浙江治岭头钼铅锌金多金属矿床矿质来源的硫、 铅同位素示踪及成矿时代

楚克磊1,2),陈小荣3),齐刚3),高翔3),胡博3),李胜进3)

1)中国地质科学院,北京,100037;2)南京大学,内生金属矿床成矿机制研究国家重点实验室,南京,210093;
 3)浙江省第七地质大队,浙江丽水,323000

内容提要:治岭头矿床上部(火山岩盖层)以铅锌矿床、硫矿床为主,中间(变质岩基底)以金银矿、铅锌矿为主, 深部(侵入岩体)以斑岩型钼矿床为主的"三层楼"模式。本文在前人研究的基础上,通过辉钼矿 Re-Os 同位素定年 技术对治岭头矿床进行了成矿时代的厘定。6件辉钼矿样品的模式年龄为113.7~114.6 Ma,加权平均年龄为 114.03±1.78 Ma,表明治岭头钼多金属矿床成矿作用发生于早白垩世晚期。该矿床辉钼矿样品的 Re 含量变化于 16.05×10⁻⁶~66.29×10⁻⁶,表明其成矿物质具有壳幔混源的特征。矿石硫同位素组成变化范围较窄(-2.1‰~ 2.6‰),具有相对均一的来源,可能主要来自上地幔或下地壳的深源岩浆,但也可能受到陆壳沉积物的混染。治岭 头钼铅锌金多金属矿床矿石铅同位素组成变化范围比较小,²⁰⁸ Pb/²⁰⁴ Pb、²⁰⁷ Pb/²⁰⁴ Pb、²⁰⁶ Pb/²⁰⁴ Pb 的范围分别为 38.765~39.137、15.523~15.751、18.450~18.667,具有明显的壳幔混合特征。治岭头钼多金属矿床成矿年龄的 厘定为下一步在区内开展同时期的斑岩成矿系统找矿勘探提供了重要的线索,同时为进一步深入研究中国东南沿 海大陆边缘成矿带成岩成矿作用提供了新的资料和证据。

关键词:治岭头钼铅锌金多金属矿床;辉钼矿 Re-Os 定年;成矿物质来源;东南沿海成矿带

中国东南沿海大陆边缘成矿带指浙江、福建、广 东等地区,属于西太平洋成矿带的重要组成部分,区 内发育大量与燕山期火山-侵入岩有关的斑岩型铜 钼矿、脉状铅锌矿、浅成低温热液型金矿,如著名的 紫金山铜金多金属矿以及治岭头钼铅锌金多金属矿 等。其中治岭头矿床是上部(火山岩盖层)以铅锌矿 床、硫矿床为主,中间(变质岩基底)以金银矿、铅锌 矿为主,深部(侵入岩体)以斑岩型钼矿床为主的"三 层楼"模式,是探索斑岩成矿系统的理想对象。

治岭头多金属矿位于绍兴-龙泉隆起带。绍兴-龙泉隆起带属于华南地块一部分,西以江山-绍兴断 裂,东以丽水断裂为界,分别与江南古岛弧和东南沿 海海西-印支褶皱带相连接。治岭头首先发现金矿, 后陆续探明黄铁矿、金银矿、铅锌矿以及钼矿等矿产 资源。迄今为止,众多的学者对治岭头多金属矿开 展了研究,已经积累了丰富的资料(Liang Zihao et al.,1985; Zheng Minghua et al.,1986; Liu Jianming, 1990; Zhang Yaxiong et al.,1995; Chen Haoshou et al.,1996; Hua Jiexiong et al.,2000; Mei Jianming, 2001; Pu Weimin et al.,2008; Wang Keqiang et al., 2014; Wang Yongbin,2014)。前人对成矿年龄进行了 不同方法的尝试,但矿床的成矿时代数据较多。Me Jianming. (2001)获得该金矿床中的石英⁴⁰ Ar/³⁹ Ar 年 龄为 139±18.6 Ma,为白垩纪早期成矿,并认为是伸 展体制下火山、次火山热液叠加作用使八都群中的 金矿化进一步富集而成,而早期的沉积变质作用仅 使金达到矿化富集。Chen Haoshou et al. (1996)也 测得治岭头金银矿床的成矿时代在 135~145Ma 左 右;而王永彬(2014)获得铅锌矿闪锌矿⁴⁰ Ar/³⁹ Ar 等 时线 年龄 为 108.3±8.3 Ma,斑岩体 侵入 年龄 114.1±1.1 Ma(Wang Keqiang et al.,2014)。

众所周知,精确测定金属矿床的成矿时代,对于

作者简介: 楚克磊, 男, 1984年生。助理研究员, 研究方向: 矿床学与矿床地球化学。Email: chukelei@163. com。

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 Chu Kelei, Chen Xiaorong, Qi Gang, Gao Xiang, Hu Bo, Li Shengjin. 2020. Sulfur and lead isotope tracing for sources of oreforming material and ore-forming age of the Zhilingtou Mo-Pb-Zn-Au polymetallic deposit in the Zhejiang Province. Acta Geologica Sinica, 94(8):2325~2340.

正确认识矿床成因和控矿因素、总结成矿规律并指 导找 矿 勘 探 工 作 都 具 有 极 为 重 要 的 意 义 (Zhai Yusheng et al., 1992; Chen Yuchuan et al., 1994; Mao Jingwen et al., 2000, 2005)。辉钼矿 Re-Os 同 位素测年技术被公认为是较为成熟的金属矿物测年 手段,已广泛应用于直接厘定不同类型钼矿床和含 钼多金属矿床成矿年龄的有效方法(Mao Jingwen et al., 2004, 2006, 2008; Xie Guiqing et al., 2007; Zheng Wei et al., 2016; Li Houmin et al., 2007; Yuan Shunda et al. ,2012;Guo Chunli et al. ,2014; Zheng Wei et al., 2013a, 2017, 2018)。本文利用辉 钼矿 Re-Os 同位素定年法,对治岭头钼铅锌金多金 属矿的成矿时限进行厘定,并对矿床中的斑岩体、围 岩以及硫化物开展系统的 S、Pb 同位素对比分析, 以期对成矿物质来源和形成机理进行初步探讨,为 完善区内成矿年代学格架提供新的信息,同时为下 一步在区域上开展同时期的多金属矿床找矿勘探提 供重要的线索。

