

罗布泊第四纪湖盆物源与盆地构造演化特征： 来自 LDK01 孔碎屑锆石 U-Pb 年龄证据

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内容提要:为揭示罗布泊盐湖第四系潜在物源区及凹陷阶段性演化过程, 本文对罗布泊第一口钾盐科探深钻 LDK01 孔更新统不同深度样品进行碎屑锆石 U-Pb 微区定年分析。Th/U 比值显示, 钻孔碎屑锆石类型主要为岩浆锆石, 少量为变质成因锆石。锆石年龄主要集中在 209~240Ma、265~304Ma、320~385Ma、406~446Ma、705~880Ma 及 2376~2405Ma 几个区间。综合分析潜在物源区的岩石属性和年龄构成, 初步认为罗布泊地区前寒武纪年龄来自北部山前库鲁克塔格地区, 加里东期碎屑锆石可能来源于阿尔金造山带和(或)南天山构造带。276Ma 的峰值记录了塔里木盆地二叠纪大火成岩省事件, 南天山最有可能为主导物源区。印支期和新生代碎屑锆石年龄暗示了北山地块和东天山, 甚至较远的帕米尔—西昆仑山等地可能也提供了物源。凹陷周缘富钾岩体广泛出露, 经风化、淋滤搬运至罗布泊, 为凹陷第四纪成钾提供了有利的物质来源。塔里木盆地内部流域带来的碎屑组分是主要的物质来源, 近源地区造山带岩体提供的物源有限。碎屑锆石年龄纵向变化特征显示, 罗布泊北部地区在中更新世发生一次明显的构造抬升, 可能是导致罗北凹地形成的重要原因。

关键词:罗布泊; 物源; 构造演化; 碎屑锆石

盆地充填的沉积物是流域水系范围内造山带岩石经物理和化学风化、剥蚀后, 通过风力和河流搬运作用, 最终沉积和埋藏的产物, 真实记录了沉积物充填过程中盆地的动力学性质和周围造山带构造活动特征(Dickinson et al., 1983; Yan Yi et al., 2002; Dai Jingen et al., 2012)。锆石具有极好的抗风化、抗磨蚀和热蚀变的能力, 在沉积循环中不易被破坏且不受分馏过程的影响, 是反映沉积物源区的良好示踪剂(Morton et al., 1996)。沉积岩中碎屑锆石年代学研究是识别盆地物源区贡献、重现构造演化和重建古地理的重要手段(Cawood et al., 2000; Berry et al., 2001; Fonneland et al., 2004; Anderson, 2005; Leier et al., 2007; Cheng Jun et al., 2017)。近年来, 随着高分辨率电子探针和激光剥蚀等离子质谱仪的发展, 碎屑锆石定年手段已经得到长足的进步和广泛应用(Li Zhong et al., 2013; Pereira et al., 2016)。如 Hietpas et al (2011) 采集美国东南部阿巴拉契亚山脉

(Appalachian Orogen) 的佛兰西布罗德河及其支流中的不同河段的现代沉积物, 进行碎屑锆石 U-Pb 定年, 发现沿着河流方向样品的年龄谱发生了很大的变化。Li Zhong et al. (2010) 和 Yang Wei et al. (2013) 通过对比南北天山前麓沉积盆地的不同时代的碎屑锆石特征值变化, 还原了南北天山自始新世以来逐渐分异的隆升剥蚀历史。在研究中国黄土高原、塔克拉玛干沙漠的物源追踪等热点问题上, 碎屑锆石亲缘性图解(MDS) 同样提供了有力的证据(Nie Junsheng et al., 2014; Rittner et al., 2016)。

罗布泊位于塔里木盆地东部, 处于塔里木地块、东天山褶皱带和北山褶皱带的交汇处。作为全球著名的干盐湖之一, 罗布泊凹陷第四纪末期形成了超大型卤水钾矿(Wang Mili et al., 2001), 深部地层蕴藏着丰富的低品位固体钾盐资源(Jiao Pengcheng et al., 2014)。新生代以来, 受新构造活动影响, 罗布泊成为塔里木“高山深盆”的次级深盆(Wang Mili et al., 2005), 演变为盆地沉积物质最终的汇集地。

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前人对罗布泊的研究主要集中在其钾盐矿床成因机制(Wang Mili et al., 2005; Liu Chenglin et al., 2006, 2015)、第四纪环境演化(Liu Chenglin et al., 2016)等。塔里木流域河流地球化学特征(Gu Xinlu et al., 2003; Liu Chenglin et al., 2010; Bo Ying et al., 2013; Liu Chenglin et al., 2015; Xiao Jun et al., 2015)和卤水同位素组成(Liu Chenglin et al., 1999)分析表明罗布泊盐湖钾离子来自周围山区含钾岩石风化淋滤作用;硫同位素研究表明湖水中硫酸根主要来自塔里木盆地西部及南天山中新生界古盐矿的风化产物(Wang Mili et al., 2001)。此外,受新生代构造活动影响,罗布泊周缘第四纪后期改造、破坏较为严重,长序列、连续的露头较为缺乏,制约了罗布泊凹陷第四纪演化研究,罗布泊盆地第四系碎屑岩建造的物源补给及其对周缘构造活动的地质响应等研究仍较为薄弱。因此,本文首次尝试对罗布泊罗北凹地内一口钻孔的更新统地层砂岩样品中的碎屑锆石进行 U-Pb 定年分析,结合罗布泊周缘造山带岩石属性和年代学及盆内流域水系单元调研结果,对罗布泊地区第四纪沉积物物源特征进行探索,探讨凹陷阶段性演化过程及其对周缘山体构造活动的构造响应。

1 区域地质

罗布泊凹陷位于塔里木盆地东部,其北部为近东西走向的库鲁克塔格山,东部为北东东走向的北山断块,南部为北东东走向的阿尔金山,西面为塔克拉玛干沙漠(图 1)。罗布泊凹陷南北分别受控于呈右行走滑的库鲁克塔格南缘断裂伴生的孔雀河断层和呈左行走滑的阿尔金断裂伴生的若羌断层,新构造运动强烈(Guo Zhaojie et al., 1995; Xia Xuncheng, 1997),深大断裂及褶皱发育,岩浆喷发及侵入活动频繁。

罗布泊凹陷主要地层为发育在太古界一下元古界中深变质岩系基底之上的几套沉积盖层、火山岩系及第四系沉积物(BGMRXUAR, 1993)。罗布泊汇水区范围内,前第四系地层主要分布在库鲁克塔格、阿尔金山和北山(图 1)。在库鲁克塔格,前古生界下部主要岩性为混合岩、片岩、片麻岩及大理岩,往上则为灰岩、大理岩等碳酸盐建造及碎屑岩建造,顶部为由碎屑岩、泥板岩、火山岩及泥岩组成的冰期沉积物。阿尔金山区为变粒岩、片麻岩、变质砂岩、粉砂岩、绿片岩、灰岩、大理岩。北山地区,下部为混合岩、片麻岩、片岩,上部为大理岩、砂岩。古生界地

层在库鲁克塔格山区,下部为灰岩,上部为砂岩、泥岩。阿尔金山区,主要为灰岩、砂砾岩。在北山,下部为灰岩、大理岩,中部为凝灰岩、安山岩及砂岩,上部为泥岩。新生界第三系为陆相沉积。

