

西藏羌塘增生杂岩带内中一晚三叠世碳酸盐岩地层的发现及大地构造意义

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内容提要:羌塘中央增生(变质)杂岩带地层的解体有助于厘定龙木错—双湖缝合带和古特提斯洋俯冲—碰撞的时空演化过程。长期以来,该缝合带三叠纪地层研究薄弱,已有认识缺乏与洋陆俯冲古地理格局的联系。蓝岭高压变质带北侧 3 个露头中新发现一套中一晚三叠世台地相碳酸盐岩夹细碎屑岩、硅质岩沉积组合,它们以断块状叠置于中二叠世蛇绿混杂岩之上,内部层理近水平状,层序并未倒转,但顶底不全。依据岩性组合、相对层厚和化石分布,将层序较完整的西南侧岩块大致划分出三个沉积亚相:局限台地相、开阔台地相和台地边缘礁相。边缘礁相砾屑、生屑灰岩中含丰富的四射珊瑚化石,已鉴定出的珊瑚总计 6 属 1 种,分别为 *Conophyllia* sp., *Omphalophyllia* sp., *Coryphyllia* sp., *Pamirastrea* sp., *Margarophyllia* cf. *capitata*, 和 *Craspedophyllia* sp., 指示地层时代为中三叠世拉丁期—晚三叠世诺利期。通过对羌塘盆地中—上三叠统的对比,发现北羌塘地体、中央增生杂岩带和南羌塘地体分别发育不同的沉积体系。构造接触关系及地层特征综合显示蓝岭北侧的中上三叠统代表北羌塘活动大陆边缘的弧前沉积,具体类型为增生楔上叠盆地。结合区域内中上三叠统的产出背景和生物化石时代,推测古特提斯洋闭合于晚三叠世诺利期中期。

关键词:羌塘;三叠纪;碳酸盐岩;珊瑚;古特提斯洋

青藏高原特提斯洋的演化具有明显的节律性,原、古、中、新特提斯洋的形成发展与南半球冈瓦纳大陆的间歇性裂解密切相关,裂离的陆块以“传送带”的形式向北运移并碰撞增生到欧亚大陆的南缘(Metcalf, 2006; Pan Guitang et al., 2012a; Pan Guitang et al., 2012b)。羌塘盆地中部的“龙木错—双湖缝合带”(图 1)为古特提斯洋俯冲消减的产物,分隔了亲冈瓦纳的羌南地块与亲扬子/华夏的羌北地块(Li Cai, 1987; Li Cai et al., 1995, 2008, 2009),不同阶段特提斯洋(∈-T)的地质演化信息均有保存(Li Cai et al., 2009; Pan Guitang et al., 2012a; Wu Yanwang et al., 2013; Zhang Xiuzheng et al., 2014; Hu Peiyuan et al., 2014, 2015; Zheng Yilong et al., 2015)。部分学者认为该缝合带与滇西昌宁—孟连缝合带相链接,继续延伸后等同于泰国—马来西亚的清莱—本通带,共同构成了早古生代—三叠纪冈瓦纳大陆与劳亚大陆的主边界—古特提斯洋主洋盆(Li Cai et al., 1995; Liu Benpei et

al., 2002; Pan Guitang et al., 2012a; Che Zicheng, 2016)。沿缝合带近 E-W 向展布着长逾 500 km, N-S 向宽近 100 km 的变质杂岩(Yin An et al., 2000; Kapp et al., 2003; Wang Genhou et al., 2009; Liang Xiao et al., 2012)。Wang Genhou et al., (2009)将羌塘中央变质杂岩带厘定为印支期古特提斯洋俯冲消减形成的增生杂岩(时代上限为晚三叠世),多期面理置换后的复理石碎屑岩作为基质包裹了基性岩、超基性岩、高压变质岩等岩块(时代下限可追溯到奥陶纪)。洋壳单向向北俯冲于北羌塘地块之下(Zhai Qingguo et al., 2011a; Liang Xiao et al., 2012),北羌塘南缘和增生杂岩带北部分布有 NWW-SEE 向,长逾 200 km 的晚三叠世诺利期—瑞替期中酸性火山岩—花岗岩系(Kapp et al., 2003; Hu Peiyuan et al., 2010; Fu Xiugen et al., 2010a; Zhang Kaijun et al., 2011; Zhai Qinghuo et al., 2013),年龄分布在 222~209 Ma 的区间内,岩石地球化学指示同碰撞背景(Zhang Xiuzheng et

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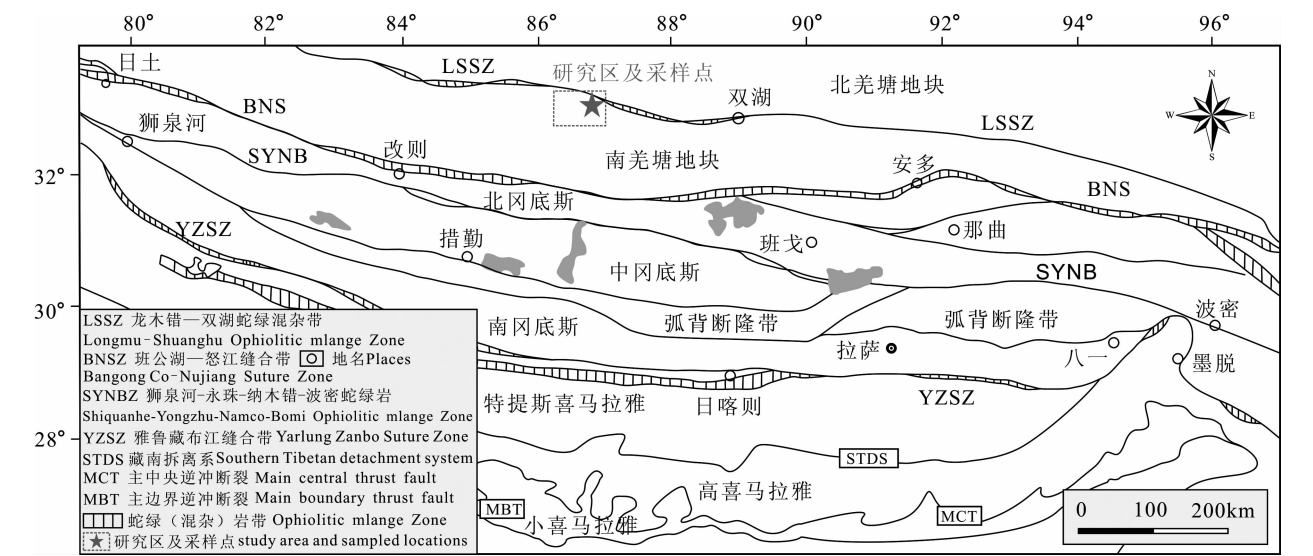


图 1 西藏大地构造分区及研究区位置(修改自 Pan Guitang et al. ,2006;Zhu Dicheng et al. ,2006)

Fig. 1 Sketched tectonic framework of Tibetan Plateau and the location of the study area
(modified from Pan Guitang et al. ,2006;Zhu Dicheng et al. ,2006)

al. ,2014;Li Guangming et al. ,2015;Li Jingchao et al. ,2015)。果干加年山上三叠统望湖岭组细碎屑岩角度不整合在增生杂岩之上,底部流纹岩获得了锆石 U-Pb 年龄 214 ± 4 Ma,代表南北羌塘地块碰撞造山的开始(Li Cai et al. ,2007)。

片石榴辉岩变质锆石 U-Pb 年龄 (237 ± 4 Ma、 230 ± 3 Ma)(Zhai Qingguo et al. ,2011b)及石榴子石 233 ± 13 Ma 的 Lu-Hf 等时线年龄(Pullen et al. ,2008)约束羌塘中部高压变质作用和古特提斯洋向北俯冲事件至少可以追溯到中三叠世。龙木错—双湖缝合带最新发现的早石炭世望果山组弧火山岩组合(Jiang Qingyuan et al. ,2014;Xu Wei et al. ,2014)指示洋壳俯冲很可能始于石炭纪。中—晚三叠世浅海—深海相沉积保留在羌塘西部的才玛尔错(Xiong Xingguo et al. ,2006),中部的荣玛乡北雪水河(Li Yuejun et al. ,1997),以及中北部的果干加年山(Ji Zhansheng et al. ,2010),显示洋壳闭合发生在晚三叠世以后。然而,中—晚三叠世沉积的构造背景一直缺乏探讨,已有认识也未能厘定出古特提斯洋俯冲相关的弧前沉积体系(Li Cai et al. ,2008,2009;Zhao Zhongbao et al. ,2014),因此羌塘中部三叠纪活动大陆边缘古地理格局仍然缺乏有效约束。