1 区域地质背景

中国东南沿海成矿带大地构造位置位于欧亚板 块东南缘,是环太平洋巨型构造-岩浆活动带的重要 组成部分。其处于政和-大浦断裂以东及邻区,自北 东向南西跨过浙江、福建和广东沿海地区。历经华 夏古陆的形成与裂解、扬子与华夏板块的碰撞拼合、 太平洋板块与欧亚大陆板块的相互作用阶段(Shui Tao, 1987;任纪舜等, 1990; Zhao Fengqing et al., 1996)。东南沿海矿集区历经自太古宙古陆核形成 以来的多期次岩浆活动,其中尤以中生代岩浆活动最 为宏伟壮观(Wu Ganguo et al., 2004)。岩石类型齐 全,从超基性-超酸性及碱性岩类均有,主要以酸性岩 类为主。其中东南沿海地区三叠纪和侏罗纪岩浆作 用相对于白垩纪较弱,前者主要由石英正长岩和石英 二长岩组成(Mao Jingwen et al., 2013; Zhu Kongyang et al., 2013),后者主要包括混合岩、花岗质岩石和火 山岩(Yui et al., 2009; Liu Qian et al., 2012; Li Zhengxiang et al., 2012; Huang Huiqing et al., 2013; Qiu Zengwang et al., 2017a)。白垩纪岩浆活动最为 强烈并广泛分布,早期岩浆岩主要由火山岩和花岗岩 组成(Wong et al., 2011; Jiang Yaohui et al., 2011; Guo Feng et al., 2012; Chen Jingyuan et al., 2014; Yang Shuiyuan et al., 2013; Zhou Jie et al., 2013; He Zhenyu et al., 2012; Zhu Kongyang et al., 2014),晚期 岩浆岩主要由花岗岩和流纹岩组成,含少量玄武岩和 辉长岩(Lan Chingying et al., 1996; Zhao Junhong et al., 2004; Guo Feng et al., 2012; Chen Jingyuan et al., 2013; Li Zhen et al., 2014;徐夕生, 2008)。

东南沿海矿产丰富而富有特色(图1),矿床的 时空分布与构造应力场的转换、多期次岩浆、沉积作 用等多因素密切相关,尤以中生代以来的成矿作用 强度大、持续时间长,占据主导地位(Wu Ganguo et al.,2004)。在浙闽地区以形成铅、锌、金、银、铜、钼 及非金属矿萤石、叶蜡石矿床为主,在粤东及粤西地 区主要形成钨、锡、铅、锌、铜、金、银为主的矿床(Chen Yuchuan et al., 2014; Zheng Wei et al., 2018)。粤东 和粤西地区成岩成矿受北东向与东西向深大断裂的 共同控制,而在浙闽地区则以北东向走滑断裂系对成 岩成矿起主导的控制作用(Xu Xiaochun et al., 1999; Zheng Wei,2018)。另外,浙闽粤沿海矿集区矿床的 元素组合及空间展布还受区域火山岩基底地层含矿 性的影响:如粤东地区的 Sn 矿化与金鸡组 Sn 的高含 量有关,而浙闽地区 Au 矿化可能与区内出露的变质 基底地层高背景值有关(Xu Xiaochun et al., 1999)。

2 矿区地质概况

治岭头矿区地层具有双层结构,基底为古元古 界八都群变质岩,盖层为中生代上侏罗统大爽组火 山碎屑岩和火山熔岩(图 2),两者呈角度不整合或 断层接触。八都群变质岩出露于矿区南部,主要岩 性包括黑云斜长片麻岩、斜长片麻岩、黑云二长片麻 岩、含石榴子石黑云斜长片麻岩、黑云变粒岩等。矿 区断裂构造、火山构造发育。尤其火山活动引起的 环状断裂、放射状断裂与基底构造相叠加形成的多 边环状断裂极为发育。矿区断裂构造具多期活动的 特点,主要有北东向、北西向、近南北向和近东西向 (Deng Xingen et al.,2010)。侵入岩主要为燕山期 的花岗斑岩、闪长玢岩、辉绿岩、霏细斑岩、霏细岩 等,侵位于基底八都群变质岩和大爽组火山岩中。

治岭头矿区矿产丰富,形成了贵金属金、银,有 色金属铅、锌、钼及硫铁矿等矿产组合。矿区已探明 的金银矿化带(图 3),位于华峰尖火山口北侧,呈 "S"型展布,多为近东西走向,倾向西南或东南,走 向上延伸超过 1400m,垂向上延深 200m 左右。金 银矿体,呈网脉状和角砾状赋存于含石榴子石黑云 斜长片麻岩等变质岩中。矿石矿物主要为金银矿、 银金矿、闪锌矿、黄铁矿、方铅矿、磁黄铁矿,次为自 然银、辉银矿、自然金、碲银矿、黄铜矿、硫锰矿等;脉 石矿物主要为石英、蔷薇辉石,次为玉髓、绢云母、菱



图 1 东南沿海地质简图及主要矿产分布图(据 Zhou Xinmin et al.,2006;Li Bin et al.,2016) Fig. 1 Geological map of the southeast coast and the distribution of major minerals (after Zhou Xinmin et al.,2006;Li Bin et al.,2016)

锰矿等。围岩蚀变包括硅化、黄铁矿化、绢云母化和 绿泥石化等。

铅锌矿分为隐爆角砾岩型铅锌矿和脉状铅锌 矿。隐爆角砾岩型铅锌矿为中型铅锌矿床,位于华 峰尖破火山口断陷边缘的西北侧,埋藏深度约 100m。铅锌矿体赋存在隐爆形成的火山岩中。脉 状铅锌矿矿体呈脉状分布于距火山中心约400~ 1000m的范围,主要集中在华峰尖北部和东部的杨 梅岗等地,另外在火山口西南治岭头以北也有分布。 华峰尖北部矿段呈脉状、分枝复合状,走向长度约 450m,厚度为0.25~11.20 m,平均厚2.84m,倾向 延深大于200m,倾向60°~70°。杨梅岗铅锌矿段为 一中小型矿床,矿体呈脉状,受北北西向一近南北向 的断裂构造控制,围岩为火山碎屑岩、变质岩与花岗 斑岩。金属矿物以方铅矿、闪锌矿和黄铁矿等为主。 脉石矿物主要包括石英、绿泥石、绿帘石、绢云母、长 石、碳酸盐类矿物。矿石构造以浸染状、脉状、细脉 状、斑杂状等为主,次为角砾状、团块状构造等。围 岩蚀变类型主要包括绿泥石化、绢云母化、绿帘石 化、碳酸盐化和萤石化等。