2 样品采集

本次测试样品来源于罗北凹地 LDK01 孔。该钻孔是由国投新疆罗布泊钾盐有限责任公司出资,中国地质科学院矿产资源研究所非金属室负责实施的罗布泊第一口钾盐科探井($40^{\circ}55'N, 90^{\circ}55'E$)。钻孔深度 781.5m,钻遇地层主要为下更新统西域组(未见底)、中更新统乌苏群以及上更新统新疆群(Lü Fenglin et al., 2015)。本文选取了钻孔中 6 个不同深度的碎屑砂岩(N2、N51、N80、N133、N215、N221)进行锆石 U-Pb 年龄测试和分析,其对应深度分别为 20.5m、72.5m、104.9m、212.05m、458.33m 以及 735.99m(图 2)。

3 测试方法

测试之前,将新鲜的砂岩样品粉碎至 80 目以下,经过人工淘洗和电磁选方法筛选锆石,之后在双目镜下仔细挑选,尽量避免包裹体多和有裂痕的锆石,将晶型较好的锆石颗粒粘在双面胶上,灌上环氧树脂制成样品靶,将靶子表面抛光,露出锆石平面。锆石的分选工作和部分阴极发光照片(CL)在北京铅年领航科技有限公司完成。

锆石 U-Pb 定年测试分析在中国地质科学院矿产资源研究所 MC-ICP-MS 实验室完成,锆石定年分析所用仪器为 Finnigan Neptune 型 MC-ICP-MS 及与之配套的 Newwave UP 213 激光剥蚀系统。激光剥蚀所用斑束直径为 $25\mu\text{m}$,频率为 10Hz,能量密度约为 $2.5\text{J}\cdot\text{cm}^{-2}$,以 He 为载气。LA-MC-ICP-MS 激光剥蚀采用单点剥蚀的方式,数据分析前用锆石 GJ-1 进行调试仪器,使之达到最优状态,锆石 U-Pb 定年以锆石 GJ-1 为外标,U、Th 含量以锆石 M127(Nashala et al., 2008)为外标进行校正。均匀锆石颗粒 $^{207}\text{Pb}/^{206}\text{Pb}$, $^{206}\text{Pb}/^{238}\text{U}$, $^{207}\text{Pb}/^{235}\text{U}$ 的测试精度(2σ)均为 2%左右,对锆石标准的定年精度和准确度在 1% (2σ)左右。详细的实验测试过程参见 Hou Kejun et al. (2009)。

锆石 U-Pb 年龄数据处理采用 ICPMSDataCal 程序(Liu Yongsheng et al., 2010),锆石年龄谐和图用 Isoplot 3.0 程序获得。样品分析过程中,Plesovice 标样作为未知样品的分析结果为 337Ma,

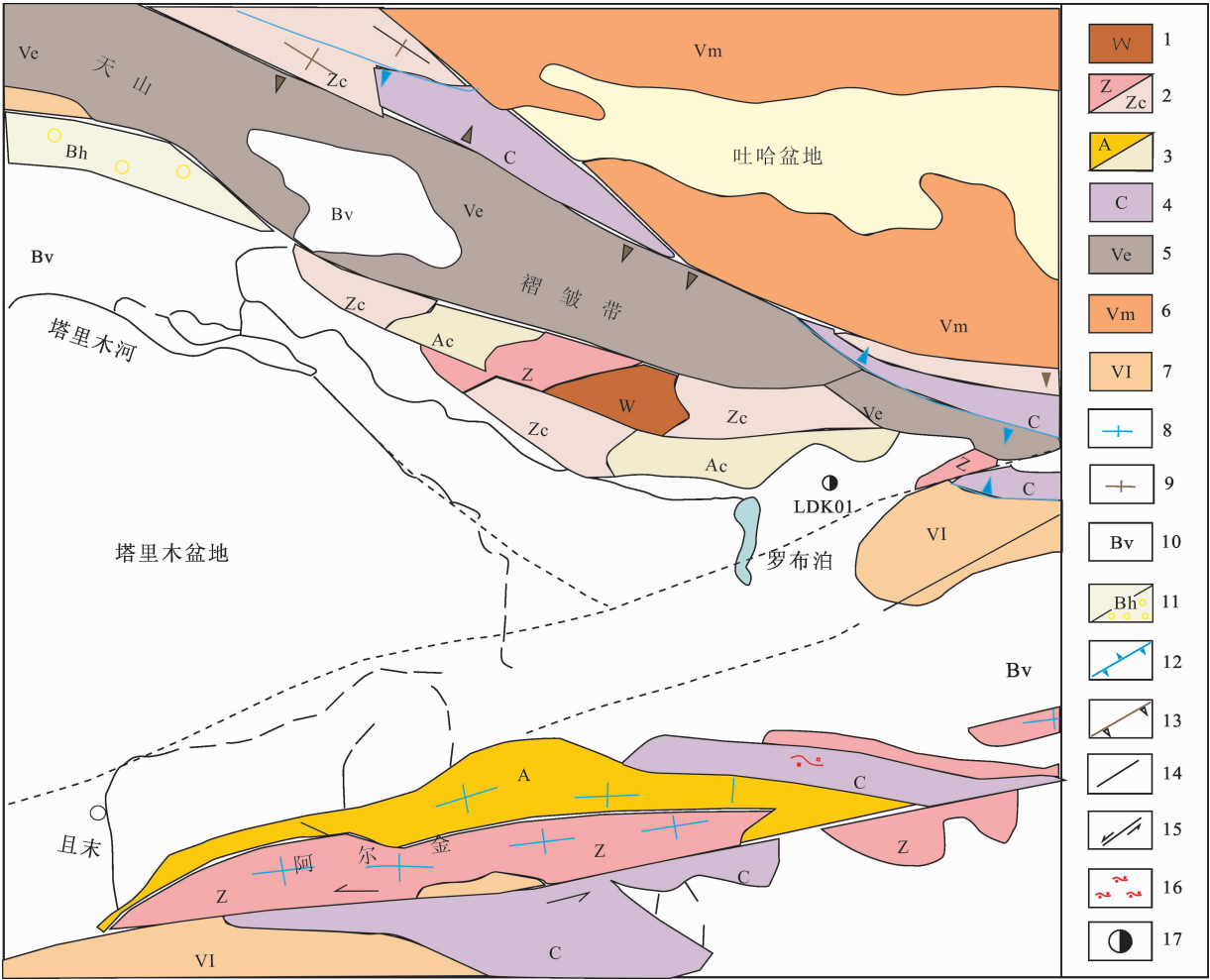


图 1 新疆罗布泊及周边地区地质简图(中国大地构造图,1 : 1200 万)及 LDK01 钻井位置
Fig. 1 Geological sketch map of Lop Nur area (China geotectonic map, 1 : 12000000),
Xinjiang and the location of the borehole LDK01

1—五台褶皱带(>2400Ma)及更老的地壳;2—中条(吕梁)褶皱带(1900~1700Ma)及盖层;3—扬子(晋宁)褶皱带(1000~800Ma)及盖层;
4—加里东褶皱带及盖层(主造山期:D/S);5—早华力西褶皱带(主造山期:C₁₋₂前);6—中华力西褶皱带(主造山期:C₃前);7—晚华力西褶皱带(主造山期:P₂/P₁);8—强烈卷入加里东造山作用;9—强烈卷入华力西造山作用;10—晚华力西旋回以来的盆地;11—喜马拉雅旋回的盆地及山前磨拉石;12—加里东缝合带;13—华力西缝合带;14—断裂带;15—走滑断裂;16—蓝闪片岩带(高压变质带);17—钻孔位置
1—Wutai fold belt (>2400Ma)and older crust; 2—Zhongtiao (Lüliang) fold belt(1900~1700Ma) and covers; 3—Yangtze (Jinning)fold belt (1000~800Ma) and covers; 4—Caledonian fold belt and covers (main orogenic period:D/S); 5—Early Variscan fold belt (main orogenic period: before C₁₋₂); 6—Mid Variscan fold belt (main orogenic period: before C₃); 7—Late Variscan fold belt (main orogenic period P₂/P₁); 8—strongly affected by Caledonian orogenic; 9—strongly affected by Variscan orogenic; 10—basin formed since late Variscan; 11—basin and piedmont molasse formed during Himalayan cycle; 12—Caledonian suture zone; 13—Variscan suture zone; 14— fault belt; 15—strike-slip fault; 16—blue schist zone(high-pressure metamorphic zone); 17—borehole location