笔者在蓝岭北侧新发现一套中三叠世拉丁期—晚三叠世诺利期浅海台地相碳酸盐岩地层,含丰富的特提斯温暖浅水型珊瑚化石组合,构造背景为增生楔上叠型海沟—斜坡盆地。对该套地层的研究有助于古特提斯洋闭合时限的确定,以及对塑造三叠

纪洋陆俯冲的时空格局具有重要意义。

1 地质背景

羌塘盆地面积近 18 万 km^2 (图 1),大面积分布着三叠纪—侏罗纪的滨浅海相沉积,是我国最大的中生代残留盆地(Wang Chengshan et al. ,2001;Zhao Zhengzhang et al. ,2001;Wang Jian et al. ,2004)。拉萨—羌塘地块的地球物理探测普遍显示羌塘盆地地壳具有“两坳夹一隆”的构造格局(Qian Hui et al. ,2016),沿盆地中央的冈玛错—戈木错/果干加年山—玛依岗日—角木日—双湖为“羌塘中央隆起带”,并由此间隔了北羌塘与南羌塘两个次级凹陷(Huang Jijun et al. ,2001;Lu Bing et al. ,2001;Zhao Zhengzhang et al. ,2001)。这一认识最近得到了沿拉萨地块北缘—羌塘盆地北缘反射地震剖面的有力支持(Lu Zhanwu et al. ,2013;Gao Rui et al. ,2013)。中央隆起带主要由早古生代—三叠纪地层混杂而成的洋壳俯冲—增生杂岩构成,(Wang Genhou et al. ,2009;Liang Xiao et al. ,2012)。古特提斯洋(早古生代—三叠纪)的洋壳残片断续状保留在增生杂岩带内部^{①②},(Zhu Tongxin et al. ,2006;Li Cai et al. ,2008,2009;Wu Yanwang et al. ,2010;Zhai Qingguo et al. ,2013)。

增生杂岩主体岩性为石炭—二叠纪被动大陆边缘碎屑岩+碳酸盐岩^③,并以大陆斜坡浊积岩或海底扇为代表^④,石炭纪地层中的“含砾板岩”(冰海杂砾岩)指示冈瓦纳冰海沉积环境^⑤(Liang Dingyi et

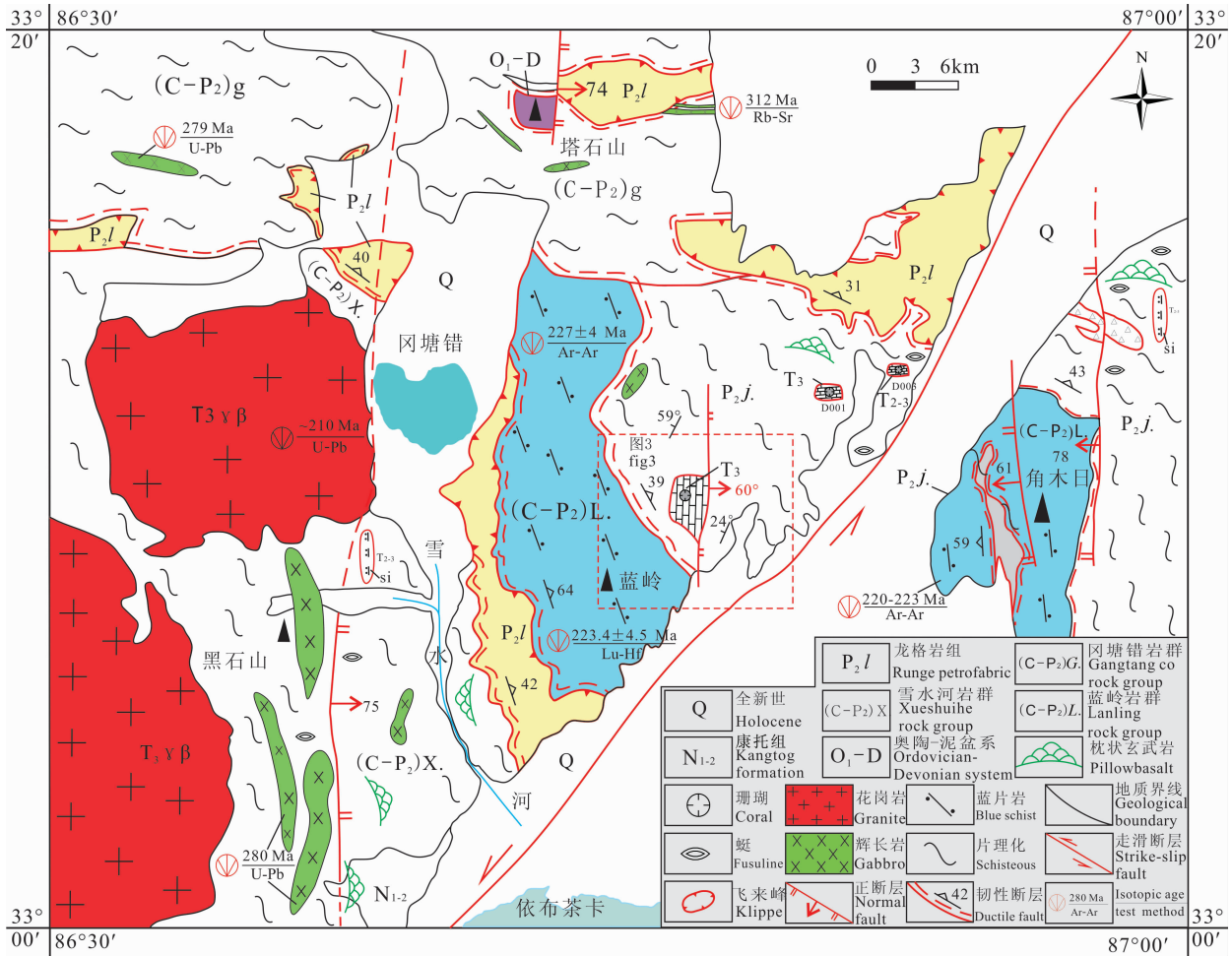


图2 西藏羌塘中部角木日地区地质—构造简图及中上三叠统的出露位置^①
Fig. 2 Simplified structure-geological map of Jiaomuri area in Central Qiangtang of Tibet and the location of the Middle-Late Triassic

al., 1983, 1994; Yin Jixiang 1997; Li Cai et al., 2008, 2009; Fan Jianjun et al., 2012, 2015), 早中二叠世浅海相灰岩层含丰富的蠕类化石组合并且具有冷暖水混生的面貌^{②③} (Geng Quanru et al., 2012; Jiao Pengwei et al., 2017)。Fan Jianjun et al. (2016, 2017) 在日湾茶卡新识别出晚二叠世—中三叠世天泉山洋岛型安山岩、玄武岩组合, 锆石 U-Pb 谐和年龄为 242~254Ma, 指示羌塘中部古特提斯大洋持续演化到中三叠世。中上三叠统呈星点状分布在增生杂岩带内, 主要为半深水、深水相灰岩和硅质岩建造 (Li Yuejun et al., 1997; Xiong Xingguo et al., 2006; Ji Zhansheng et al., 2010)。荣玛乡以南, 增生杂岩被上三叠统日干配错组及侏罗纪海相沉积以角度不整合所覆盖^{④⑤} (Chen Wenxi et al., 2009)。侏罗纪, 羌塘“中央隆起”仍然作为显著的古地理界线, 两侧的南北羌塘海相沉积存在明显区别^⑥ (Wang Chengshan et al., 2001; Wang Jian et

al., 2004; Chen Wenxi et al., 2009)。白垩纪以后, 羌塘盆地整体上逐渐转变为陆相沉积盆地 (Wang Chengshan et al., 2001; Wang Jian et al., 2004)。

蓝岭高压变质带西临果干加年山一片石山含榴辉岩高压变质带, 总体呈 N-S 向展布, 东西两侧均与二叠纪—三叠纪蛇绿混杂岩相拼贴 (图 1、2)。东侧新发现的中—上三叠统的三个露头中, 岩层均以“断块”形式与中二叠世海山岩块 (枕状玄武岩/浅水灰岩) 相接触, 主体岩性为一套灰—深灰色碳酸盐岩 (图 2), 夹有浅灰色硅质岩薄层, 部分与灰绿色粉砂质泥岩呈互层状。其中, 西南端的断块出露规模较大, 东侧被向东陡倾的正断层所切割 (图 3), 内部层理向西缓倾, 层序并未倒转, 但未见顶底出露。