钼矿呈细脉浸染状主要赋存于花岗斑岩和变质 岩内。赋存标高在 100m 以下,其中 ZK1203 矿体 总垂厚达 78.71m,最大单层垂厚为 40.99m,Mo 品 位一般为 0.03%~0.10%,最高达 0.208%。新近 完成的钻孔 ZK1204 揭示出 6 层钼矿体,累计垂直 厚度 超过 200m,其中 IV 号钼矿体垂直厚度 196.2m,平均含 Mo 0.06%。该孔最高品位处在斑 岩体与变质岩接触带,含 Mo 达 0.63%。钼矿化主 要以石英-钾长石-辉钼矿脉、石英-辉钼矿脉、辉钼矿 细脉形式产出。围岩蚀变种类丰富,包括钾化、绢云 母化、硅化、绿泥石化和绿帘石化等典型的斑岩型矿 化蚀变类型。



图 2 治岭头多金属矿床矿区地质图(据华杰雄等,2000)





图 3 治岭头金银矿体中段展布图(据钱建民等,2007)

Fig. 3 The layout of the middle section of the gold-silver orebodies in the Zhilingtou deposit(after Qian Jianmin et al. ,2000)

3 样品采集及分析方法

3.1 Re-Os 同位素测年

为了精确厘定治岭头钼多金属矿床的成矿年龄,本文对该矿床进行了 Re-Os 同位素测年。用于

Re-Os 测年的 6 件辉钼矿样品均采自 ZK1204 钻 孔,辉钼矿主要以浸染状和脉状等形式产出。Re-Os 同位素测年在国家地质实验测试中心 Re-Os 同 位素实验室完成,分析方法和流程主要有样品的分 解、Os 的蒸馏分离、Re 的萃取分离、质谱测定 4 个 步骤,详细流程见有关文献(Shirey et al.,1995;Du Andao et al.,2001;Du Andao et al.,2004)。Re、 Os 含量的不确定度包括样品和稀释剂的称量误差、 质谱测量的分馏校正误差、稀释剂的标定误差、待分 析样品同位素比值测量误差。模式年龄的不确定度 还包括衰变常数的不确定度(1.02%),置信水平 95%。辉钼矿的模式年龄计算采用公式如下: $t=1/\lambda$ [$\ln(1+^{187} \text{Os}/^{187} \text{Re}$]。上式中¹⁸⁷ Re 衰变常数 $\lambda =$ 1.666×10⁻¹¹ a⁻¹(相对不确定度 1.02%)(Shen et al.,1996)。

3.2 S、Pb 同位素分析

本次工作主要分析了石英硫化物阶段的矿石矿 物黄铁矿、闪锌矿和方铅矿以及花岗斑岩和赋矿围 岩片麻岩样品的 S、Pb 同位素。从中选取具有代表 性的样品,经手工进行逐级破碎、过筛在双目镜下挑 选 40~60 目,纯度>99%的单矿物样品 0.5g 以上。 将挑纯后的单矿物样品在玛瑙钵里研磨至 200 目以 下,送实验室进行分析。S、Pb 同位素在核工业北京 地质研究院分析测试研究中心完成。硫同位素由 MAT-251 同位素质谱仪进行测定,采用 VCDT 国 际标准,分析精度为±0.2%;铅同位素比值利用 ISOPROBE-T 热电离质谱仪测定,²⁰⁸Pb/²⁰⁴Pb、 ²⁰⁷Pb/²⁰⁴Pb、²⁰⁶Pb/²⁰⁴Pb比值误差小于 0.05%。

4 分析结果

4.1 辉钼矿 Re-Os 年龄

选自治岭头多金属矿床的 6 个辉钼矿样品的 Re-Os 同位素测试结果如表 1 所示。由表 1 可见, 辉钼矿中¹⁸⁷ Re 和¹⁸⁷ Os 含量变化非常小,¹⁸⁷ Re 为 $10087 \times 10^{-9} \sim 41662 \times 10^{-9}$,¹⁸⁷ Os 为 19.14×10⁻⁹ ~79.29×10⁻⁹。6 个辉钼矿样品的 Re-Os 模式年龄 介于 113.7~114.6 Ma 之间,利用 Isoplot 软件 (Ludwig,2003)得到其等时线年龄为 114.3±1.1 Ma (图 4)。

表 1 浙江治岭头多金属矿床中辉钼矿 Re-Os 同位素测年结果

 Table 1
 Re-Os isotopic analyses of molybdenites from the Zhilingtou polymetallic deposit in Zhejiang Province

样号	样重(g)	Re ($\times 10^{-9}$)		普 Os (×10 ⁻⁹)		187 Re (×10 ⁻⁹)		$^{187}\mathrm{Os}$ ($ imes 10^{-9}$)		模式年龄(Ma)	
		测定值	不确定度	测定值	不确定度	测定值	不确定度	测定值	不确定度	测定值	不确定度
ZLT-1	0.01655	59708	787	0.0114	0.0009	37528	495	71.72	0.43	114.6	2
ZLT-2	0.01507	66286	852	0.0219	0.001	41662	536	79.29	0.47	114.1	2
ZLT-3	0.0151	31831	316	0.0327	0.001	20006	199	37.98	0.23	113.8	1.7
ZLT-4	0.01517	45051	367	0.0318	0.001	28316	231	53.93	0.33	114.2	1.6
ZLT-5	0.01522	16049	115	0.0215	0.0009	10087	72	19.14	0.12	113.8	1.6
ZLT-6	0.0152	49529	472	0.0518	0.0009	31130	297	59.03	0.42	113.7	1.8





4.2 硫同位素组成

治岭头矿床矿石硫化物的 S 同位素组成见表

2。32件样品 δ^{34} S 值变化于 2.3‰~7.6‰之间,极 差为 5.3‰,平均值为 5.2‰。其中,6件西矿段含 金石英脉中的黄铁矿 δ^{34} S 值为 3.3‰~4.8‰,极差 为 1.5‰,平均值为 3.95‰;7件片麻岩中的黄铁矿 δ^{34} S 值为 6.3‰~7.6‰,极差为 1.3‰,平均值为 7‰;8件方铅矿 δ^{34} S 值为 2.3‰~4.8‰,极差为 2.5‰,平均值为 3.8‰;8件闪锌矿 δ^{34} S 值为 5.8‰ ~6.8‰,极差为 1‰,平均值为 6.2‰;5件花岗斑 岩 δ^{34} S 值为 5.8‰~6.2‰,极差为 0.4‰,平均值 为 5.9‰;3件片麻岩 δ^{34} S 值为 5.1‰~6.2‰,极差 为 1.1‰,平均值为 5.63‰。王永彬(2014)曾测得 8件辉钼矿 δ^{34} S 值为 6.9‰~7.6‰,极差为 0.7‰, 平均值为 7.2‰。

