

Research Advances

The Redefinition on Formation Time of the Lapeiquan Group in the Hongliugou Area, North Altyn: Constrains from New Detrital Zircon LA-ICP-MS U-Pb Ages



DONG Shunli¹, YAN Zhaokun^{2*}, REN Jing^{3,4}, WANG Xinwei¹, ZHAO Shaoze¹ and DENG Tao¹

¹ State Key Laboratory of Oil and Gas Reservoir Geology and Exploitation, Chengdu University of Technology, Chengdu 610059, China

² State Key Laboratory of Nuclear Resources and Environment, East China University of Technology, Nanchang 330013, China

³ Key Laboratory of Sedimentary Basin and Oil and Gas Resources, Ministry of Land and Resources, Chengdu 610081, China

⁴ Chengdu Center, China Geological Survey, Chengdu 610081, China

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Objective

The Early Paleozoic tectonic events such as oceanic subduction and collision-orogeny in the North Altyn have been extensively studied. The Lapeiquan Group, distributed in the Hongliugou ophiolite mélangé belt (Fig. 1), was interpreted to form along with the Ordovician Caledonian tectonic events. Based on few evidence of macrofossils, the formation age of the Lapeiquan Group was the Late Ordovician, which remains controversial. This work reports initially the detrital zircon LA-ICP-MS U-Pb ages of the sandstone sample from this group, the youngest U-Pb ages enable us to redefine the formation age of the Lapeiquan Group, further provide new basic information for studies on Early Paleozoic tectonic evolution in the North Altyn.

Methods

Based on detailed geological field survey in the Hongliugou area, we selected fresh coarse sandstone samples from the Lapeiquan Group for subsequent zircon selection and U-Pb dating. Zircons were separated using standard heavy liquid, magnetic techniques and hand picking at the Langfang Regional Geological Survey Institute in Hebei Province, China. Cathodoluminescence images and LA-ICP-MS U-Pb dating of detrital zircons were both carried out at the State Key Laboratory of Lithospheric Evolution, Institute of Geology and Geophysics, Chinese Academy of Sciences, China. U-Th-Pb isotopic ratios were calculated using the Glitter program. Data reduction was conducted using Isoplot.

Results

The Lapeiquan Group is generally divided into three

segments from the bottom up based on the dominant lithology: volcanic-rock member, clastic-rock member and carbonate-rock member. The sandstone sample (sample No. HLG-03) used for U-Pb dating in this study was collected from the clastic-rock member in Qiashikansayi canal, Hongliugou area (Fig. 1). The sample HLG-03 is greywacke with matrix proportion >25%. Quartz grains yield 35% of the whole grains, while feldspars and lithic fragments yield proportions of 30% and 35%, respectively. Generally speaking, most grains are characterized by poor sorting and roundness, indicative of the immature nature and short transportation distance.

This sandstone sample yields 93 effective detrital zircon U-Pb ages in a wide range from 393 Ma to 2080 Ma (Fig. 1c), with 90% of Paleozoic ages among the whole data. As a whole, it yields a most evident peak age of 475 Ma with several Precambrian ages distributing sporadically. It's noteworthy that this sandstone sample includes 15 youngest U-Pb ages (discordance from -7% to 10%) between 393 Ma and 436 Ma with a peak age of 430 Ma (Fig. 1d), i.e., ages younger than Ordovician, which occupies 16% of the whole ages. The zircons yielding these young ages mostly present clear oscillatory zoning (Fig. 1e) and have wholly Th/U ratios of greater than 0.1, indicative of their magmatic origin.

Conclusions

The sandstone sample collected from the clastic-rock member of the Lapeiquan Group, North Altyn, yields a wide age ranging from 393 Ma to 2080 Ma with 15 young U-Pb ages between 393 Ma and 436 Ma, and CL images and Th/U ratios of the young zircons suggest their magmatic origin. Thus, the Lapeiquan Group must have formed after 393 Ma. In addition, it is inferred that the deposition time of the Lapeiquan Group is the Devonian or younger age but not the Late Ordovician.