2 地层特征

2.1 实测剖面简介及地层特征

对地层层序较为连续的西南侧岩块进行了实测

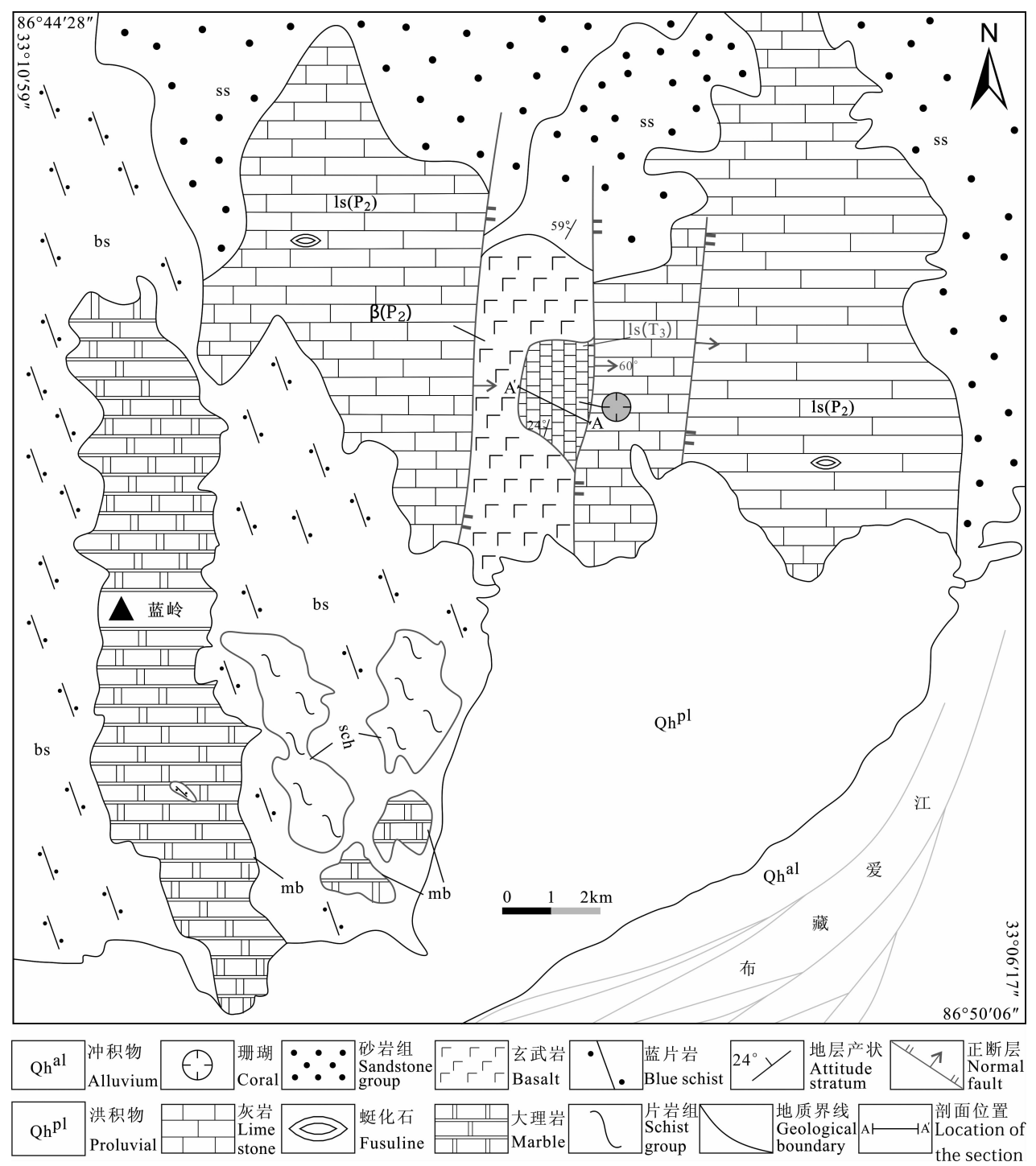


图 3 西藏羌塘蓝岭高压变质带及北侧蛇绿混杂岩地质简图及实测剖面位置

Fig. 3 Simplified geological map of Lanling in central Qiangtang of Tibet high-pressure metamorphic belt and ophiolitic melange on its northern flank

剖面。地理位置:西藏自治区尼玛县荣玛乡蓝岭北侧,坐标: N33° 08′ 44. 03″, E86° 47′ 20. 04″, 海拔 4966m。地层呈断块状产出于浅变质混杂岩之中,出露面积较小,近东西向延伸约 100m,出露厚度大于 45m。顶底与中二叠世海山枕状玄武岩+灰岩建造呈断层接触,东侧被新生代高角度正断层所切割,

总体产状平缓(305∠24°)。该剖面岩石组合主要为碳酸盐岩夹碎屑岩和少量硅质岩条带,剖面(图 4)具体描述如下:

二叠系枕状玄武岩

=====断 层=====

8. 浅灰—灰色薄层泥晶灰岩,偶夹泥质条带。单层厚 5~

- 10cm。灰岩含双壳类、海百合茎、腕足类碎片。3m
7. 灰—深灰色巨厚层含生屑、内碎屑泥晶灰岩,生屑为腕足类、双壳类、菊石、苔藓虫,采集含单体和复体珊瑚,计有 3 属,*Conophyllia* sp.,*Omphalopyllia* sp.,*Coryphyllia* sp. 7m
6. 深灰色中厚层泥晶灰岩与灰绿色粉砂质泥岩互层,局部夹有 1~4cm 硅质条带。自下而上泥岩夹层明显增大。20m
5. 深灰色厚层生屑、内碎屑亮晶灰岩,生屑破碎明显,见有菊石、双壳类、苔藓虫、海百合茎,内碎屑以砂屑为主,砂屑少量,均具有一定的磨圆。该层发现复体珊瑚 1 件,*Pamirastraea* sp. 2m
4. 灰—深灰色中薄层泥晶灰岩与浅灰色粉砂质泥岩互层,灰岩单层厚 3~16cm,粉砂质泥岩 1~4cm。4m
3. 灰—深灰色薄层白云质灰岩,变质重结晶呈微晶状,厚约 20cm。3m
2. 灰—深灰色中厚层泥晶灰岩与浅灰—灰绿色薄层粉砂质泥岩互层,灰岩厚 16~20cm,略微变质重结晶,粉砂质泥岩厚 1~5cm,发育顺层劈理。2m
1. 灰—深灰色薄层泥晶灰岩夹黑色硅质条带或团块,灰岩厚 2~6cm,略变质重结晶,硅质岩 2~4cm,石香肠化。5m
- =====断 层=====

二叠系枕状玄武岩

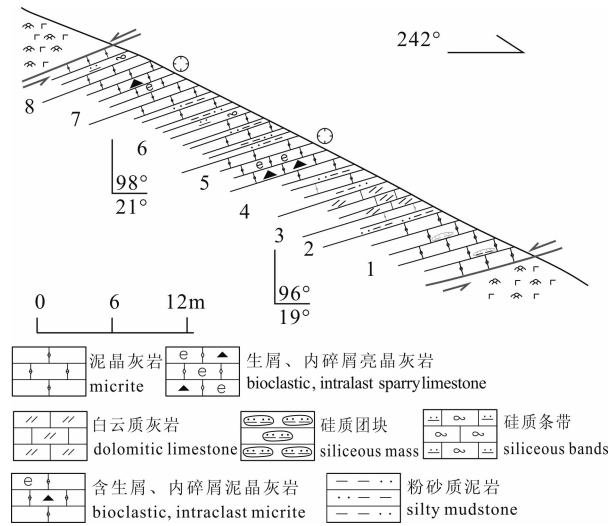


图 4 西藏羌塘蓝岭上三叠统碳酸盐岩实测剖面
Fig. 4 Measured section of carbonate rock the Upper Triassic of Lanling in central Qiangtang of Tibet

实测剖面中观察到岩层发生了浅变质变形作用,大部分灰岩变质重结晶后呈微晶状,硅质岩多被压扁拉长后呈对称的石香肠构造(图 5c),第 7 层所含的珊瑚化石内部结构被不同程度破坏,因此不能鉴定到种。泥岩中发育一组平行于原生层理 S_0 的板劈理 S_1 ,第 2 层与第 6 层的灰岩中发育一组斜交层面 S_0 的板劈理,依据“劈理比层理陡”的关系,判断露头中地层层序并未倒转(图 5f)。由底到顶,岩