4.3 铅同位素组成

本次研究对治岭头矿床中的24个金属硫化物 (6个含金矿石英脉中的黄铁矿,6个斑岩体及赋矿 围岩中的黄铁矿,7个闪锌矿以及5个方铅矿)进行 了普通铅同位素分析,并对比了3个赋矿围岩和5个

表 2 浙江治岭头多金属矿床硫同位素测试结果

		•	•	1 5	•				
序号	样号	测定矿物	岩性		δ^{34} SV-CDT($\%_0$)				
				数值	变化范围	极差	平均	, , ,	
1	V-1	黄铁矿	含金石英脉	4.2					
2	V-2-1	黄铁矿	含金石英脉	3.5					
3	V-2-2	黄铁矿	含金石英脉	4.8					
4	V-3-2	黄铁矿	含金石英脉	3.3					
5	V-3-3	黄铁矿	含金石英脉	4.2					
6	F1-(1-4)	黄铁矿	断层中的含金石英脉	3.7					
7	D07-1	黄铁矿	含铅锌矿片麻岩	6.7	3.3~7.6	4.3	5.6		
8	D08	黄铁矿	含铅锌矿片麻岩	7.6					
9	540-1-2	黄铁矿	含铅锌矿片麻岩	6.7					
10	540-3-1	黄铁矿	含铅锌矿、黄铁矿钾化片麻岩	7.1					
11	2007-23	黄铁矿	片麻岩	7.1					
12	2007-24	黄铁矿	片麻岩	6.3					
13	2007-15	黄铁矿	片麻岩	7.2					
14	D08	方铅矿	片麻岩	4.5				7	
15	D09-1	方铅矿	片麻岩	4.3					
16	D09-2	方铅矿	片麻岩	2.3	9.9-4.9	2.5	3.8		
17	540-1-1	方铅矿	片麻岩	4.8	2.3~4.8				
18	540-3-1	方铅矿	片麻岩	4.3				5.43	
19	2007-7	方铅矿	片麻岩	2.4					
20	D06-1	闪锌矿	片麻岩	6.1				1	
21	D07-1	闪锌矿	片麻岩	5.8					
22	D07-2	闪锌矿	片麻岩	6.4					
23	D08	闪锌矿	片麻岩	5.8	5 0 4 0				
24	540-1-2	闪锌矿	片麻岩	6.2	5.8~6.8	1	6. Z		
25	540-3-1	闪锌矿	片麻岩	6					
26	2007-2	闪锌矿	片麻岩	6.3					
27	2007-15	闪锌矿	片麻岩	6.8					
28	zk12-01	花岗斑岩	花岗斑岩	6.2				1	
29	zk12-06	花岗斑岩	花岗斑岩	5.7					
30	zk12-09	花岗斑岩	花岗斑岩	5.8	5.8~6.2	0.4	5.9		
31	zk12-11	花岗斑岩	花岗斑岩	6					
32	zk12-14	花岗斑岩	花岗斑岩	5.8					
33	zk12-35	片麻岩	片麻岩	5.6				1	
34	zk12-37	片麻岩	片麻岩	6.2	5.1~6.2	1.1	5.63		
35	zk12-41	片麻岩	片麻岩	5.1					

Table 2 Sulfur isotopic analytical results from the Zhilingtou polymetallic deposit in Zhejiang Province

花岗斑岩样品的 Pb 同位素数据,测试结果如表 3 所示。

含金石英脉中的 6 个黄铁矿样品²⁰⁶ Pb/²⁰⁴ Pb 比 值为 18.015~18.354,²⁰⁷ Pb/²⁰⁴ Pb 比值为 15.638~ 15.783,²⁰⁸ Pb/²⁰⁴ Pb 比值为 38.477~39.469; 斑岩 体及赋矿 围岩中的黄铁矿样品²⁰⁶ Pb/²⁰⁴ Pb 比值为 17.999~18.177,²⁰⁷ Pb/²⁰⁴ Pb 比值为 15.412~ 15.716,²⁰⁸ Pb/²⁰⁴ Pb 比值为 38.841~39.12; 7 个闪 锌矿样品²⁰⁶ Pb/²⁰⁴ Pb 比值为 17.985~18.073, ²⁰⁷ Pb/²⁰⁴ Pb 比值为 15.658~15.761,²⁰⁸ Pb/²⁰⁴ Pb 比 值为 38.914~39.219; 5 个方铅矿样品²⁰⁶ Pb/²⁰⁴ Pb 比值为 18.012~18.095,²⁰⁷ Pb/²⁰⁴ Pb 比值为 15.643 ~15.732,²⁰⁸ Pb/²⁰⁴ Pb 比值为 38.875~39.17; 5 个 花岗斑岩样品²⁰⁶ Pb/²⁰⁴ Pb 比值为 18.109~18.335, ²⁰⁷ Pb/²⁰⁴ Pb比值为 15.636~15.657,²⁰⁸ Pb/²⁰⁴ Pb 比 值为 39.092~39.544;3 个片麻岩样品²⁰⁶ Pb/²⁰⁴ Pb 比值为 18.249~18.457,²⁰⁷ Pb/²⁰⁴ Pb 比值为 15.639 ~15.662,²⁰⁸ Pb/²⁰⁴ Pb 比值为 39.259~39.705。

5 讨论

5.1 成矿时代和成矿期次

利用金属矿物进行同位素测年可以直接获取成 矿年龄,并随着分析测试技术的快速发展正逐渐成为 成矿年代学研究的趋势(Yuan Shunda et al.,2008, 2012;Zheng Wei et al.,2013b,2017)。本次测试的 6 件辉钼矿样品等时线年龄为 114.3±1.1 Ma,为治岭