对应年龄推荐值为 $337.13 \pm 0.37 (2\sigma)$ (Sláma et al., 2008),两者在误差范围内一致。对于碎屑锆石年龄较老(>1000Ma)的样品,由于大量放射性成因 Pb 的存在因而采用 $^{207}\text{Pb}/^{206}\text{Pb}$ 表面年龄,而对于较年轻锆石(<1000Ma)的样品,由于可用于测量的放射性成因 Pb 含量较低和普通 Pb 校正的不确定性,因而采用更为可靠的 $^{206}\text{Pb}/^{238}\text{U}$ 年龄 (Sircombe et al., 1999)。以不谐和度 $\leq 10\%$ 为标

准遴选 U-Pb 年龄数据,共获得 367 个有效数据点。

4 分析结果

锆石年龄分布的范围介于 $33.4 \pm 0.4\text{Ma}$ 和 $2746 \pm 5\text{Ma}$ 之间,分别集中在 209~240Ma、265~304Ma、320~385Ma、406~446Ma、705~880Ma 和 2376~2405Ma(早元古代基底年龄)这六组区间(图 3)。

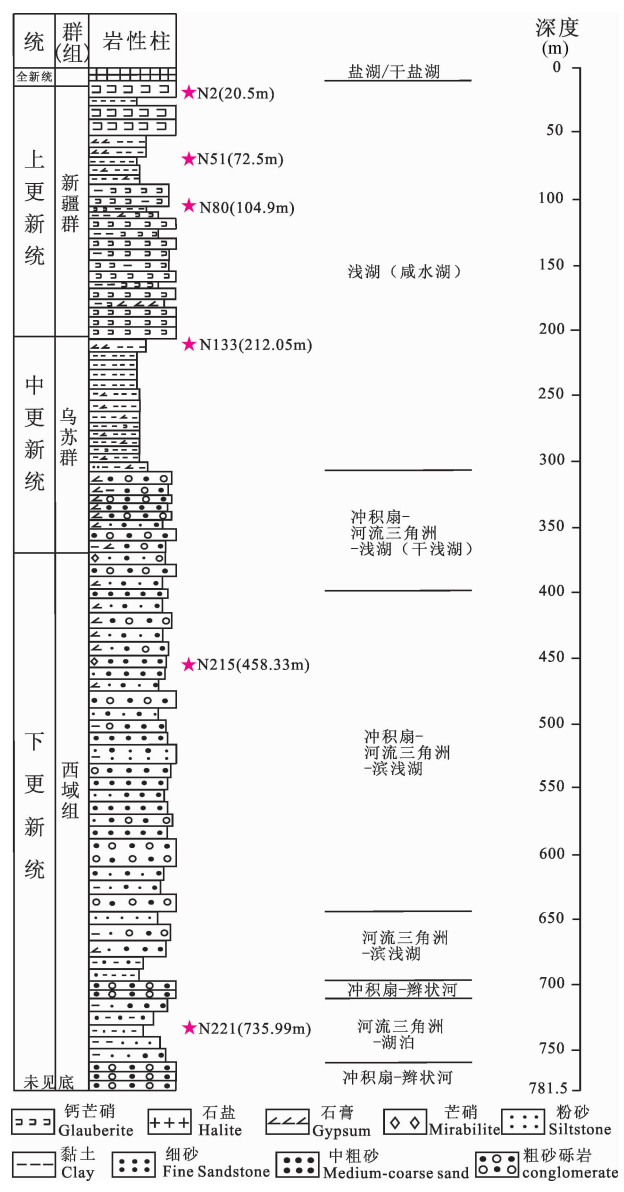


图2 新疆罗布泊 LDK01 孔岩性柱状图及采样深度
(据 Lü Fenglin et al., 2015 修改)

Fig. 2 Lithological column and sampling depth of LDK01 core, Lop Nur, Xinjiang (Modified from Lü Fenglin et al., 2015)

样品 N2 锆石年龄值变化于 $263 \pm 2 \sim 2746 \pm 5$ Ma, 集中分布于 $263 \sim 306$ Ma (占总有效数据的 33%)、 $420 \sim 472$ Ma (占总有效数据 21%) 和 $328 \sim 339$ Ma (占总有效数据 12.5%) 三个区间(图 3)。锆石颗粒形态(图 4)及 Th/U 比值(0.4~1.6)(图 5), 显示该样品中锆石具有岩浆锆石特征。样品 N51 锆石 U-Pb 年龄值分布变化于 $247.8 \pm 2.4 \sim 795 \pm 8.8$ Ma, 集中分布于 $279 \sim 304$ Ma (40%)、 $247 \sim 255$ Ma (12%) 和 $437 \sim 443$ Ma (12%) 三个区间(图 3)。在阴极发光下颗粒晶形完整, 呈圆状—浑圆状, 显示搬运磨蚀迹象(图 4)。整体 Th/U 比值介于

0.46~1.52 之间(图 5), 显示岩浆锆石特征。样品 N80 锆石晶形较为破碎, 浑圆状一次棱角, 半自形—不规则状较多(图 4)。年龄值分布变化于 $207.2 \pm 4.8 \sim 2387.3 \pm 7.7$ Ma 之间, 集中分布在 $271 \sim 309$ Ma (34%)、 $362 \sim 441$ Ma (28%) 两个区间(图 3, 图 6a)。样品 N133 碎屑锆石形态多次浑圆—棱角状, 粒径大小一般为 $30 \sim 80 \mu\text{m}$ (图 4), 整体 Th/U 比值介于 0.09~2.87 之间(图 5)。锆石年龄值分布较为广泛变化, 于 $33.4 \pm 0.4 \sim 2570 \pm 6$ Ma 之间, 明显的概率峰值出现在 265Ma、446Ma、414Ma、291Ma 等(图 3, 图 6b)。样品 N215 锆石晶型较为完整, 浑圆状—一次圆状, 自形—半自形, 粒径多为 $50 \sim 90 \mu\text{m}$ 。CL 图像显示, 以分析点 1 和 6 为典型的岩浆锆石振荡韵律环带(图 4e), 部分锆石呈面状分带(图 4e 分析点 2)均质无分带特征(图 4e 分析点 5)。整体 Th/U 比值介于 0.04~3.84 之间(图 5)。年龄值分布变化于 $40 \pm 0.5 \sim 2573 \pm 37$ Ma 之间。主要集中在 $205 \sim 295$ Ma (31%) 和 $391 \sim 443$ Ma (15%) 两个区间(图 3, 图 6c), 其余年龄分布较为广泛。样品 N221 锆石颗粒晶型较为完整, 次圆状—一次棱角状, 自形—半自形, 粒径较大, 多为 $40 \sim 110 \mu\text{m}$ 。CL 图像显示, 以分析点 1 和 6 为典型的岩浆锆石振荡韵律环带(图 4f); 分析点 3 呈边部变质重结晶锆石切割原岩岩浆锆石的环带(图 4f); 图 4f 分析点 5 呈典型面状分带的变质锆石特征, 其 Th/U 仅有 0.02。整体 Th/U 比值介于 0.01~3.21 之间(图 5)。年龄值分布变化于 $103 \pm 1.3 \sim 2670 \pm 9$ Ma 之间, 主要集中在 $209 \sim 289$ Ma (45%) 和 $367 \sim 407$ Ma (20%) 两个区间(图 3, 图 6d)。