* Corresponding author. E-mail: yzk517@163.com

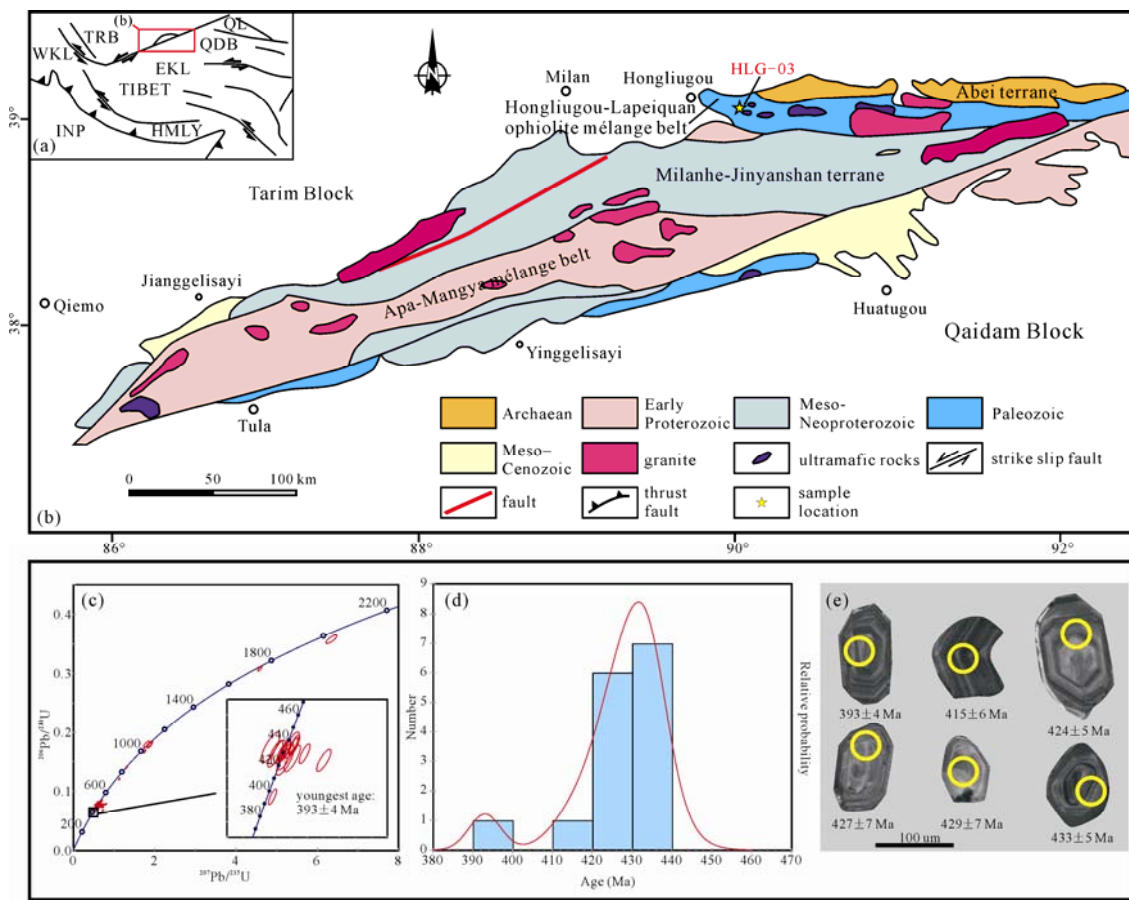


Fig. 1. (a) Generalized geological map of western China; (b) simplified geological map of the Altyn orogenic belt, showing the location of sample in this study; (c) concordia diagrams (one for whole age data, another for the ages younger than 440 Ma) of zircon U-Pb for the sandstone sample; (d-e) age spectra and cathodoluminescence images of the young zircons (younger than 440 Ma) from the sandstone sample, respectively.

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Appendix 1 Detrital zircon U-Pb dating result of the studied samples in the Lapeiquan Group

Sample No.	²³² Th/ ²³⁸ U	Isotope ratios						Apparent ages (Ma)					
		²⁰⁷ Pb/ ²⁰⁶ Pb		²⁰⁷ Pb/ ²³⁵ U		²⁰⁶ Pb/ ²³⁸ U		²⁰⁷ Pb/ ²⁰⁶ Pb		²⁰⁷ Pb/ ²³⁵ U		²⁰⁶ Pb/ ²³⁸ U	
1	0.19	0.0581	0.0012	0.6259	0.0114	0.0782	0.0009	533	20	494	7	485	6
2	0.16	0.0676	0.0010	0.7221	0.0100	0.0775	0.0008	732	43	525	7	479	5
3	0.22	0.0552	0.0010	0.5985	0.0098	0.0786	0.0009	421	18	476	6	488	5
4	0.16	0.0574	0.0013	0.6114	0.0129	0.0773	0.0010	507	25	484	8	480	6
5	0.57	0.0598	0.0011	0.6712	0.0107	0.0814	0.0009	596	17	521	7	505	5
6	0.15	0.0567	0.0016	0.5753	0.0146	0.0735	0.0011	481	31	461	9	457	6
7	0.15	0.0571	0.0009	0.5991	0.0086	0.0761	0.0008	494	15	477	5	473	5
8	0.16	0.0580	0.0011	0.6129	0.0101	0.0766	0.0009	531	18	485	6	476	5
9	0.13	0.0584	0.0010	0.6023	0.0092	0.0748	0.0008	546	16	479	6	465	5
10	0.12	0.0512	0.0014	0.5476	0.0142	0.0776	0.0011	250	34	443	9	482	7
11	0.16	0.0572	0.0018	0.5517	0.0158	0.0700	0.0011	498	36	446	10	436	7
12	0.11	0.0678	0.0009	1.3126	0.0148	0.1404	0.0014	863	10	851	6	847	8
13	0.20	0.0557	0.0012	0.5333	0.0105	0.0695	0.0009	438	23	434	7	433	5
14	0.20	0.0592	0.0011	0.5945	0.0096	0.0729	0.0008	573	18	474	6	453	5
15	0.33	0.0576	0.0018	0.6395	0.0180	0.0806	0.0013	513	35	502	11	500	8
16	0.20	0.0555	0.0009	0.5774	0.0083	0.0755	0.0008	433	16	463	5	469	5
17	0.13	0.0602	0.0009	0.6350	0.0086	0.0765	0.0008	612	14	499	5	475	5
18	0.26	0.0678	0.0009	1.1309	0.0129	0.1211	0.0012	861	11	768	6	737	7
19	0.23	0.0563	0.0020	0.5989	0.0193	0.0772	0.0013	463	41	477	12	479	8
20	0.16	0.0541	0.0016	0.5316	0.0140	0.0713	0.0010	374	34	433	9	444	6
21	0.19	0.1287	0.0020	6.3544	0.0922	0.3582	0.0049	2080	11	2026	13	1974	23