表 3 浙江治岭头多金属矿床铅同位素测试结果

	_				-		-		
序号	样号	样品描述	$^{206}Pb/^{204}Pb$	$^{207}{\rm Pb}/^{204}{\rm Pb}$	$^{208}{ m Pb}/^{204}{ m Pb}$	T/Ma	μ	ω	Th/U
1	V-1	含金石英脉中的黄铁矿	18.354	15.638	38.477	255	9.54	41.1	4.17
2	V-2-1	含金石英脉中的黄铁矿	18.068	15.783	39.469	620	9.87	44.9	4.4
3	V-2-2	含金石英脉中的黄铁矿	18.028	15.738	39.382	598	9.79	44.29	4.38
4	V-3-2	含金石英脉中的黄铁矿	18.015	15.698	39.195	563	9.71	43.11	4.3
5	V-3-3	含金石英脉中的黄铁矿	18.037	15.748	39.297	603	9.81	43.96	4.34
6	F1-(1-4)	含金石英脉中的黄铁矿	18.073	15.735	39.233	564	9.78	43.29	4.28
7	D07-1	片麻岩中的黄铁矿	18.078	15.412	39.096	174	9.13	39.26	4.16
8	D08	片麻岩中的黄铁矿	18.084	15.716	39.12	536	9.74	42.51	4.22
9	540-1-2	片麻岩中的黄铁矿	18.177	15.711	39.105	466	9.71	41.79	4.17
10	540-3-1	片麻岩中的黄铁矿	18.052	15.678	38.993	515	9.66	41.76	4.18
11	2007-23	片麻岩中的黄铁矿	17.999	15.651	38.841	521	9.62	41.15	4.14
12	2007-15	片麻岩中的黄铁矿	18.014	15.663	38.88	524	9.64	41.35	4.15
13	D06-1	闪锌矿	18.057	15.687	39.005	522	9.68	41.87	4.19
14	D07-1	闪锌矿	18.073	15.708	39.074	534	9.72	42.3	4.21
15	D07-2	闪锌矿	18.064	15.694	39.03	525	9.69	42.01	4.2
16	D08	闪锌矿	17.985	15.761	39.219	652	9.84	44.11	4.34
17	540-1-2	闪锌矿	18.029	15.677	38.974	530	9.66	41.81	4.19
18	540-3-1	闪锌矿	18.017	15.658	38.914	517	9.63	41.43	4.16
19	2007-15	闪锌矿	18.05	15.717	39.063	560	9.74	42.49	4.22
20	D08	方铅矿	18.012	15.662	38.936	525	9.64	41.6	4.18
21	D09-1	方铅矿	18.044	15.675	38.953	517	9.66	41.6	4.17
22	D09-2	方铅矿	18.047	15.68	38.958	521	9.67	41.66	4.17
23	540-1-1	方铅矿	18.095	15.732	39.17	546	9.77	42.83	4.24
24	540-3-1	方铅矿	18.024	15.643	38.875	495	9.6	41.06	4.14
25	ZK12-01	花岗斑岩	18.291	15.644	39.423	307	9.56	41.72	4.22
26	ZK12-06	花岗斑岩	18.109	15.649	39.092	442	9.6	41.52	4.19
27	ZK12-09	花岗斑岩	18.335	15.643	39.544	274	9.56	41.94	4.25
28	ZK12-11	花岗斑岩	18.248	15.657	39.348	353	9.59	41.82	4.22
29	ZK12-14	花岗斑岩	18.23	15.636	39.115	341	9.56	40.72	4.12
30	ZK12-35	片麻岩	18.457	15.659	39.705	206	9.57	42.02	4.25
31	ZK12-37	片麻岩	18.426	15.639	39.564	204	9.54	41.42	4.2
32	ZK12-41		18.249	15.662	39, 259	358	9.6	41.48	4.18

Fig. 3 Pb isotope data in ore mineral from the Zhilingtou polymetallic deposit in Zhejiang Province

头钼多金属矿床提供了一个准确的形成时限。王科 强等(2014)获得治岭头花岗斑岩体的 SHRIMP 锆 石 U-Pb 同位素年龄为 114.1±1.1 Ma,其与治岭 头矿床成矿年龄在误差范围内基本一致,反映出该 矿区成岩成矿作用主要发生于早白垩世晚期。

大量高精度成岩成矿年龄数据的积累,对认识 重大成岩成矿事件非常重要。本文在梳理前人工作 成果的基础上,认为东南沿海成矿带燕山晚期矿床 可归纳为四个成矿系列:

(1)浙闽粤与侏罗纪斑岩-次火山岩有关的 Cu、 Mo、Au、Sn 矿床成矿系列,矿床类型主要以斑岩-砂 卡岩型为主。例如,粤东地区新识别出 161 Ma 的 新寮岽斑岩 Cu-Mo 矿(Wang Xiaoyu et al.,2016), 157 Ma 的钟丘洋斑岩型 Cu 矿、156 Ma 年的鸿沟山 和鸡笼山斑岩型 Cu-Au 矿(Mao Jingwen et al., 2017)以及鹅地铜多金属矿、玉水铜矿等;粤西新识 别出的旗鼓岭斑岩-矽卡岩型 Cu-Mo 矿形成于 165 Ma(Zheng Wei et al., 2018)以及茶地、芒饿岭、文 光岭、地豆岗和陂头面等铜多金属矿(Zheng Wei et al.,2015a、b); 闽西古田斑岩 Cu-Mo 矿形成于 158 ~161 Ma之间(Li Bin et al., 2016),峰岩和丁家山 等钼铅锌多金属矿斑岩成矿系统(与金东公司内部 交流,156~160 Ma之间)等;浙江地区的岭后矽卡 岩型 Cu-Mo 矿的辉钼矿 Re-Os 年龄为 162.2±1.4 Ma(Tang Yanwen et al., 2017a),桐村斑岩 Mo-Cu 矿绢云母⁴⁰ Ar-³⁹ Ar 年龄为 155.5±0.9 Ma(Tang Yanwen et al., 2017b); 在江西地区产出有 170 Ma 左右的德兴斑岩 Cu-Au-Mo 矿床(Guo Shuo et al., 2012)、156 Ma的永平矽卡岩型 Cu-Mo-W 矿(L Xiaofeng et al., 2012)和154 Ma 的龙头岗矽卡岩型 Cu-Mo-W 矿(Wu Shenghua et al., 2015),表明越 来越多的侏罗纪斑岩铜多金属矿床正逐渐被识别和 发现。斑岩型矿床是一种极为重要的指示俯冲作用 的矿床类型。越来越多的侏罗纪斑岩铜多金属矿床 逐渐被发现,表明沿着中国东部东南沿海可能存在 一个中晚侏罗世的大陆岩浆弧和相关的斑岩铜矿 带。不过这个斑岩铜矿带的大部分被白垩纪的安山 质岩石所覆盖,使得对东部陆缘不同区段的岩浆与 成矿特征、时空变化、成因以及相互间对比还缺乏很 好的限定,因此在白垩纪火山岩覆盖区的深部存在 侏罗纪斑岩铜矿。这也与 180 Ma 左右 Izanagi 板 块或古太平洋板块开始向欧亚大陆发生斜向俯冲, 中国东部大陆边缘成为活动大陆边缘的认识相吻合 (Maruyama et al.,1997;Mao Jingwen et al.,2014, 2017;Zheng Wei et al.,2018)。

