

Research Advances

A Marine or Continental Nature of the Deltas in the Early Cretaceous Lingshanda Formation—Evidences from Trace Elements

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Objective

A succession with well-developed soft-sediment deformation structures and well-exposed gravity-flow deposits in the Early Cretaceous Lingshanda Formation has caused considerable controversy on whether the non-disputed deltaic environment was marine or continental. This dispute cannot be solved by mere sedimentological analysis and, moreover, too few fossils are present to be decisive. The main objective of the present contribution is to shed light on this problem on the basis of trace-element analysis.

Methods

Field work was carried out for collecting samples of dark shales from both the middle and the upper Lingshanda Fm., in the western Yellow Sea. Six of the samples were used for Analysis of 16 trace elements. The trace-elements analysis was performed by means of inductively coupled plasma mass spectrometry (ICP-MS). The sedimentary environment during the Early Cretaceous was determined on the basis of the trace-element analysis in combination with the ratios of some specific elements and plant fragments. The paleosalinity (Sp) of the Lingshanda Fm. was calculated from the boron (B) concentration following Adamas Formula ($Sp=0.0977*B-7.043$).

Results and Analysis

The ratios of the 16 trace elements from the six samples are provided in Table 1. These data can be used for analysis of main sedimentary environments, such as marine vs. continental facies, but also for determining the paleosalinity and oxidation-reduction potential.

The boron can be used to distinguish deposits with a marine facies (usually 80–125 ppm) from those with a continental facies (usually less than 60 ppm). The boron (B) concentration in the Lingshanda Fm. varies from 82.734 to 174.672 ppm, thus being always higher than 80 ppm. The paleosalinity (Sp), calculated by using Adamas Formula, ranges from 1.040‰ to 10.022‰, which indicates a brackish-water environment. The high concentration of strontium (Sr) in the six samples is high, varying from 229.197 to 408.056 ppm, thus pointing at a marine environment (>160 ppm) rather than a continental one (usually < 90 ppm). The concentration of vanadium (V), ranging from 78.165 to 115.089 ppm is intermediate between the values known for marine (<86 ppm) and continental sediments (>110 ppm). The gallium (Ga) content in the Lingshanda Fm., however, seems to point consistently to a continental environment (15–23 ppm), rather than to a marine environment (<15 ppm).

Although the concentrations of some trace elements may be indicative of a marine or continental environment, the ratio of some particular elements is more reliably conclusive. The B/Ga ratio varies from 4.347 to 7.732, which points to a marine facies (>4.2), rather than a continental facies (<3.3). The Sr/Ba ratio ranges from 0.296 to 0.709, mostly being higher than 0.35, indicating a marine environment (>0.35), rather than a continental one (<0.2). The Rb/Sr ratio, ranging from 0.299 to 0.663 also infers a marine environment, as well as a humid paleoclimate.

The Cu/Zn, V/Cr, Ni/Co, U/Th, and V/(V+Ni) ratios are indicative of the oxidation-reduction potential. The Cu/Zn ratio in the Lingshanda Fm. ranges from 0.235 to 0.390, thus indicating a slightly reducing environment. The V/(V+Ni) ratios vary between 0.724 and 0.800, equally implying a slightly reducing environment, as well as a stratified flow of seawater. However, the ratios of V/Cr (<2), Ni/Co (<5) and U/Th (<0.75) imply a slightly

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Table 1 Trace elements and ratios of some specific elements in the Lingshanda Fm.

Element/Ratio	Blank (ppb)	ZC1 (ppm)	ZC2 (ppm)	ZC3 (ppm)	DT1 (ppm)	DT2 (ppm)	DT3 (ppm)
Sc	0.078	11.041	11.717	14.514	11.057	12.781	12.516
V	0.013	78.165	95.167	115.089	82.243	96.383	93.760
Cr	0.633	66.046	54.105	65.194	58.537	59.101	68.918
Co	0.031	7.597	13.176	13.287	5.120	12.133	7.084
Ni	0.350	29.861	25.360	30.801	21.549	27.349	23.473
Cu	0.065	15.805	29.119	29.363	17.342	28.185	17.288
Zn	0.289	41.061	74.668	81.207	73.518	72.997	73.649
Ga	0.004	18.148	18.201	22.591	18.949	20.252	19.031
Rb	0.125	118.515	122.213	184.621	126.721	154.172	128.414
Sr	0.015	251.878	408.056	278.553	313.242	282.225	229.197
Ba	0.084	729.813	575.561	638.569	738.532	700.540	773.345
La	0.003	45.792	41.865	45.999	35.359	44.047	34.576
Pb	0.087	12.632	25.587	23.095	8.865	23.654	9.556
Th	0.003	12.893	13.210	16.546	12.873	14.358	14.005
U	0.001	3.286	4.074	3.155	3.189	3.450	3.356
B	/	119.946	97.313	174.672	103.977	107.395	82.734
Sp(‰)	/	4.676	2.464	10.022	3.116	3.449	1.040
B/Ga	/	6.609	5.347	7.732	5.487	5.303	4.347
Sr/Ba	/	0.345	0.709	0.436	0.424	0.403	0.296
Rb/Sr	/	0.471	0.299	0.663	0.405	0.546	0.560
Cu/Zn	/	0.385	0.390	0.362	0.236	0.386	0.235
V/Cr	/	1.183	1.759	1.765	1.405	1.631	1.360
Ni/Co	/	3.931	1.925	2.318	4.209	2.254	3.314
U/Th	/	0.255	0.308	0.191	0.248	0.240	0.240
V/(V+Ni)	/	0.724	0.790	0.789	0.792	0.779	0.800

*Analyses were performed at the State Key Laboratory of Continental Dynamics, Northwest University, Xi'an, China.

oxidizing environment.

Field observations provide additional clues regarding the sedimentary environment of the Lingshanda Fm.. Numerous remains of carbonized plant fragments found on top of sandy debrites in the formation suggest that the sediments originated from the continent. This may explain the apparently inconsistent conclusions drawn on the basis of the trace elements: the salinity of the sea water was periodically lowered by the input of fresh-water floods from the nearby continent.

Conclusion

The Early Cretaceous Lingshanda Fm. was deposited in a deep sea instead of marine or continental deltas. The

salinity of the sea water was lower than normal because of the inflow of fresh-water floods from the nearby continent. The oxidation-reduction potential in the depositional environment varied from slightly reducing to slightly oxidizing-flow events and fresh-water input. A humid climate is indicated by both the trace elements and the abundance of plant fragments.

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