图 7 为 LDK01 孔不同层位锆石年代与锆石类型分布对应图, 可以看出, 锆石数量集中在古生代和早中生代, 其中石炭纪、二叠纪和三叠纪年龄区间所占比例为 54%, 大部分表现为典型的岩浆锆石清晰的或稍弱的振荡环带特征, 古元古代—中元古代和古生代泥盆纪发现大量捕获老锆石或新生锆石核部的继承锆石, 表明该时期提供源区母岩性质主要以变质岩为主。

5 讨论

5.1 物源区岩石属性和年龄构成

样品 N80、N133、N215 中存在加权年龄峰值在 $2376 \sim 2405$ Ma 的锆石年龄(图 3), 部分锆石呈现出保留岩浆锆石形态的次生加大边(图 4), 可能为原生锆石经后期热液蚀变改造而成, 显示变质锆石的

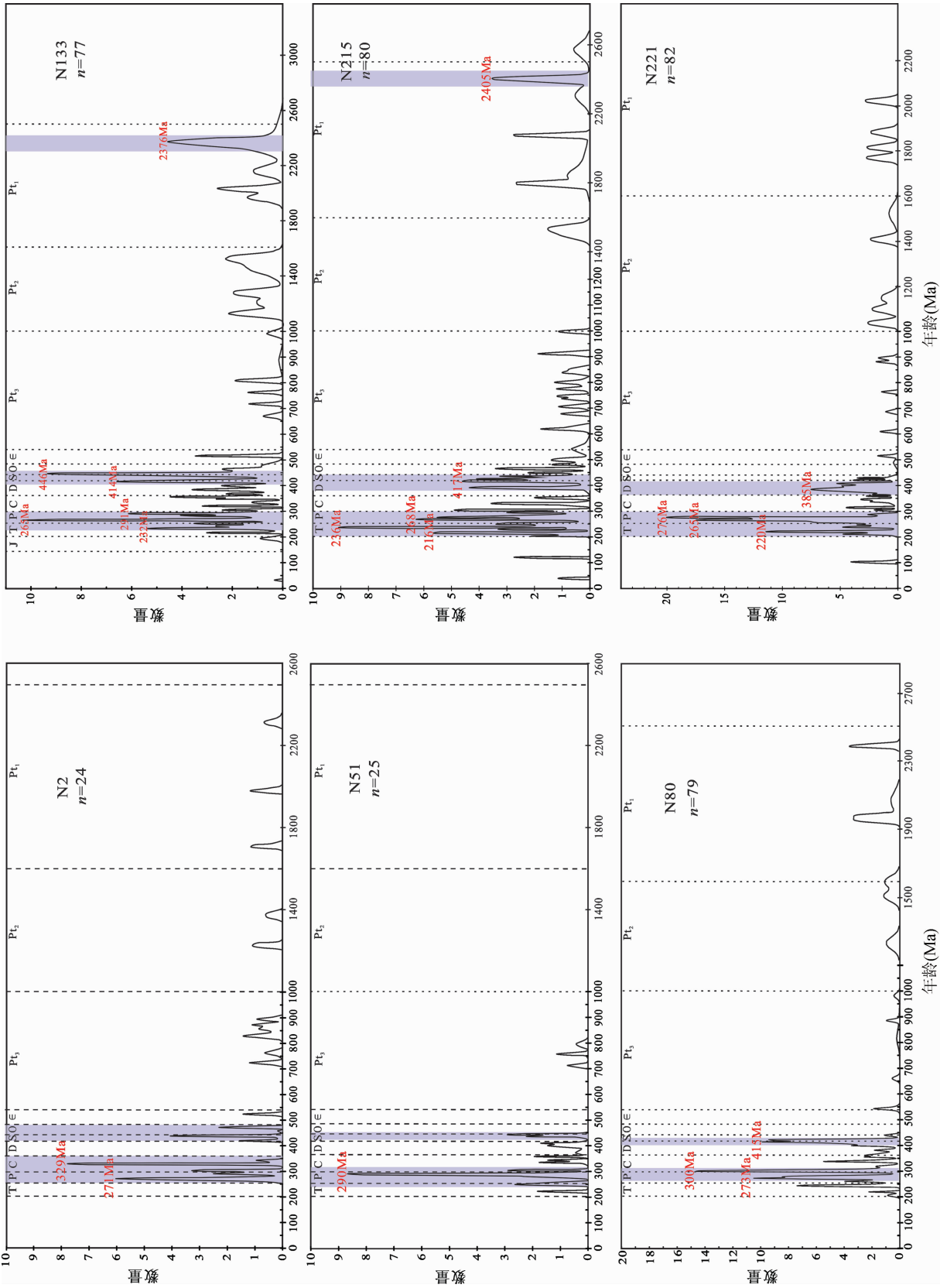


图3 LDK01孔碎屑锆石年龄分布图

Fig.3 Probability plots and number histograms of detrital zircons of U-Pb ages of LDK01 core