(2)浙闽粤早白垩纪与侵入岩有关的 Cu、Mo、 Au、W、Sn、Nb、Ta 矿床成矿系列,主要分布在福建 南部和广东东部,矿床类型比较复杂,其中,与中酸 性次火山岩浅成斑岩侵入体有关的斑岩-浅成低温 热液铜金矿床称为紫金山式,与斑岩-次火山岩有关 的钨矿称为莲花山式、锡铅锌银矿床称为厚婆坳式 (Wang Denghong et al., 2005), 与花岗岩有关的锡 铜铅锌多金属矿床成为金坑式。福建紫金山矿田内 的岩浆活动分为晚侏罗世花岗岩和白垩纪火山侵入 杂岩两个体系,矿化围绕着中酸性次火山岩分布,与 英安玢岩以及隐爆角砾岩关系密切。其中火山-次 火山活动发生于 110~100 Ma,也是紫金山-二庙沟 铜、金主要成矿期(Li Bin et al., 2017), 与本文研究 的治岭头多金属矿床形成时代相近。罗卜岭 Cu-Mo 矿形成于 105 Ma(Zhong Jun et al., 2014),其 与福安赤路斑岩钼矿床(105~106 Ma, Zhang Keyao et al., 2009)、石菉和天堂斑岩-砂卡岩型铜 钼多金属矿(102~104 Ma, Zheng Wei et al., 2013a; Zheng Wei, 2016) 形成时代相一致。粤东莲 花山斑岩钨矿矿化赋存在石英斑岩的内外接触带, 具有斑岩型矿床的蚀变分带(钾化、云英岩化、绿泥 石化、绢云母化和青磐岩化),形成于133 Ma(Liu Peng et al., 2018)。粤东主要形成与斑岩-次火山 岩有关的锡矿称为长埔式。长埔锡多金属矿床是一 中型热液脉型锡多金属矿床,与石英斑岩(145 Ma) 密切相关,蚀变类型主要包括电气石化、硅化、绿泥 石化、绢云母化、碳酸盐化(Qiu Zengwang et al., 2016)。同类型的与高分异 I 型花岗岩有关的金坑 Sn-Cu-Pb-Zn 矿形成于 139~145Ma 之间(Qiu Zengwang et al., 2017b), 以及塌山 Sn-Pb-Zn 矿形 成于 138 Ma(Shen Weizhou et al., 1994)。东南沿 海的白垩纪火山岩很发育,强烈喷发的陆相火山岩 一般伴随有次火山岩,并且保存有发育良好的火山 机构,有利于斑岩型铜、金多金属矿床和浅成低温热 液型金矿床的形成与保存,但目前已知的矿床很少, 是已经被剥蚀?还是由于勘查力度不够、自然条件 等外部原因导致?

(3) 浙闽粤晚白垩纪与火山岩有关的 Pb、Zn、 Ag 以及非金属矿床成矿系列。包括浙江后岸银矿 床、千官岭金银矿床、大岭口银铅锌矿床、五部式铅 锌银矿、青田式叶蜡石矿床、矾山式明矾石矿床、老 虎头式沸石矿床和武义式萤石矿床等。浙江后岸银 矿床形成于 75~85 Ma 之间,大岭口银铅锌矿床形 成于 100 Ma(Yin Jiaheng et al.,1989)。浙江东部 与火山岩有关的非金属超大型矿床,如武义盆地内 的萤石矿矿化年龄为 80~95 Ma(Zhang Yongjia, 1995)。

(4)浙闽粤与晚白垩纪侵入岩有关的 W、Sn、 Mo、Bi 多金属矿床成矿系列。包括广东鹦鹉岭锡 钨多金属矿的形成时代为 83.0±1.7 Ma(Zheng Wei et al., 2013a),阳春脉状锡矿为 76 Ma(Yu Jinsheng et al., 1988),银岩斑岩型锡矿的矿化年龄 为 78.65±0.98 Ma(Zheng Wei et al., 2016),大金 山钨锡矿辉钼矿的 Re-Os 年龄为 82.5±3.1 Ma (Yu Changfa et al., 2012),锡坪钼多金属矿床的辉 钼矿 Re-Os 年龄为 89.9±3.4 Ma(Zheng Wei et al., 2017),小南山以及锡山钨锡矿的形成时代分别 为 78.3±2.7 Ma 和 79.41±1.11 Ma(Zheng We et al., 2015; Zheng Wei et al., 2017)。综上所述, 白垩纪晚期粤西地区发生了大规模的钨锡矿化。

5.2 成矿物质来源

成矿物质来源是矿床地球化学研究中的一项重 要内容,也同时制约着成矿地质体的确定。硫是矿 床中成矿金属元素迁移、富集、沉淀的重要矿化剂, 通过硫同位素组成的研究,可以帮助了解和判断成 矿物质的来源(Ohmoto,1986),进而探讨矿床成因 (Zhang Ligang et al.,1985; Zheng Wei et al., 2012)。目前硫同位素组成特征已经成为判断硫源 及成矿物质来源的主要工具,然而对于矿床中硫的 来源讨论必须根据硫化物沉淀期间热液的总硫同位 素组成加以探讨分析(Zheng Yongfei et al.,2000)。 由于热液成矿作用过程中的硫同位素分馏效应取决 于体系总硫同位素组成、离子强度、氢离子浓度、氧 逸度及温度等多方面因素,因此由热液成矿作用形 成的硫化物和硫酸盐的 δ³⁴S值一般不等于热液总 的 δ³⁴S 值 (Ohmoto, 1972)。但是在一定条件下可 以根据矿床的矿物组合关系估计成矿流体的硫同位 素组成。Ohmoto(1972)提出如果热液体系中不存 在硫酸盐类矿物,同时矿物组合比较简单时,硫化物 的 δ³⁴S 值可以大致代表热液的总硫同位素组成。 Zhang Ligang et al., (1985)和 Zheng Yongfei et al.,(2000)认为在硫同位素分馏达到平衡的条件 下,共生硫化物的δ³⁴S值按辉钼矿→黄铁矿→磁黄 铁矿→闪锌矿→黄铜矿→方铅矿的顺序递减。治岭 头钼多金属矿床发育的硫化物主要为黄铜矿、辉钼 矿、黄铁矿和斑铜矿,矿物组合比较简单,反映出成 矿热液中不同价态的硫之间以及不同成矿阶段的硫 同位素分馏较弱。同时,治岭头矿床矿石矿物硫同 位素值δ³⁴S显示辉钼矿>黄铁矿(片麻岩中)>闪 锌矿 > 方铅矿的富集规律,表明含矿流体在成矿过 程中达到硫同位素分馏平衡(图 5),硫化物的硫同 位素基本能代表成矿流体的硫同位素组成。硫化物 的硫同位素组成变化范围较窄(-2.1‰~2.6‰), 表明治岭头钼多金属矿床硫同位素可能主要来自上 地幔或下地壳的深源岩浆,具有相对均一的硫源,但 也可能受到陆壳沉积物源硫的混染。