石组合大致可以分为两个旋回,第 1~4 层构成了第一个岩性旋回(图 5b、d)。第 1~4 层为韵律状层序,互层状的灰—深灰色中薄层—薄层状泥晶灰岩和薄层状粉砂质泥岩与中薄层—薄层状泥晶灰岩重复出现(图 5b),部分泥晶灰岩层间含有浅灰色薄层状的硅质岩(图 5b)。第 5 层为厚层状生屑、内碎屑灰岩,碎屑粒度普遍在 2mm 以上,最大可达 2cm,大小混杂,形状十分不规则,大部分为棱角状(图 5d)。第 6~7 层构成了另一个岩性旋回(图 5e、f),其中第 6 层为灰—深灰色薄层—中薄层状泥晶灰岩灰绿色粉砂质泥岩互层(图 5f),自下而上粉砂质泥岩的层厚由薄层增大为中薄层,最大可达 16cm。第 7 层岩性与第 5 层类似,但层厚明显增加,碎屑的直径增大,最大可达 5cm,并且生物碎屑含量更高(图 6l),例如,大量分布有单体或复体珊瑚化石。第 8 层岩性主要为浅灰—灰色薄层泥晶灰岩,灰岩含有生物碎片。

2.2 生物组合及时代

剖面的第 5 层和第 7 层含有丰富的珊瑚化石,共采集 28 件。另外两件化石样品 D001-HB1(图 6a、b)和 D003-HB1(图 6c)是距离剖面位置北东方向约 40°,约 10km 处采集(图 2)。化石薄片由北京大学教育部重点实验室贾秋月老师制作,鉴定由中国地质大学(北京)地球科学与资源学院王训练老师完成。共 12 件薄片含有珊瑚化石,但仅有 6 块化石薄片具有鉴定意义,其中 1529-HB13、1529-HB26 可以见到珊瑚的完整的横切面(图 7)。可鉴定珊瑚化石分别为 *Conophyllia* sp.,*Omphalopyllia* sp.,*Coryphyllia* sp.,*Pamirastraea* sp.,*Margarophyllia* cf. *capitata*,*Craspedophyllia* sp.,从化石的组合特点及国内外主要产出层位看(表 1),该珊瑚化石组合的地层时代应为中—晚三叠世,生活环境为正常温暖浅海。

3 沉积相分析

依据其主体岩性,可以判断中—上三叠统沉积环境为浅海台地相。现以西南侧地层连续的岩块为例,对其沉积环境进行分析。根据实测剖面中岩性组合(图 6d~6i、6l、m)、镜下薄片特征、生物化石组合特征,由下而上可将其沉积相大致分为三个亚相:局限台地、半局限—半开阔台地、开阔台地和台地边缘礁。

第 1、3 层灰—深灰色中薄层状泥晶灰岩普遍夹有浅灰色的硅质岩薄层,生物碎屑或化石含量少(图

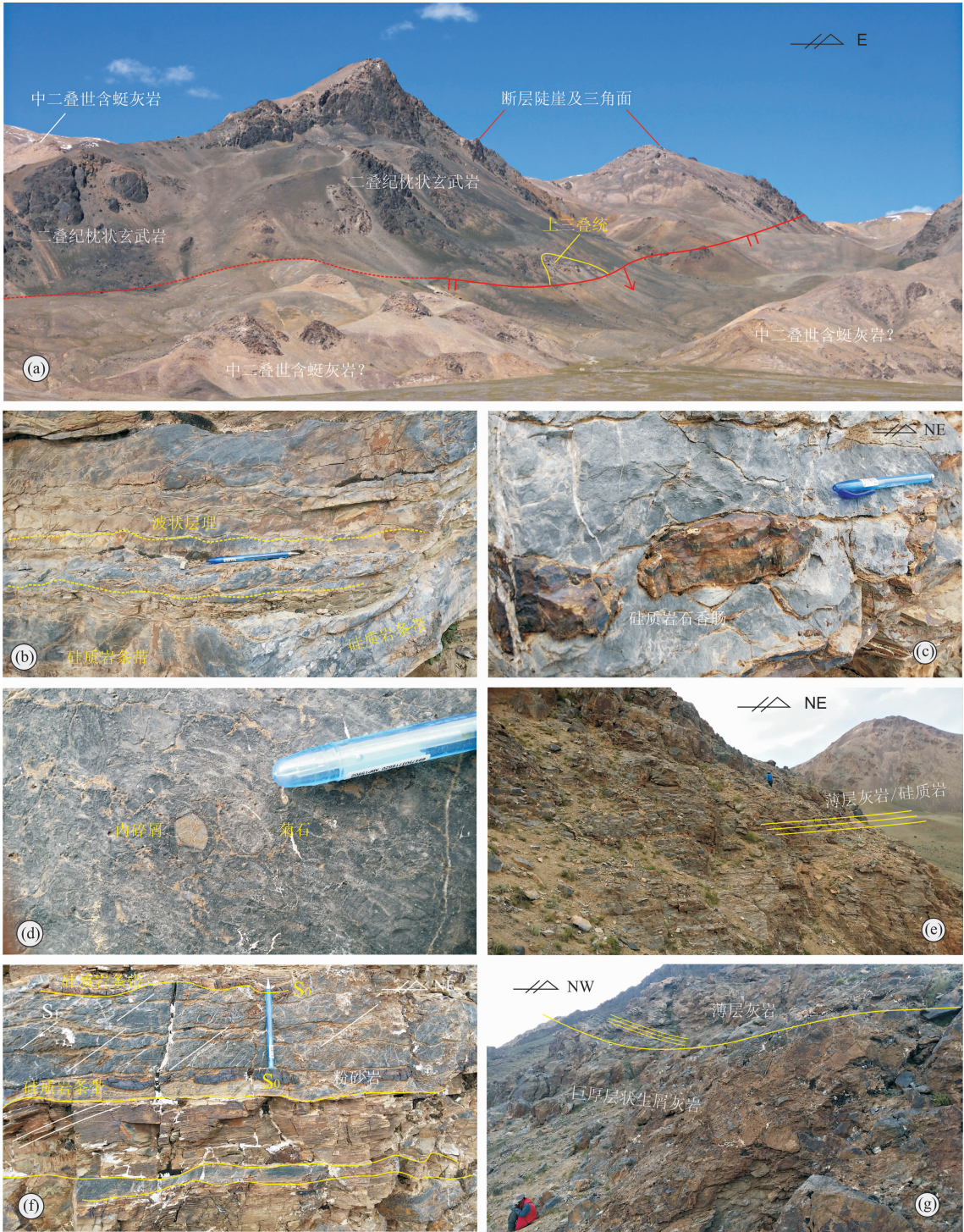


图 5 西藏羌塘蓝岭晚三叠世碳酸盐岩野外露头照片

Fig. 5 Outcrop and stratigraphical characteristics of the Late Triassic carbonatite of Lanling in central Qiangtang, Tibet
(a)—上三叠统宏观露头照片;(b)—剖面第 1~3 层岩性及波状层理;(c)—第 3 层泥晶灰岩中的硅质岩石香肠;
(d)—第 5 层生屑、内碎屑灰岩;(e)—第 6 层宏观照片;(f)—第 6 层中薄层状粉砂质泥岩与泥晶灰岩互层;(g)—第 7、8 层岩性组合特征
(a)—Macroscopic outcrop characteristic of the upper Triassic which demonstrates a clear contacting relationship between it and the surrounding Middle Permian seamount pillow basalt/carbonatite;(b)—lithology of the first-third layers in section which show a current bedding;(c)—silicite boudinages within the third layer of micrite;(d)—the fifth layer of bioclastic/intraclast limestone;(e)—outcrop characteristics of the sixth layer;(f)—the interbedding of thin silty mudstone and mud limestone of the sixth layers;(g)—lithological characteristics of seventh-eighth layers

