很少,这说明包裹体捕获的流体为不混溶作用(沸腾)导致岩浆发生强烈脱气后的残余流体。

6 结论与认识

采用显微测温与低温原位拉曼光谱技术分析和鉴定了东营凹陷平南地区火山岩中的含石盐子晶流体包裹体。通过对含石盐子晶流体包裹体研究,揭示了该区岩浆中存在高盐度流体。结合超临界流体质,对高盐度 CO₂-H₂O-NaCl 体系的不混溶作用进行了分析讨论,高盐度会促使流体发生不混溶作用并减小 CO₂ 在流体中的溶解度,流体的不混溶作用可以加速 CO₂ 的脱气过程。火山岩中含石盐子晶流体包裹体的发现,证实了平南地区的岩浆 CO₂ 脱气过程与高盐度流体有密切的关系,这个发现对研究平南地区乃至整个中国东部岩浆脱气机制及 CO₂ 成藏有重要的意义。由于缺少其它类岩浆中的高盐度流体数据,关于含 CO₂ 高盐度流体不混溶行为与岩浆脱气的作用机制还需进一步深入研究。

本文在成文过程中得到中国石油大学(华东)吕洪波教授的指导,在此特表谢意。

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The Discovery of Bearing Halite Daughter Minerals Fluid Inclusions in Volcanic Rocks at Pingnan Area in Dongying Depression and Implications to Magma Degassing

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Abstract

This paper has reported the discovery of bearing halite daughter mineral fluid inclusions in volcanic rocks at Pingnan area in Dongying depression. The samples were collected from BGX-15 well drill cores and the interested fluid inclusions are hosted in quartz of diorite-porphyrite. The daughter minerals are NaCl crystals which were identified by observation under microscope and analyzed by in situ Raman spectroscopy at -185°C. The results of micro-thermal analysis show that the homogenization temperatures of primary fluid inclusions are between 359°C and 496°C, and the salinities of fluid inclusions are from 43.26‰ to 54.51‰. All fluid inclusions in studied samples can be divided into five types including primary fluid inclusions and secondary fluid inclusions. Five type fluid inclusions were symbiotic in a same quartz grain, which imply that immiscibility happened in magma. Due to the decreases of temperature and pressure during magma ascending, the fluids became intensively immiscible. This process will accelerate degassing of CO₂ from magma, and the result of the CO₂-rich fluids had got away from magma by degassing, but the remnant fluids with high salinity were saved in fluid inclusions. So, the primary fluid inclusions are mainly in NaCl-H₂O fluids, poor in CO₂. The results of our study indicate that the degassing of magma and accumulation of CO₂ gas at Pingnan area are relative to the immiscibility of high salinity fluids. This discovery is important to deep understand the mechanism of magma degassing and accumulation of inorganic CO₂ in eastern China.

Key words: halite daughter minerals; fluid inclusion; magma degassing; immiscibility; Dongying depression; Pingnan area