金属硫化物中通常含有一定量的 Pb,且 U、Th 含量很低,所以金属硫化物结晶以后通过衰变作用





Fig. 5 The δ^{34} S values from the Zhilingtou polymetallic deposit in Zhejiang Province (References for the ranges of major sulphur reservoirs are from Li Shengrong et al., 2014)



图 6 治岭头钼多金属矿床的赋矿岩石和矿石²⁰⁸ Pb/ ²⁰⁴ Pb-²⁰⁶ Pb/²⁰⁴ Pb (a) 和²⁰⁷ Pb/²⁰⁴ Pb-²⁰⁶ Pb/²⁰⁴ Pb (b) 构造模 式图解(Zartman et al., 1981)

Fig. 6 ²⁰⁸ Pb/²⁰⁴ Pb-²⁰⁶ Pb/²⁰⁴ Pb (a) 和²⁰⁷ Pb/²⁰⁴ Pb-²⁰⁶ Pb/²⁰⁴ Pb-²⁰⁶ Pb/²⁰⁴ Pb (b) tectonic pattern illustration of ore-forming rocks and ore in the Zhilingtou molybdenum polymetallic deposit(after Zartman et al., 1981)

所产生的放射性成因铅的含量非常低,对硫化物铅 同位素组成的影响可以忽略不计(Wei Juying et al.,1988)。同时利用岩浆岩、矿石、地层和基底铅 同位素综合对比的方法来研究矿石铅来源,能够更 精确地探析矿床成矿流体的来源,因此 Pb 同位素 也被广泛应用于成矿物质来源的示踪研究。²⁰⁸ Pb/²⁰⁴ Pb-²⁰⁶ Pb/²⁰⁴ Pb留同位素构造模式图解(图 6a) 显示该矿床的矿石铅比较集中,落在造山带和下地 壳之间。同样在²⁰⁷ Pb/²⁰⁴ Pb-²⁰⁶ Pb/²⁰⁴ Pb 铅同位素 构造模式图解(图 6b)中,矿石铅也比较集中,落在 地幔和上地壳之间。

同时,特征值中μ值等的变化也可以提供地质 体经历的某些地质作用的信息,反映铅的来源。从 表 3 中由两阶段铅演化模式计算的参数(μ、w、Th/ U)看,不同样品的 μ 、w、Th/U模式值均有一定的变化,说明其来源并不单一,无论是黄铁矿、辉钼矿,还是方铅矿、闪锌矿,它们的铅同位素均具有非单一阶段的演化,表明可能是有不同的源区或在演化过程中有不同源区物质的混入(Meng Xiangjin et al., 2006)。一般认为来自下部地壳或者上地幔的铅 μ 值比较低(Zhu Shangqing et al.,1988)。据 Doe et al. (1979)资料显示,地幔环境的 μ 值为 8.92,造山带 μ 值为 10.87。由表 3 可知,治岭头钼多金属矿床的矿石矿物铅同位素 μ 值变化较小,比较集中, μ 值从 9.13~9.87,明显高于地幔 μ 值(8.92)而低于造山带 μ 值(10.87),同样显示铅的来源可能为壳幔混源。其他特征值如 μ 、w、Th/U、U/Pb 和 Th/Pb 比值均基本相同,反应来源的相似性(Fei Guangchun et al.,2011)。

Re-Os 同位素体系不仅可以精确确定矿床形成 的时间,还可以示踪成矿物质来源以及指示成矿过 程中不同来源的物质混入的程度。Mao Jingwen et al. (1999)和 Stein et al. (2001)的研究认为,从壳 源、壳幔混源到幔源,辉钼矿的 Re 含量各递升一个 数量级,从 n×10⁻⁶→n×10⁻⁵→n×10⁻⁴ (Mao Jingwen et al.,2003)。由表1可以看出,治岭头钼 多金属矿床6个辉钼矿的 Re 含量基本都在16.05 ×10⁻⁶与 66.29×10⁻⁶之间,同样显示该矿床的成 矿物质可能来自于壳幔混合来源。

5.3 成矿机制

矿区内与多金属矿化紧密共生的围岩主要为八 都群变质岩、白垩纪火山岩、花岗斑岩及各种脉岩。 Zheng Minghua et al. (1986)研究发现代表八都群 原始含金背景的矿区外围大面积变质岩,以含金平 均值低和数据集中为特征,但当夹有相当数量的火 山物质时,其含量则可能升高。能够代表矿区变质 岩含金原始丰度的远矿无蚀变变质岩,含金量普遍 较高,其平均值为11.9×10⁻⁹,是区域背景值的六 倍,是地壳克拉克值的三倍,而此地段变质岩的含银 量更为突出(平均达 600×10⁻⁹)。可见,原岩以火 山物质为主的治岭头矿区变质岩,是八都群地层大 面积贫金的背景上,金银等成矿元素初步聚集的原 始异常地段,表明金银矿床在该区产出绝非偶然。 曾受矿液活动影响的矿带两侧的变质岩,含金量则 显著降低,平均仅 4.77×10⁻⁹。据此可以推断,此 地段变质岩中的金曾发生过大规模的迁移。Hua Jiexiong et al. (2000)对矿区内不同地质体开展成 矿元素分析,结果显示流纹岩、花岗斑岩、霏细斑岩 的 Pb、Zn 平均含量高于其他地质体,比维氏地壳酸 性岩 Pb、Zn 丰度值高几倍-几十倍。显然这些岩浆 分异形成的热液具备提供成矿元素的条件。结合本 次 S、Pb 同位素测试结果,可以推测铅锌矿化、钼矿 化与该地区早白垩世晚期(110 Ma)的火山-岩浆活 动有关。Zhao Chao et al. (2014)对治岭头斑岩钼 矿体流体包裹体研究表明,成矿期发生了强烈的流 体沸腾作用,导致钼矿化富集沉淀,同时发育大量的 钾化和黄铁绢英岩化,为典型的斑岩型矿化。氢氧 同位素研究表明该成矿流体由成矿前、成矿期的岩 浆热液演化为成矿后的大气降水为主。Pu Weimin et al. (2008)和 Mao Weixiong et al. (2001)针对隐 爆角砾岩型铅锌矿和脉状铅锌矿的流体包裹体研究 表明,成矿热液为中低温中低盐度,流体混合可能是 铅锌矿化沉淀的主要原因。