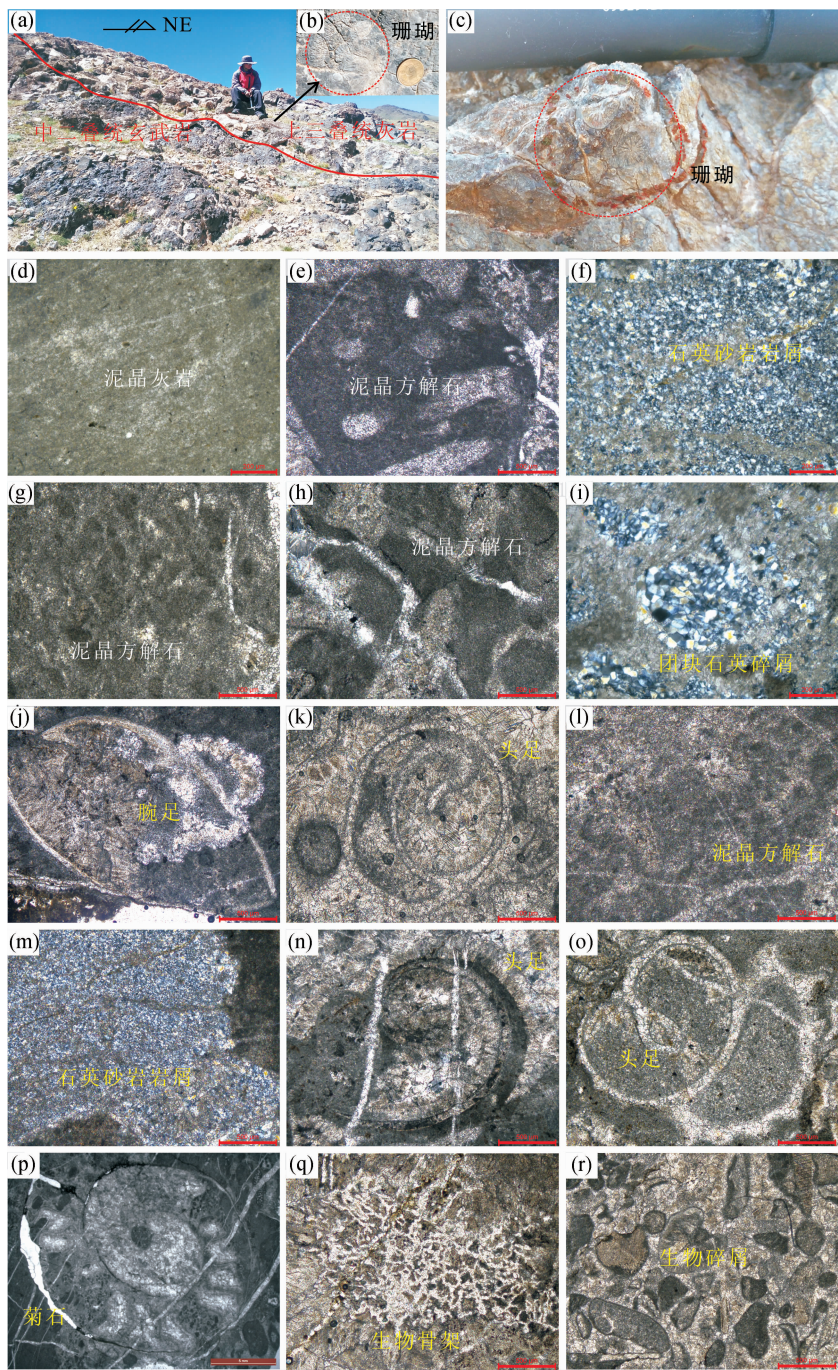


图 6 西藏羌塘蓝岭碳酸盐岩野外和显微照片

Fig. 6 Pictures of in the filed and micrograph the carbonatite rocks of Lanling in central Qiangtang of Tibet

(a)、(b)—D001-HB1 野外宏观和珊瑚近景照片；(c)—D003-HB1 珊瑚化石照片；(d)—剖面 1 层泥晶灰岩显微照片(+)；(e)、(f)—剖面 2 层泥晶方解石(+)、石英砂岩岩屑(+)；(g)—剖面 3 层泥晶方解石(+)；(h)、(i)—剖面 4 层泥晶方解石(+)、团块石英碎屑(+)；(j)、(k)—剖面 5 层腕足(+)、头足(+)；(l)、(m)—剖面 6 层泥晶方解石(+)、石英砂屑岩屑(+)；(n)、(o)、(p)、(q)—剖面 7 层头足(+)、菊石(+)、生物骨架(+)；(r)—剖面 8 层生物碎屑(+)。

(a)and(b)—Photos of in the filed the geological point D001-HB1 and coral fossil；(c)—Photos of the geological point D003-HB1 coral fossil；(d)—micrograph of micrite first layer of profile (orthogonal polarization)；(e)、(f)—micrograph of micritic calcite and quartz sandstone debris second layer profile (orthogonal polarization)；(g)—micrograph of micritic calcite third layer of profile (orthogonal polarization)；(h)、(i)—micrograph of micritic calcite and briquettes quartz clastic fourth layer of profile (orthogonal polarization)；(j)、(k)—micrograph of brachiopoda and cephaopod fifth layer of profile (orthogonal polarization)；(l)、(m)—micrograph of micritic calcite and quartz sandstone debris sixth layer profile (orthogonal polarization)；(n)、(o)、(p)、(q)—micrograph of cephalopod, ammonite and organic framework seventh layer of profile (orthogonal polarization)；(r)—micrograph of bioclasts eighth layer profile (orthogonal polarization)

表 1 西藏羌塘蓝岭地区已鉴定珊瑚化石的属种和时代
Table 1 The classification of genus/specie,age of the identified coral fossils of Lanling area in central Qiangtang of Tibet

样品编号	珊瑚属种	时代	国内外产地及层位
D1529-HB2	<i>Conophyllia</i> sp.	T ₂ -T ₃	奥地利 Norian 阶的 Dachstein 组; 希腊 Pantokrator 的上三叠统 Carnian 阶; 贵阳青岩中三叠统青岩组
D1529-HB13	<i>Omphalopyllia</i> sp.	T ₂ -T ₃	西藏改则原展金组; 意大利北部上三叠统
D1529-HB15	<i>Coryphyllia</i> sp.	T ₃ -J ₁	希腊、塔吉克斯坦和斯洛文尼亚上三叠统
D1529-HB26	<i>Pamirastraea</i> sp.	T ₃	四川德格县上三叠统图姆沟组
D001-HB1	<i>Margarophyllia</i> cf. <i>capitata</i>	T ₃	四川白玉县上三叠统曲嘎寺组
D003-HB1	<i>Craspedophyllia</i> sp.	T ₂ -T ₃	奥地利、斯洛文尼亚三叠系

5b),结合泥晶灰岩镜下特征(图 6d、6g),根据泥晶灰岩含有机质含量多少,推测第 3 层泥晶灰岩沉积水体比第 1 层水深,指示一种低能闭塞环境,推测为局限台地亚相。第 2、4 层为灰—深灰色中薄层—薄层状泥晶灰岩与灰绿色薄层状粉砂质泥岩互层,泥晶灰岩中偶夹有浅灰色的硅质岩薄层(图 5b);与第 1、3 层相比,灰岩层厚减小,并且粉砂质、泥质成分大量出现,代表水深增加的相对低能环境;另外,波状层理构造则反映沉积环境处于浪基面以上(图 5b),海浪活动较为频繁。结合镜下陆源碎屑物特征和泥晶灰岩(图 6f、i),因此,综合判断第 2、4 层代表开阔台地相。第 5 层深灰色厚层状生屑、内碎屑灰岩(图 5d),生物碎屑和内碎屑以砾屑为主,含少量砂屑,部分具有明显的磨圆,其中化石破碎明显,整体形态显示异地埋藏特点,反映海水活动剧烈,并且生物种类繁多,含有珊瑚、腕足类、头足类、海百合茎等化石(图 6j、k、6n~6p),显示海水流通性好。这些特征指示第 5 层沉积环境为开阔台地。第 6 层为深灰色中薄层—薄层状泥晶灰岩与灰绿色粉砂质泥岩互层,泥晶灰岩中多含有硅质岩薄层,其中粉砂质泥岩层厚明显(图 5e),镜下可见泥晶方解石和石英砂屑岩屑(图 6l、m),判断其代表半局限半开阔台地相沉积。第 7 层岩性和生物组合特征与第 5 层相似,但层厚及总体厚度明显增加,生物种类更多,大量出现了具有鉴定意义的单体、复体珊瑚化石(图 7(4b))以及窄适性生物头足类、腕足类(图 6n~6p),沉积环境为台地边缘礁。第 8 层与第 1、3 层岩性类似,指示开阔台地相。