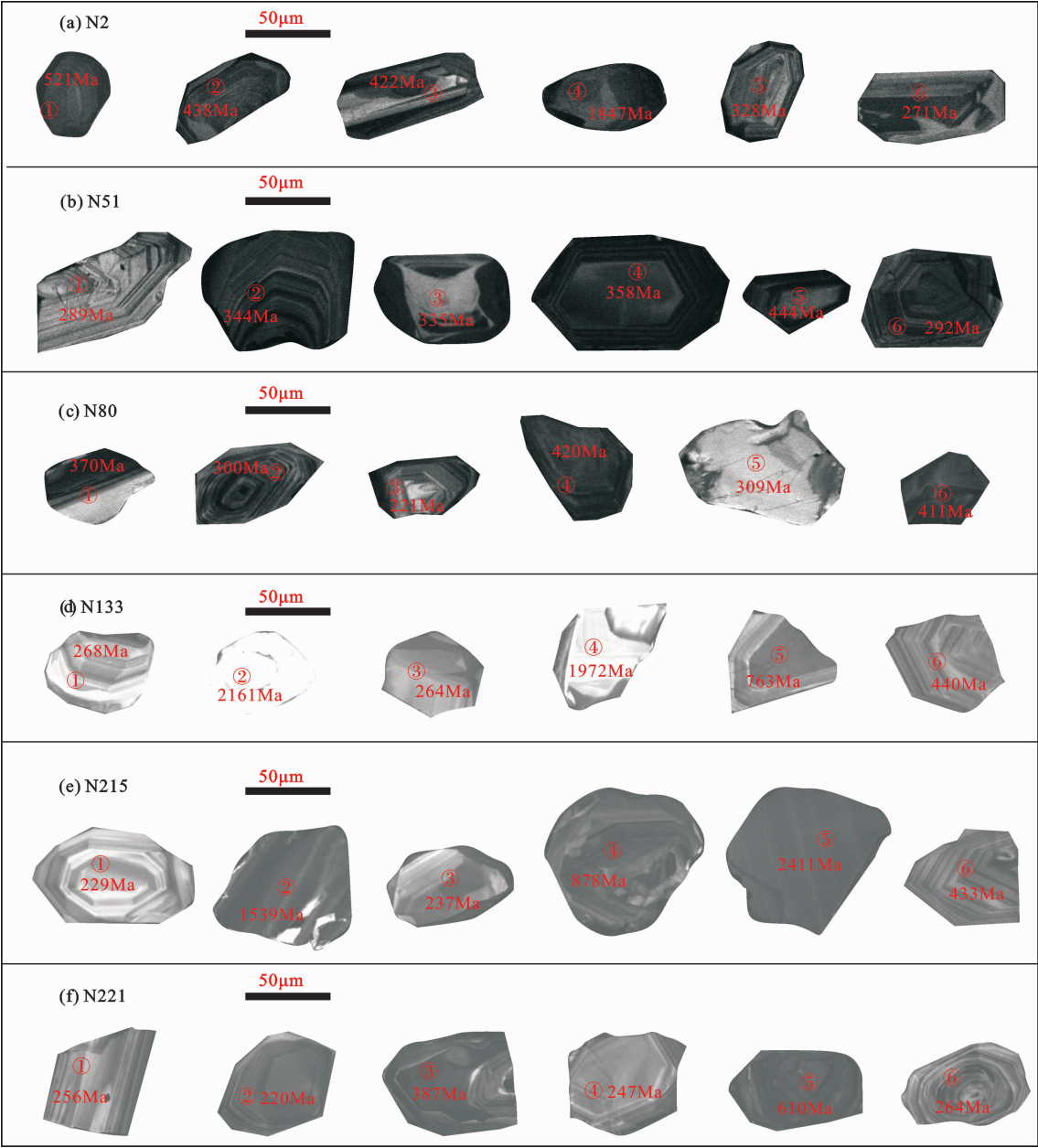


图 4 典型碎屑锆石 CL 图像

Fig. 4 CL images of detrital zircons of LDK01 core

特征 (Hoskin et al., 2000; Wu Yuanbao et al., 2004)。对比发现,这一年龄区间可能与全球大陆增生事件相关 (Guo Chuntao et al., 2015)。古元古代是塔里木盆地地壳快速增长的重要时期 (Tang Liangjie, 1996),塔里木盆地东北缘库鲁克塔格、兴地断裂南阔克苏以及辛格勒南部等地,分布有年龄 2308~2486Ma 的片麻状花岗岩、灰色片麻岩、斜长角闪岩和麻粒岩 (Hu Aiqin et al., 2006; Zhang Yingli et al., 2011; Shu Liangshu et al., 2011; Cao Xiaofeng, 2012),暗示了库鲁克塔格古元古代基底岩石对罗布泊物源的贡献。

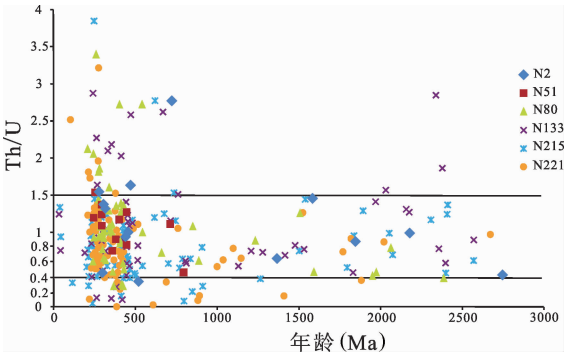


图 5 LDK01 孔碎屑锆石 U-Pb 年龄与 Th/U 比值

Fig. 5 U-Pb ages vs. Th/U ratios of detrital zircons

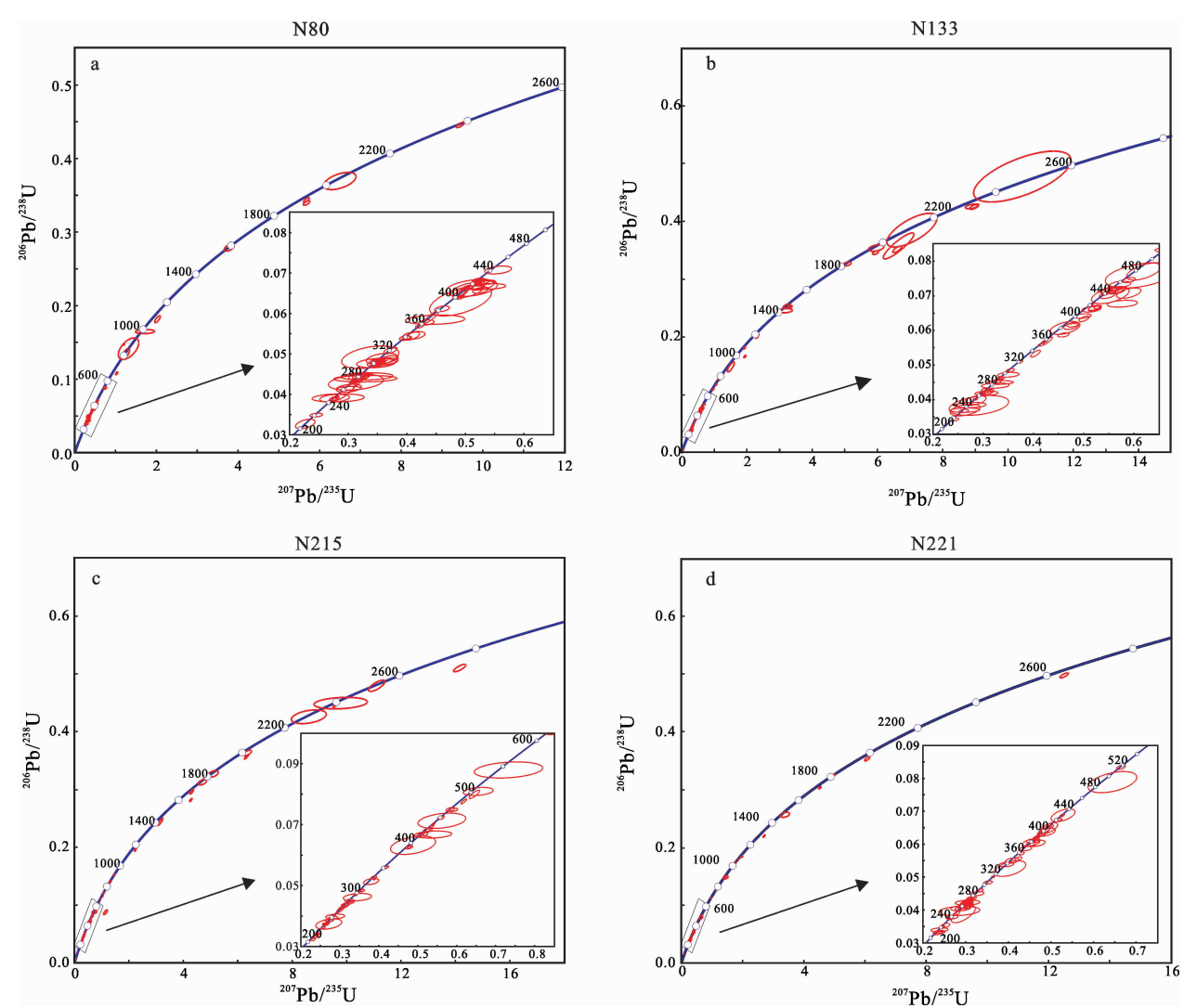


图 6 LDK01 孔 N80、N133、N215、N221 样品碎屑锆石谐和曲线和 U-Pb 年龄谱图
Fig. 6 Concordia plots for detrital zircons from N80、N133、N215、N221