135 Ma之后,由于太平洋板块沿 NNE 向走滑 断裂带向北大规模走滑,使得中国大陆处于持续伸 展阶段(Mao et al., 2008a, 2008b)。在早白垩世晚 期,治岭头矿区同样处于拉张伸展环境,软流圈上涌 引发上地幔部分熔融,所产生的玄武质岩浆侵位于 下地壳底部,诱发中下地壳的广泛熔融而形成含成 矿物质的酸性岩浆房,早期受到酸性岩浆屏蔽作用 的阻碍而难以继续上升,主要是酸性岩浆沿断裂上 升并喷发到地表,形成隐爆角砾岩筒和断裂组成的 火山机构,成为良好的导矿-容矿空间(Wang Yongbin et al., 2014); 后期, 玄武质岩浆和酸性岩 浆混合形成的花岗质岩浆,其演化分异出携带大量 金属元素的高温、高盐度的 H₂O-NaCl 流体发生沸 腾作用,形成斑岩钼矿化(Zhao Chao et al., 2014)。 同时,一部分流体沿隐爆角砾岩筒向上迁移,形成角 砾岩型铅锌矿;一部分流体沿火山机构伴生断裂向 上迁移,与加热循环的大气水混合,形成脉状铅锌矿 (Wang Yongbin, 2014)。治岭头矿区斑岩钼矿、浅 成低温热液型铅锌矿和远程热液型金银矿共存,在 垂向上和平面上均具有从斑岩体向外出现斑岩钼 矿、浅成低温热液型铅锌矿和远程热液型金银矿的 分布规律。在岩浆活动中心的外围断裂系统,尤其 是古韧性剪切带中的引张部位,是寻找远程热液型 金矿的重要部位。

6 结论

(1)治岭头钼铅锌金多金属矿床的6件辉钼矿 样品模式年龄的变化范围为113.7~114.6 Ma,等 时线年龄为114.3±1.1 Ma,表明成矿作用发生于 早白垩世晚期。

(2)治岭头矿床辉钼矿的¹⁸⁷ Re 含量为 16.05× $10^{-6} \sim 66.29 \times 10^{-6}$,显示成矿物质可能主要来自于 壳幔混源。矿石硫同位素组成变化范围较窄 (-2.1‰~2.6‰),达到同位素分馏平衡,可能主要 来自上地幔或下地壳的深源岩浆,具有相对均一的 硫源,但也受到陆壳沉积物源硫的混染。治岭头钼 铅锌金多金属矿床矿石铅同位素组成变化范围比较 小,²⁰⁸ Pb/²⁰⁴ Pb、²⁰⁷ Pb/²⁰⁴ Pb、²⁰⁶ Pb/²⁰⁴ Pb 的范围分别 为 38.765 ~ 39.137、15.523 ~ 15.751、18.450 ~ 18.667,具有明显的壳幔混合特征。

(3)治岭头矿区斑岩钼矿、浅成低温热液型铅锌 矿和远程热液型金矿共存是同一构造-岩浆-热事件 的产物,成矿时间具有一致性,成矿空间具有明显的 分带性:即从岩体内外接触带向外或向上依次为斑 岩钼矿、浅成低温热液型铅锌矿和远程热液型金矿, 构成一个矿床组合模型。在这个模型中,不同类型 矿床互为找矿标志。在岩浆活动中心的外围断裂系 统,是寻找远程热液金矿的重要部位。

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Sulfur and lead isotope tracing for sources of ore-forming material and ore-forming age of the Zhilingtou Mo-Pb-Zn-Au polymetallic deposit in the Zhejiang Province

CHU Kelei^{*1,2)}, CHEN Xiaorong³⁾, QI Gang³⁾, GAO Xiang³⁾, HU Bo³⁾, LI Shengjin³⁾

1) Chinese Academy of Geological Sciences, Beijing, 100037;

2) State Key Laboratory for Mineral Deposits Research, Nanjing University, Nanjing, 210093;

3) No. 7 Geological Party of Zhejiang Province, Lishui, Zhejiang, 323000

* Corresponding author: chukelei@163.com

Abstract

The Zhilingtou deposit has a "three-story" mode: the upper part (volcanic rock caprock) is dominated by lead-zinc deposits and sulfur deposits, the middle section (metamorphic rock basement) by gold-silver and lead-zinc deposits, and the deep part (intrusive rock) by porphyry molybdenum deposits. Based on the previous studies, this paper determined the ore-forming age of the Zhilingtou deposit by molybdenite Re-Os isotope dating. The model age of the six molybdenite samples was 113.7 \sim 114.6 Ma, with a weighted average age of 114.03 \pm 1.78 Ma, indicating that the mineralization of the Zhilingtou Mo polymetallic deposit occurred in the late stages of the Early Cretaceous. The Re content of the molybdenum sample in this deposit changed from 16.05×10^{-6} to 66.29×10^{-6} , indicating that its ore-forming material has the characteristics of crust-mantle mixed source. The sulfur composition of sulphide has a narrow range of changes $(-2.1\% \sim 2.6\%)$, which reaches the equilibrium of isotope fractionation and may be derived mainly from the deep magma of the upper mantle or the lower crust. It has a relatively uniform source of sulfur, but may also be affected by the continental crust. The range of lead isotope composition of the ore minerals from the Zhilingtou Mo-Pb-Zn-Au polymetallic deposit is relatively small. The values of ²⁰⁸ Pb/ ²⁰⁴ Pb, ²⁰⁷ Pb/²⁰⁴ Pb, and ²⁰⁶ Pb/²⁰⁴ Pb are 38.765~39.137, 15.523~15.751, 18.450~18.667, respectively, indicate crust-mantle mixing characteristics. The determination of the ore-forming age of the Zhilingtou Mo polymetallic deposit provides an important clue for the prospecting of contemporary porphyry oreforming systems in the region, and provides new information for further studying the petrogensis and metallogeny of the southeast coastal metallogenic belt.

Key words: Zhilingtou Mo-Pb-Zn-Au polymetallic deposit; molybdenite Re-Os dating; sources of oreforming materials; southeast coastal metallogenic belt

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