总体上,该套地层从下到上显示 2 个沉积旋回,

不仅体现为岩性组合的大致重复,也体现为水深和能量的韵律性变化,具体沉积环境为:局限台地—开阔台地—半局限半开阔台地—边缘礁—开阔台地(图 8)。

4 讨论

4.1 羌塘中部中一晚三叠世地层划分对比

结合区调资料和前人研究成果(Chen Wenxi et al.,2009;Ji Zhansheng et al.,2010;Li Cai et al.,2007)^{①②④⑤},对比羌塘中部中一晚三叠统岩石地层、时代和沉积相等特征,发现增生杂岩带与南、北羌塘地块中一晚三叠世的沉积作用存在明显分带(表 2)。

表 2 西藏羌塘盆地中一晚三叠统地层划分对比表

Table 2 Stratigraphic division and correlation of the Middle-Late Triassic in Qiangtang basin, Tibet

作者		北羌塘地块	中央增生杂岩带	南羌塘地块
地层年代单位		陈文西等 (2009)	①②	④
下一中侏罗统(J ₁₋₂)		雁石坪群		色哇组
上三叠统 (T ₃)	瑞替阶	那底岗日组	望湖岭组	日干配错组
	诺利阶			
	卡尼阶			
中三叠统 (T ₂)	拉丁阶	肖茶卡组	硅质岩 放射虫 碳酸盐岩	
	安尼阶	康南组	增生杂岩(中奥陶世—晚三叠世)	

羌塘中央增生杂岩带内,中一晚三叠统呈零星状分布,呈断块状与围岩相接触(Deng Wanming et al.,1996;Li Yuejun et al.,1997;Li cai et al.,2007;Ji Zhansheng et al.,2010),岩石组合类型、沉积相差异较大,从深海盆地相硅质岩到浅海碳酸盐岩、细碎屑岩均有出露。才玛尔错两侧、雪水河西岸原定为早二叠世地层中均发现了中一晚三叠统放射虫硅质岩(Deng Wanming et al.,1996;Li Yuejun et al.,1997);才玛尔错东发现有上三叠统砂屑灰岩,含 *Thecosmilia* aff. *clathrata* Emmrich (格子剑鞘珊瑚相似种)(Xiong Xingguo et al.,2006);本次从蓝岭北侧下一中二叠统解体出中一晚三叠统碳酸盐岩。果干加年山出露的上三叠统望湖岭组角度不整合于增生杂岩之上,底部流纹岩夹层锆石 U-Pb 年龄 214±4 Ma 及下伏杂岩阳起石 Ar-Ar 坪年龄

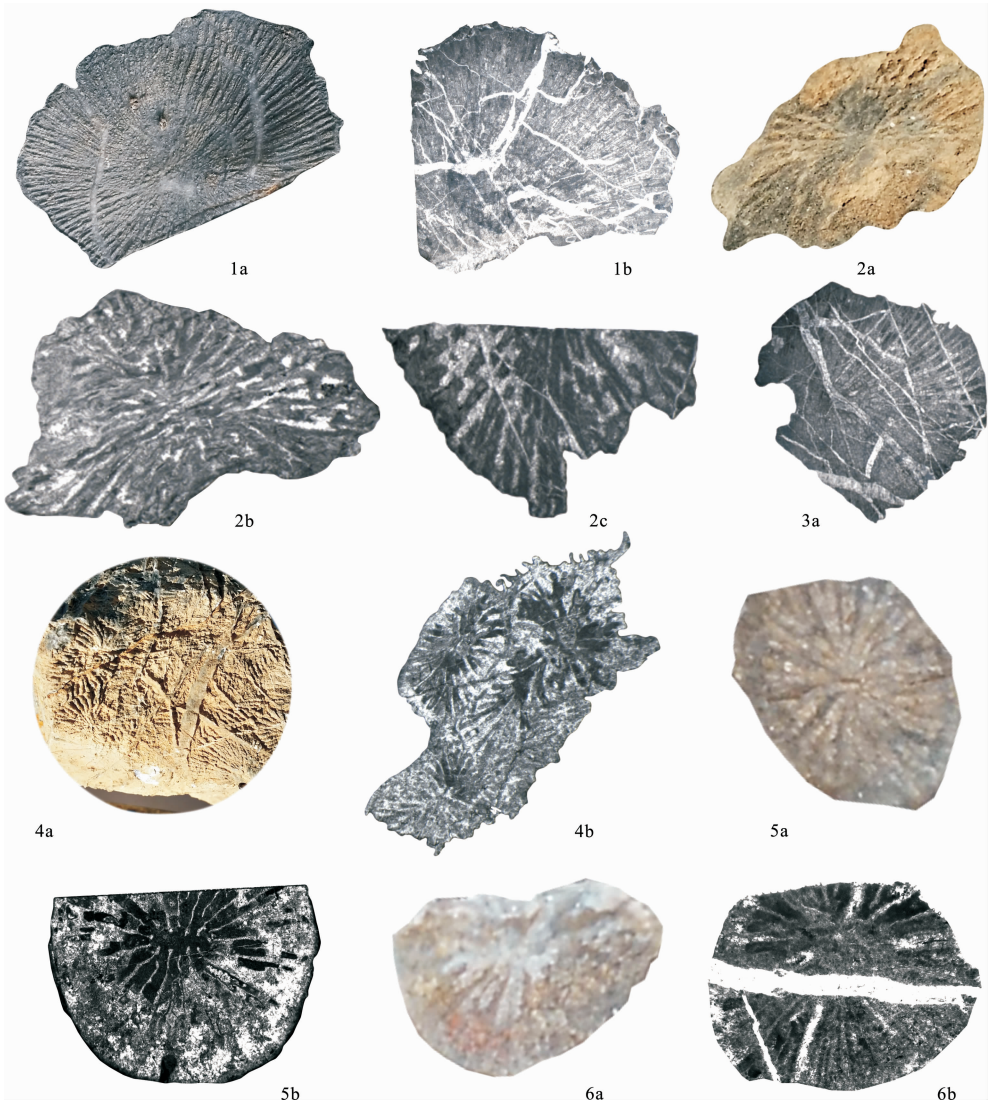


图 7 西藏羌塘蓝岭地区已鉴定珊瑚化石的标本、显微尺度照片特征

化石说明格式:图片中编号,化石拉丁文名称,野外样品编号,室内编号,产出层位,放大倍数

Fig. 7 The mesoscopic and microscopic photograph assembly of the identified coral fossils of the Middle-Late Triassic carbonatite from the three locations of Lanling area in central Qiangtang of Tibet. The format of explanation of each coral:figure number,genus and species,the original number of the specimen,chamber number of specimen,magnification,bed
(1a)、(1b)—*Conophyllia* sp. 野外编号 1529-HB2,室内编号 1529-HB2-a,横切面×5, 7 层;(2a)、(2b)、(2c)—*Omphalophyllia* sp. 野外编号 1529-HB13,室内编号 1529-HB13-a,室内编号 1529-HB13-b,横切面×2, 7 层;(3a)—*Coryphyllia* sp. 室内编号 1529-HB15-a,横切面×5, 7 层;(4a)、(4b)—*Pamirastraea* sp. 野外编号 1529-HB26,室内编号 1529-HB26-a,横切面×2, 5 层;(5a)、(5b)—*Margarophyllia* cf. *capitata* 野外编号 D001-HB1,室内编号 D001-HB1-a,横截面×2;(6a)、(6b)—*Craspedophyllia* sp. 野外编号 D003-HB1,室内编号 D003-HB1-a,横截面×2

(1a)and(1b)—*Conophyllia* sp. 1529-HB2,1529-HB2-a,×5,7 层;(2a)、(2b)、(2c)—*Omphalophyllia* sp. 1529-HB13,1529-HB13-a,×2, 7 层;(3a)—*Coryphyllia* sp. 1529-HB15-a, × 5, 7 层;(4a)、(4b)—*Pamirastraea* sp. 1529-HB26, 1529-HB26-a, × 2, 5 层;(5a)、(5b)—*Margarophyllia* cf. *capitata*,D001-HB1,D001-HB1-a,×2;(6a)、(6b)—*Craspedophyllia* sp. D003-HB1,D003-HB1-a,×2

219.7±6.5Ma 共同指示角度不整合时代在诺利中期,代表洋壳俯冲结束后的碰撞期沉积(Li Cai et al.,2007)。Ji Zhansheng et al.(2010)在望湖岭组碎屑岩所夹的灰岩岩块中采集到了晚三叠世卡尼期—诺利期珊瑚化石组合和诺利期牙形石。

北羌塘三叠纪地层时代连续并且出露完整,菊花山—沃若山—江爱达日那一带的中上三叠统肖茶卡组(菊花山组)整合于中三叠统安尼阶康南组之上,主要岩性为一套碳酸盐岩,并见含煤碎屑岩,化石组合时代为拉丁期—诺利期,被上三叠统那底岗