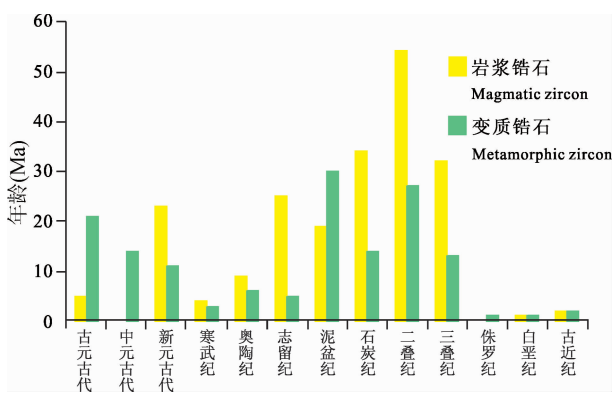


图 7 碎屑锆石时代划分与锆石类型分布
Fig. 7 The histogram distributions of the detrital zircon ages vs. types

713~913Ma 区间的锆石在 6 个样品中均有存在,主要表现为岩浆锆石的弱环带特征。该期构造事件在盆地内库鲁克塔格(Xu Bei et al. , 2009;

Zhang Yingli et al. , 2011)、阿克苏(Chen Yan et al. , 2004)、塔中(Li Yuejun et al. , 2003; Wu Guanghui et al. , 2010)、巴楚(Song Wenjie et al. , 2003)、铁克里克(Xiao Aifang et al. , 2010)等均有广泛的分布,可能与 Rodinia 超大陆裂解事件相关(Yang Jingsui et al. , 2004; Zeng Jianyuan et al. , 2006; Zhang Chuanlin et al. , 2012))。其中库鲁克塔格地块新元古代中期岩浆活动表现最为强烈,例如其东部大平梁地区出露的 830~800Ma 的富钾花岗岩及类埃达克岩(Cao Xiaofeng et al. , 2011),南华系底部发育 760Ma 的陆内裂谷火山岩(Lu Songnian et al. , 2003),均指示塔里木盆地北缘大量而广泛的岩浆活动。

460~400Ma 区间的锆石形态主要表现为岩浆锆石振荡环带特征,并呈现出 415Ma、422 Ma 和 446Ma 的加权年龄峰值(图 3),推测该年龄区间可

能对应早古生代的“天山洋”和“库地洋”俯冲消减和微板块的碰撞所产生的加里东中期运动(Wan Tianfeng et al., 2012)。Liu Jingyan et al. (2012)通过塔北隆起泥盆系中碎屑锆石年龄揭示出塔北周缘 460~414Ma 存在强烈的构造事件。库鲁克塔格地区钾长花岗岩体 U-Pb 锆石年龄为 $430.6 \pm 1.6\text{Ma}$ (Xiao Peixi et al., 2006)。巴音布鲁克岩体是南天山北缘代表性早古生代花岗岩体,主要侵入岩为二长花岗岩和钾长花岗岩,形成年龄为 464~424Ma(Wei Qiang et al., 2017; Li Ping et al., 2018)。此外,该时段锆石年龄亦在阿尔金断裂带红柳沟—冰沟地区大量出现,可能与奥陶—志留纪发生的大规模的构造事件与火成岩活动相关(Guo Zhaojie et al., 2003; Liu Yongjiang et al., 2003; Yang Yi, 2003),其中 461~410Ma 左右的俯冲—碰撞与后碰撞构造—岩浆事件(Sobel et al., 1999; Chen Xuanhua et al., 2002)、产出的混杂蛇绿岩及中酸性火山碎屑岩也可能向北为罗布泊输送了碎屑物质。

400~205Ma 年龄为本次研究中获得数量最多,峰值最为集中,主峰表现为 220、232、265、273、291、300、329、385Ma 等(图 3),塔里木盆内在这一时间段表现为以大型逆冲推覆隆升剥蚀和强烈岩浆活动为特色。其中 345~300Ma 年龄可能对应晚泥盆世—石炭纪期间古南天山洋盆闭合(Xia Linqi et al., 2006; Gao Jun et al., 2006),南天山造山带西段产出一套蛇绿岩、高压变质岩带(Li Yuejun et al., 2009)。另外, N2 样品中 329Ma 的峰值也可能与东天山石炭纪企鹅群火山岩建造有关(Li Yuan et al., 2011)。

晚石炭世—早二叠世是塔里木盆地岩浆活动最强烈的时期,与海西期 306~260Ma 南天山造山带的碰撞造山及后碰撞的拉张作用、大规模的 A 型花岗岩及岩浆—热液成矿活动紧密相关(Luo Jinhai et al., 2008; Zhu Zhixin et al., 2008)。尤其是早二叠世的大火成岩省岩浆活动, 292~285Ma 期间以形成大陆溢流玄武岩为特征,主要出露于柯坪地区及塔西南其木干等地区(Xu Yigang et al., 2014; Li Zilong et al., 2008; 2011; Yang Shufeng et al., 2007); 284~270Ma 期间以巴楚地区广泛分布的辉绿岩为代表的浅成侵入活动为主,同时伴有同时期的酸性火山活动(Yang Shufeng et al., 1996; Chen Hanlin et al., 1997; Liu Dongdong et al., 2013; Tian Wei et al., 2010),如塔北高钾钙碱性长英质

火山岩(~276Ma)(Yu Junchuan et al., 2011)。另外,在东天山地区(包括星星峡断裂),同样出露一系列后碰撞期石炭—二叠纪的 300~270Ma 偏铝质和钙碱性高钾火山岩(Ma Xuxuan et al., 2015)。而在罗布泊周边地块,只有北山黑山岭地区晚石炭世火山岩报道(Li Yuhang et al., 2008)以及以橄榄岩、辉长岩为主的 289~260Ma 镁铁—超镁铁质杂岩体(Tang Qingyan et al., 2015)。

盆内印支期多为陆内裂谷型火山活动,板块消减及碰撞后造山阶段有关的岩浆作用可能一直持续到三叠纪(Li Shan et al., 2010)。报道见北山地区 210Ma~250Ma 岩浆岩成岩和成矿年龄(Jiang Sihong et al., 2006; Miao Laicheng, 2014),东天山鄯善、尾亚等地钾长花岗岩(Li Wenming et al., 2002)及中天山东段天湖花岗岩(Zhao Honggang et al., 2017)等同时段岩体年龄。侏罗纪年龄缺失,中白垩世—渐新世一共获得 6 个年龄(119Ma, 103Ma, 33.4~46.1Ma),可能指示了托云盆地新生代火山岩(Ji Jianqing et al., 2006; Liang Tao et al., 2007)和皮羌火山岩(Luo Zhaozhua et al., 2003)等对罗布泊盆地的贡献。

本次获得的锆石 U-Pb 年龄集中在 209~240Ma、265~304Ma、320~385Ma、406~446Ma、705~880Ma 区间,这与碎屑锆石限定的区域上发生的大部分岩浆活动时间相吻合,强烈的构造运动和风化作用导致岩体被剥蚀后带入盆地。综合分析研究区及邻区已有同位素年代学资料,表明罗布泊地区前寒武纪年龄可能主要来自于近源的库鲁克塔格地区,加里东期碎屑锆石可能来自于盆地南北两个方向的供给即阿尔金造山带和南天山构造带或其一。276Ma 的峰值记录了塔里木盆地大火成岩省二叠纪酸性岩浆底侵作用时间,东天山、北山、南天山西段均有同时段岩体产出年龄报道,印支期年龄锆石可能来自于北山和东天山。白垩纪—古近纪碎屑锆石年龄可能指示较远的帕米尔—西昆仑山等地也提供了少量物源。