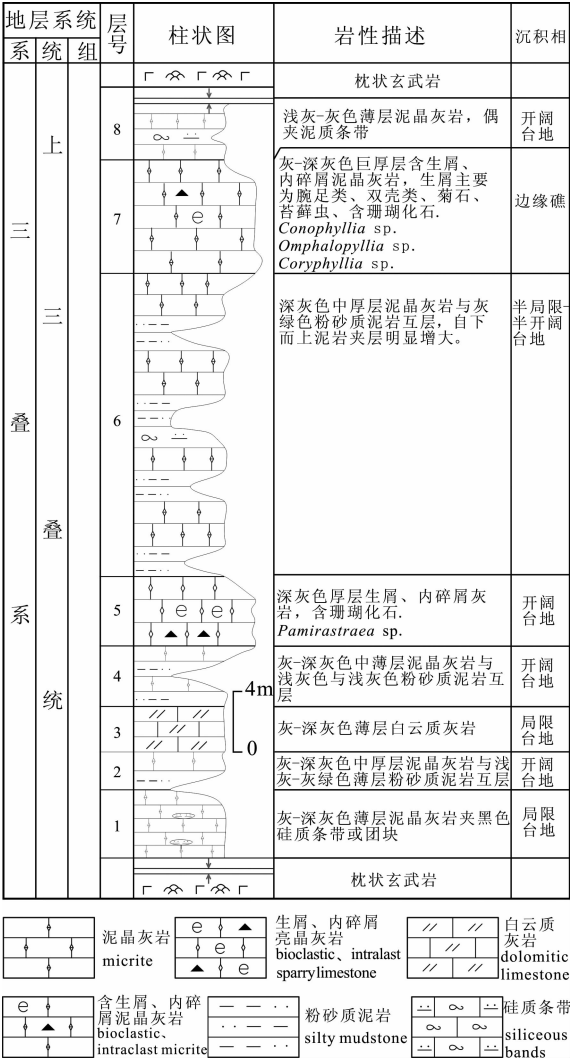


图 8 西藏羌塘中部蓝岭北侧上三叠统地层柱状图
Fig. 8 Stratigraphic column of the Upper Triassic on the north flank of Lanling, Central Qiangtang of Tibet

日组不整合覆盖^②(Chen Wenxi et al. ,2009)。那底岗日组中酸性火山岩呈 NWW-SEE 向断续状出露于北羌塘地块南缘,锆石 U-Pb 年龄集中在 223~202 Ma,上部与下一中侏罗统雁石坪群底部砂砾岩平行不整合接触。

南羌塘地体,上三叠统日干配错组普遍以角度不整合覆盖于上古生界^③(Chen Wenxi et al. ,2009),珊瑚等生物化石组合时代为卡尼期—瑞替期^④(Chen Wenxi et al. ,2009),下段见基性火山岩,中段为浅海碳酸盐台地沉积,上段为陆棚相(钙质)粉砂岩、泥岩夹岩屑长石砂岩透镜体(Chen Wenxi et al. ,2009)。

4.2 龙木错—双湖缝合带三叠纪古地理格局

对羌塘中部高压变质带的研究表明:龙木错—双湖缝合带为古特提斯洋向北羌塘地块之下俯冲消

减的产物(Li Cai et al. ,2008;Zhai Qingguo et al. ,2011a; Liang Xiao et al. ,2012),榴辉岩峰期变质时间为 237~230 Ma(Zhai Qingguo et al. ,2011b),指示俯冲启动的时间早于中三叠世。晚三叠世,沿龙木错—双湖缝合带涌现了剧烈的中酸性岩浆事件,带内沿桃形湖—蜈蚣山—果干加年山—冈塘错—双湖近东西向展布有中酸性—酸性的花岗岩株或流纹质火山岩,侵入于增生杂岩之中或覆盖其上,时代集中在 222~209 Ma(Kapp et al. ,2003;Hu Peiyuan et al. ,2010;Li Guangming et al. ,2015);北羌塘的南缘,沿菊花山—那底岗日—各拉丹东出露那底岗日组中酸性火山岩(Fu Xiugen et al. ,2010b;Zhang Kaijun et al. ,2011;Zhai Qingguo et al. ,2013),锆石 U-Pb 年龄区间为 223~202 Ma,但普遍大于 208 Ma。它们的岩石地球化学特征指示同碰撞背景(Li Guangming et al. ,2015;Li Jingchao et al. ,2015),极有可能为古特提斯洋俯冲结束后深部发生板片断裂(slab break-off)的结果(Zhang Kaijun et al. ,2011)。尽管目前仍未在龙木错—双湖缝合带发现可靠的三叠纪弧火山岩组合,但该套中—酸性岩浆岩的时空格局明确指示古特提斯洋具有向北的俯冲极性,羌塘中部三叠纪古地理格局可能类似于南美安第斯山型洋陆俯冲造山带(图 9)。弧火山岩的赋存状态可对比雅鲁藏布江缝合带,即在拉萨地块的南缘大面积分布同碰撞型冈底斯花岗岩基,而火山弧岩岩浆遭受了强烈的抬升侵蚀,只在岩体边缘零星出露(Yin An et al. ,2000;Searle et al. ,2011;Pan Guitang et al. ,2012a)。

研究表明:活动大陆边缘的沉积作用具有明显的分带性,弧前增生系统包含三类沉积体系,从海沟到岩浆弧依次为海沟盆地、海沟—斜坡盆地与弧前盆地(Miall,1984;Dickson et al. ,1979;Liu Chiyang et al. ,2015)。蓝岭北侧的这套中—晚三叠世台地相碳酸盐岩产出于龙木错—双湖缝合带内部,以断块状上叠于下一中二叠统蛇绿混杂岩,露头中产状十分低缓,呈正常的单斜岩层,与围岩的构造样式不协调,表明其产出的构造背景为洋壳俯冲带增生楔。岩层变质程度微弱,显示并未卷入俯冲作用,很可能受浅表层次逆冲断层改造呈断块状。其次,珊瑚化石组合显示的中—晚三叠世沉积作用发生时洋壳并未关闭(那底岗日、冈塘错晚三叠世中酸性岩浆事件、望湖岭组角度不整合反映洋壳闭合时间为晚三叠世诺利期)。再者,羌塘中部典型的台地相碳酸盐岩产出在南、北羌塘地块的边缘,分别为晚三叠世日

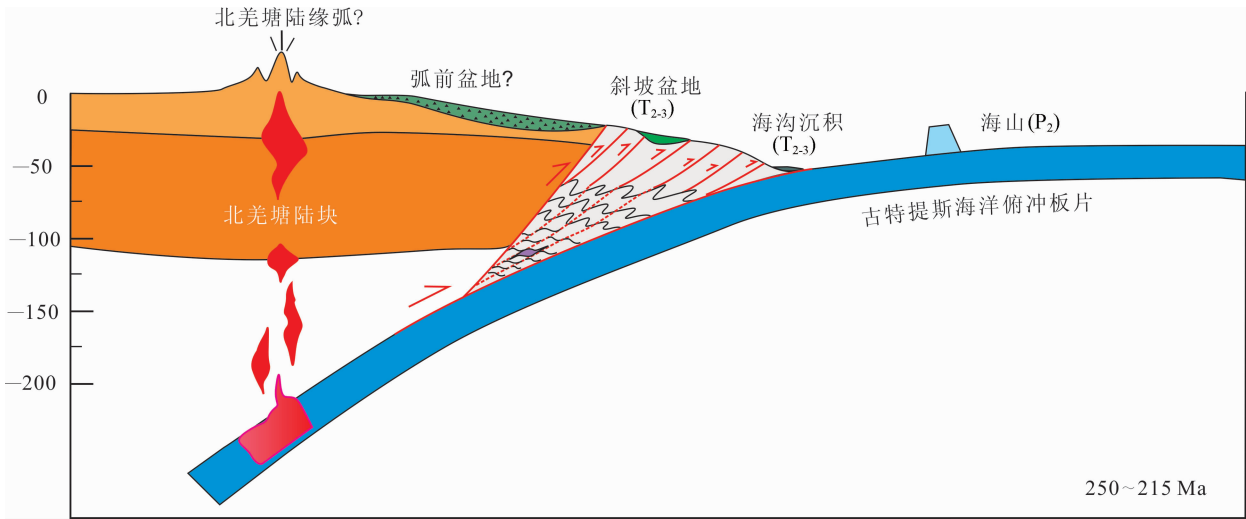


图9 西藏羌塘中部中一晚三叠世斜坡盆地模型及其大地构造位置

Fig. 9 The tectonic setting of Middle-Late Triassic slope basin in Central Qiangtang of Tibet

干配错组、中一晚三叠世菊花山组或肖茶卡组，它们的时代、岩性组合、与围岩接触关系，明显区别于蓝岭北侧的中—上三叠统。因此，笔者推断蓝岭中—上三叠统代表北美塘活动大陆边缘的弧前沉积，具体类型为岩浆弧前增生楔上叠盆地。鉴于其浅水沉积环境，进一步推测盆地紧邻弧前盆地发育(图9)。

4.3 增生杂岩时代与古特提斯洋闭合时限

玛依岗日一角木日一带的古生界为印支期古特提斯洋向北西俯冲消减形成的增生杂岩(Wang Genhou et al., 2009; Liang Xiao et al., 2012)。杂岩中地层时代呈单向收敛的无限区间，上限为洋壳的最终闭合期，下限并不确定，即不同时代、不同大地构造相的岩层沿构造面相互拼贴混杂(Wang Genhou et al., 2009)。这种时代的混乱或多元性从上世纪90年代以来羌塘中部地质调查的历史可见一斑，1:100万改则幅^⑥厘定阿木岗群(戈木日群)为前泥盆系，Wu Ruizhong et al., (1986)将其解体为变质相和岩性存在差别的三个亚群：前泥盆系、前志留系、前寒武系。油气地质会战中众多学者认为其为羌塘盆地的前泥盆系结晶基底，并认为下部为中深变质的结晶基底，上部为浅变质褶皱基底(Wang Chengshan et al., 1987; Zhao Zhengzhang et al., 2001; Wang Jian et al., 2004)，两者之间为角度不整合接触(Huang Jijun et al., 2001)。Li Cai et al., (1990、1995、2001)将其划分为前震旦系与上石炭统，上石炭统一二叠系。Li Yuejun et al., (1997)在荣玛乡北部雪水河的戈木日群硅质岩中发现了典型的中一晚三叠世放射虫组合。Deng Wanming et al., (1996)也在雪水河与才玛尔错东南硅质岩中发

现了中晚三叠世放射虫。增生杂岩主体为上古生界，含丰富的早—中二叠世蜓类动物群^⑦，基本对应日土地区的冈瓦纳冷水相霍尔巴错群和吞龙共巴组(Liang Dingyi et al., 1983)，解体出的中奥陶统一泥盆系属于冈瓦纳大陆北缘陆表海沉积^⑧。1:5万角木日区调^⑨从中解体出一套台地相的中二叠统碳酸盐岩—龙格岩组。Geng Quanru et al., (2012)在玛依岗日西侧的果干加年山大沙河，发现了晚石炭世珊瑚，证实了展金组的存在。笔者在蓝岭北侧新发现了中—上三叠统碳酸盐岩。

总之，很难对增生杂岩的原岩时代予以统一约束，随着地质研究程度的深入，构造拼贴而成的无序岩系(混杂岩)面临着时代的解体。因此，厘定增生杂岩时代区间将有助于约束大洋的演化时间，依据卷入俯冲的最新地层可大致限定洋盆的关闭时间。日湾茶卡地区天泉山洋岛火山岩的发现指示羌塘中部晚二叠世—中三叠世仍存在成熟洋盆，反映洋壳消减完毕的时间在中三叠世以后(Fan Jianjun et al. 2016, 2017)。蓝岭高压变质带东侧新发现的中—上三叠统呈断块状混杂于下一中二叠统增生杂岩之中，与蓝岭西侧中一晚三叠世放射虫硅质岩时代基本一致，指示龙木错—双湖古特提斯洋的关闭应在晚三叠世诺利期以后。果干加年山地区，上三叠统望湖岭组陆棚相细碎屑岩角度不整合在增生杂岩之上，代表洋盆闭合后的残留海沉积，底部流纹岩获得的锆石 SHRIMP U-Pb 谐和年龄：214±4 Ma 限制了闭合的最新时限(Li Cai et al., 2007)。Ji Zhansheng et al., (2010)在望湖岭组灰岩块中发现了晚三叠世卡尼期—诺利期珊瑚化石。

另外,龙木错—双湖缝合带南北两侧同碰撞型岩浆事件十分活跃,同样反映洋盆的关闭时间为诺利期中期(Cui Yuliang et al.,2017)。冈塘错花岗岩株侵入于增生杂岩之中,时代集中在 210 Ma 左右^{①②}(Kapp et al.,2003;Li Jingchao et al.,2015)。北羌塘南缘,NWW 向展布的那底岗日组中酸性火山岩时代集中在 217~208Ma(Fu Xiugen et al.,2010a,2010b;Wang Jian et al.,2007,2008;Zhang Kaijun et al.,2011;Zhai Qingguo et al.,2007;Zhai Qingguo et al.,2013)。

5 结论

(1)羌塘中部蓝岭北侧新发现一套中一晚三叠世台地相碳酸盐岩夹细碎屑岩、硅质岩沉积,呈断块状上叠于中二叠世蛇绿混杂岩,顶底不全,被脆性断层所截切。岩性组合、相对层厚和化石保存状态的差异指示三个沉积亚相:局限台地相、开阔台地相和半局限—半开阔台地相和台地边缘礁相。

(2)砾屑(生屑)灰岩中含丰富的特提斯温暖浅水型珊瑚化石组合,以及腕足类、头足类、菊石、苔藓虫、有孔虫等。共鉴定出 6 个属种(含未定种)的珊瑚化石,分别为 *Conophyllia* sp.,*Omphalophyllia* sp.,*Coryphyllia* sp.,*Pamirastraea* sp.,*Margarophyllia* cf. *capitata*.,*Craspedophyllia* sp.。反映地层时代为中三叠世拉丁阶一晚三叠世诺利阶。

(3)羌塘中央增生杂岩带和两侧的南、北羌塘地体在中晚三叠世分别发育不同的沉积体系。综合构造和地层特征,推测蓝岭北侧的中上三叠统代表北羌塘活动大陆边缘增生楔上叠盆地,古特提斯洋在晚三叠世诺利期中期发生闭合。

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注 释

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Discovery of the Middle-Late Triassic Carbonatite Formation in the Central Qiangtang Accretionary Complex, North Tibet and its Tectonic Significance

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Abstract

The disintegration of the strata within the Central Qiangtang accretionary complex (CQAC) provides important role in determining the temporal-spatial evolution of the Longmu Co-Shuanghu suture and the Paleo-Tethys Ocean. For a long time, however, the study of the Triassic strata in this area remains poor, particularly in lacking a correlation with the existing knowledge about the Triassic ocean-continent subduction/collision tectonics. Three outcrops on the north of Lanling were found to host one set of a Middle-Late Triassic platform-facies carbonatite intercalated with fine-grained clastic rocks and silicalite. They occur as brittle fault-bounded blocks and are superimposed on the Middle Permian ophiolitic mélangé, with an approximately horizontal and un-inverted bedding, but its top and bottom are tectonically removed away. Based on the lithological assemblages, thickness variation of limestone layers and fossil distribution, three sedimentary subfacies were determined: restricted platform, open platform and shoal in the platform margin. The shoal facies bioclastic/intraclast limestones contain abundant hexacoralia fossils consisting of six genera including one species such as *Conophyllia* sp., *Omphalophyllia* sp., *Coryphyllia* sp., *Pamirastraea* sp., *Margarophyllia* cf. *capitata*, and *Craspedophyllia* sp., which suggest an age of Ladinian of Middle Triassic-Norian of Late Triassic. A comparative analysis on the Middle-Upper Triassic strata of the Qiangtang basin shows different sedimentation systems occurring in North Qiangtang terrane, central accretionary complex (CQAC), and South Qiangtang terrane. Both structural relationship and stratigraphic characteristics indicate that the newly discovered Middle-Upper Triassic carbonatite characterizes a portion of the fore-arc deposition in the active continental margin in south of North Qiangtang terrane, that is, slope basin immediately above the Triassic accretion wedge, which therefore implies a northward subduction polarity for the Paleo-Tethys Ocean. On the basis of tectonic background of the Middle-Upper Triassic in Central Qiangtang and time of the fossil assemblages, the closure of the Paleo-Tethys ocean was hence speculated to happen in middle Norian of Late Triassic.

Key words: Qiangtang; Triassic; Carbonatite; Coral; Paleo-Tethys Ocean