5.2 罗布泊凹陷演化过程

罗布泊更新统碎屑锆石数据显示中生代以来缺失侏罗纪年龄,白垩纪只有 2 个年龄产出(119Ma 和 103Ma),推测可能有两个原因:①中、新生代以来,塔里木盆地进入陆内前陆盆地—挤压造山阶段(Li Yuejun et al., 2009)。盆内燕山期岩浆活动鲜有报道,罗布泊周缘地区仅见东天山白山铼钼矿区 181Ma 的黑云母斜长花岗岩体(Li Huaqin et

al., 2005), 指示侏罗纪以来塔里木盆地花岗岩分布规模可能十分有限。②古近纪之前, 塔里木盆地呈现东高西低的古地理面貌 (Mu Guijin et al., 2001), 这个时期的罗布泊连同其周缘地区库鲁克塔格处于抬升区内 (Wu Guanghui et al., 2007), 一同作为高地经历剥蚀—准平原化, 这也可能是侏罗纪年龄缺失的原因。始新世时期由于受到印度板块对欧亚板块的俯冲与碰撞后远程效应的影响, 天山和昆仑山地区经历了广泛快速的抬升 (Tang Liangjie, 1996; Li Zhong et al., 2010; Yang Wei et al., 2013), 东部地区开始相对沉降。受到这种构造反转的约束 (Liu Chenglin et al., 2008), 罗布泊开始演变为塔里木盆地东端的最低洼处, 完成了沉积中心自西向东的转移, 成为塔里木盆地的汇水区, 最终接受了大量的碎屑沉积。

从罗布泊 LDK01 孔碎屑锆石年龄分布纵向演化来看 (图 3), 早更新世早期的样品 N215 主峰值约为 276Ma, 对应南天山造山带后碰撞高钾花岗质岩浆活动 (Wang Chao et al., 2007), 说明大量该时段的物源通过塔里木河流带入到罗布泊; 早更新世晚期 N221 主峰为 236Ma, 与东天山罗觉塔格地区强烈的印支期花岗岩侵入活动时间吻合 (Li Huaqin et al., 2006), 除此之外还观察到 2405Ma 峰值为库鲁克塔格结晶基底年龄。这说明早更新世时期罗布泊正处于断陷湖盆演化阶段, 近源山体剥蚀物质和河流带来的碎屑组分在罗布泊快速堆积下来。中更新世样品 N133 特征峰值为 266Ma, 指示主要物源可能同样来自南天山造山带。但碎屑锆石年龄表现为各个时代分布十分广泛, 古元古代年龄相对较为集中, 2376Ma 峰值明显。推测造成这种现象的原因是, 中更新世时期一次新构造运动可能使罗布泊古湖盆基底发生不均衡抬升, 罗北凹地开始形成 (Liu Chenglin et al., 1999)。该次构造活动剧烈, 周缘山系隆起幅度增大, 切割基岩程度加剧, 带动卷入库鲁克塔格前寒武变质基底年龄。除此之外, 早—中更新世样品碎屑锆石颗粒普遍大于晚更新世样品 (N80 和 N2) (图 4), 这也可能说明了盆地早期处于断陷阶段, 构造运动较为强烈, 经较强的水动力条件搬运来的粗颗粒锆石数量较多。之后亚洲内陆加速干旱化进程 (Li Jijun et al., 1999), 罗布泊地区随着气候极端干旱化与构造活动性减弱, 湖盆演化也进入充填萎缩阶段, 周缘物质输入趋于稳定。

5.3 物源供给方式初探

分析物源的传统做法是将沉积岩的碎屑锆石年

龄谱与周缘造山带结晶岩体 (包括岩浆岩和变质岩) 的年龄进行对比, 若年龄匹配, 则该造山带可解释成潜在物源之一 (Vermeesch, 2012)。但实际上分析过程必须考虑多种因素, 因为同一个造山带可能经历多期岩浆活动, 沉积物再旋回造成与年轻岩体的混合, 尤其是来自不同物源区的河流系统在沉积区汇合 (Thomas, 2011)。

罗布泊在第四纪早期曾是一个统一的大湖区, 面积覆盖了现今的罗北凹地、东西台地、大耳朵湖盆地 (Wang Mili et al., 2001), 北连库鲁克塔格山, 南到阿尔金山, 东至阿奇克谷地, 发源于昆仑山的和田河、克里雅河由南往北注入塔里木河, 盆地西部的喀什噶尔河、叶尔羌河、阿克苏河向东并入塔里木河, 之后与发源于阿尔金山西段的车尔臣河在台特玛湖汇合后, 与发源于博斯腾湖的孔雀河最终汇聚于罗布泊凹陷 (Bo Ying et al., 2013; Liu Chenglin et al., 2015, 2016; Mischke et al., 2017), 因此历史上处于泛湖期的罗布泊古湖面积很大, 接收了塔里木流域内多条河流的碎屑物质输入。

为验证这一假设, 选取了 4 个潜在物源区与罗布泊 LDK01 孔碎屑锆石 U-Pb 年龄组成进行概率密度估计曲线对比 (图 8a), 其中墩阔坦乡位于现代塔里木河和孔雀河流域, 其全新世河道砂样品代表了现代河流信息。这五组样品的 2 个主要年龄峰基本重合, 介于 250Ma~320Ma 和 410~460Ma 之间, 代表了这些地区对罗布泊物源均有一定程度的贡献。但累计年龄分布曲线 (图 8b) 更明显地展示了碎屑锆石年龄之间的微弱差距 (Vermeesch, 2012)。为了考证西昆仑山前物源对东端罗布泊盆地的贡献, 挑选了麻扎塔格山上新世砾岩年龄谱 (Si Jialiang et al., 2009) 与罗布泊样品进行比较。结果显示, 两者相距最远, 麻扎塔格山主峰最年轻碎屑锆石年龄为 ~250Ma, 可能是由于其上新世物源主要来自于西昆仑西段及北缘晚古生代—中生代印支期花岗质岩浆作用的产物 (Kang Lei et al., 2015; Huang Jianguo et al., 2016), 并非罗布泊的主要供给物源。准噶尔盆地南缘塔西河剖面早更新世西域组砾岩的碎屑锆石主峰年龄 ~306Ma, 与 LDK01 主峰 ~276Ma 之间也存在一定的差异, 这是由于天山自中生代以来就持续为两侧沉降盆地提供物源 (Hendrix, 2000), 北天山主要出露泥盆纪—石炭纪地层 (Li Jinyi et al., 2006), 所以“就地取材”成为准噶尔盆地南缘主要物源区 (Yang Wei et al., 2013)。而罗布泊 LDK01 孔与南天山西段黑英山

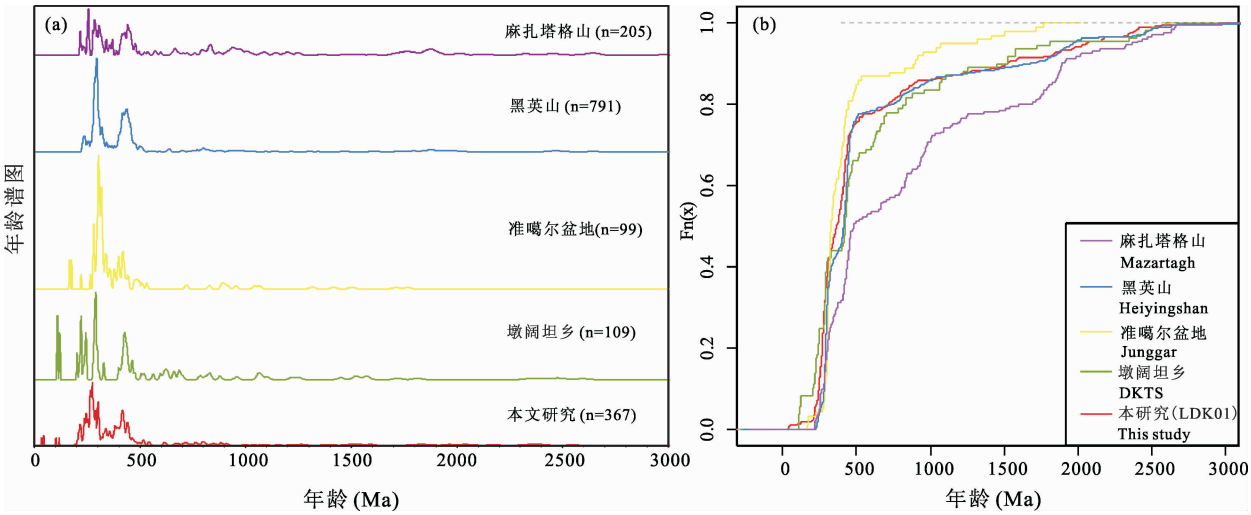


图 8 新疆罗布泊 LDK01 孔碎屑锆石年龄与其他地区对比图

Fig. 8 Comparison between the zircon ages of Lop Nur sediments and other areas in Xinjiang

(a)—KDE 图;(b)—CAD 图 数据来源:麻扎塔格山上新世砾岩(Si Jialiang et al. , 2009);南天山西段黑英山—卡普沙良二叠纪—侏罗纪沉积岩(Liu Dongdong et al. , 2013);准噶尔盆地南缘北天山北麓西域组砾岩(Yang Wei et al. , 2013);DKTS(墩阔坦乡)全新世河道砂(未发表数据)

(a)—Kernel density estimation plot; (b)—cumulative age distributions plot. Data from: A Lower Pliocene conglomerate layer of Mazartagh (Si Jialiang et al. , 2009); the Permian-Jurassic sediments of the western South Tianshan (Liu Dongdong et al. , 2013); the lower Pleistocene Xiyu formation in the southern Junggar basin(Yang Wei et al. , 2013); DKTS (Dunketangxiang) Holocene river sand (unpublished data)

一卡普沙良剖面基岩年龄(Liu Dongdong et al. , 2013)和墩阔坦乡河道砂碎屑锆石年龄最为接近。南天山样品表现为 300~260Ma 的年龄峰,是中—南天山岩浆岩地质体的特征年龄(Li Zilong et al. , 2011),而 450~410Ma 加里东期年龄峰在天山和阿尔金山均有大量出露,如起源于中天山的开都河两岸的早泥盆世中酸性火山岩(Zhu Zhixin et al. , 2009)。墩阔坦乡河道砂碎屑锆石虽表现为年龄构成稍宽,但与其下端的罗布泊物源表现高度一致,这同样说明了该河流对罗布泊的主要贡献。

总的来说,研究区周缘北部的库鲁克塔格山可能主要提供前寒武纪的年龄和南部的阿尔金山主要提供加里东期的碎屑锆石年龄,虽然罗布泊东部北山地块也存在大量的 289~260Ma 镁铁—超镁铁质杂岩体(Tang Qingyan et al. , 2015),但是处于伸展背景下产出的铁镁质岩体不可能产生大量锆石(Moecher et al. , 2006),物源供给方向来自于东部的可能性也较小。所以这一分析结果最有可能说明,罗布泊的主要物质来源是来自塔里木盆地北缘南天山山前的塔里木河,这一点从样品早二叠世的碎屑锆石一定磨蚀的形态也可以证明;而盆地北部、东北部造山带为罗布泊提供的物源十分有限。需要说明的是,现今看到的塔里木河道是风沙和河流改造共同作用的结果,具体哪条河流率先发育并主导

了第四纪时期罗布泊的物源供给还需要进一步的研究,建立塔里木河流域上、下游的大数据库,这不仅对第四纪地貌反演以及对罗布泊第四纪环境变化甚至楼兰古国的消亡都有重要意义。

6 结论

(1)罗布泊更新统碎屑锆石类型以岩浆锆石为主。年龄分布比较广泛,物源年龄构成复杂,主要集中在 209~240Ma、265~304Ma、320~385Ma、406~446Ma、705~880Ma 五组。这些年龄与塔里木盆地几期重要的构造运动有关,可分为古元古代增生造山事件,Rodinia 大陆裂解期,加里东中期,海西中晚期及燕山期,其中海西期岩浆活动最为强烈,碎屑锆石年龄最为集中。综合分析研究区及邻区已有同位素年代学资料,认为罗布泊地区前寒武纪碎屑锆石年龄可能主要来自库鲁克塔格,加里东期碎屑锆石主要来自阿尔金造山带和(或)南天山构造带;276Ma 的峰值指示主导物源可能来自南天山;印支期年龄锆石可能来自于北山和东天山;新生代碎屑锆石可能有来自于较远的帕米尔—西昆仑山等地的物源。

(2)碎屑锆石年龄分布纵向演化特征显示,早更新世时期罗布泊正处于断陷湖盆演化阶段,近源山体剥蚀物质和河流带来的碎屑组分在罗布泊快速堆

积下来,是造成年龄分布宽泛的原因;中更新世罗布泊凹陷北部山前经历过一次明显的构造抬升,可能是导致罗北凹地形成的重要原因。锆石形态同样揭示了进入晚更新世以来构造活动强度减弱的过程。

(3)第四纪以来罗布泊成为塔里木盆地最低洼处汇水盆地,其碎屑锆石组成特征以及潜在源区对比表明,塔里木盆地北缘南天山山前的塔里木河对罗布泊物源贡献较大,而罗布泊北部山前尤其是东北部对凹陷物源供给十分有限。

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Provenance of the Quaternary Lake Basin and Tectonic Evolution of the Basin
in Lop Nur: Evidence from Detrital Zircon U-Pb Age of Core LDK01

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Abstract

Detrital zircon U-Pb age analyses for the samples from the first scientific potash exploration drill core in Lop Nur were conducted to reveal the potential source area of the Quaternary salt lake and stage tectonic evolution of the depression in Lop Nur. Th/U ratios show that the detrital zircons are dominantly magmatic origin, with minor metamorphic zircons. Zircon ages are concentrated mainly in the ranges of 209~240Ma, 265~304Ma, 320~385Ma, 406~446Ma, 705~880Ma and 2376~2405Ma. Combined with the rock nature and age range of the potential source area, it can be preliminarily concluded that the pre-Cambrian materials mainly sourced from the Kuruktag area, and the Caledonian sources derived from the Altyn and South Tianshan tectonic belts. 276Ma of age peak records a representative event of Permian large igneous province (LIPs) in the Tarim basin and South Tianshan is likely the major provenance. Indosinian and Cenozoic detrital zircon ages imply that the Beishan block and East Tianshan, even distant Pamir and West Kunlun, were the potential source areas. Due to large-scale weathering and leaching, widely exposed K-rich rocks in the periphery of the depression were transported to Lop Nur, providing important materials for the Quaternary potash formation in Lop Nur. In addition, sediment compositions transported by drainage system of the Tarim basin were also major provenance, with limited amount from the orogenic belts in northern and northeastern Lop Nur. The vertical variety features of detrital zircon ages suggest an obvious tectonic uplift in the north part of Lop Nur, probably resulting in the formation of the Luobei depression.

Key words: Lop Nur; provenance; tectonic evolution; detrital zircon