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柴达木盆地北缘都兰地区榴辉岩中 透长石+石英包裹体:超高压变质作用的证据

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内容提要 近年来在祁连地体与柴达木地块之间大柴旦—都兰北一线确定出一条榴辉岩带。在此带东端都兰北部地区野马滩一带榴辉岩中,我们发现产在石榴子石内部由透长石和多晶石英(柯石英假像)构成的包裹体。包裹体的内部结构特征及周围的放射状裂纹反映早期超高压变质相特征,据此,推测柴达木盆地北缘榴辉岩带可能是一形成深度超过80 km的超高压变质带。

关键词 榴辉岩 透长石+多晶石英包裹体 超高压变质 柴达木盆地北缘 都兰

自Chopin (1984)和Smith (1984)分别在变沉积岩和榴辉岩中发现柯石英包裹体之后,岩石中出现超高压特征矿物柯石英被认为是鉴别该岩石是否经历超高压变质作用最直接的标志。然而,在陆内碰撞带深部形成的榴辉岩折返地表的过程中,复杂和多阶段退化变质及水的参与往往使大部分的柯石英难以保存,一些时代较老的榴辉岩带尤其如此。实验表明,只要有极少量水的参与,柯石英很快就会变成多晶石英而以柯石英假像出现(Mosenfelder et al., 1997),因此,在论证岩石是否经历过超高压变质作用时,柯石英假像显得与柯石英同等重要(Liou et al., 1998)。例如,Enami等(1990)和Schmadicke(1991)分别利用柯石英假像论证了苏鲁地区和Saxonian Erzgebirge地区的榴辉岩是超高压变质的产物。

在榴辉岩及其相关岩石中,还有其他矿物及矿物组合能够反映岩石经历了超高压变质作用(Schreyer, 1995; Liou et al., 1998)。透长石和柯石英组合最早被发现于南非金伯利岩的榴辉岩基质以及金刚石的包裹体中,反映透长石可以形成于超高压的地幔环境(Prinz et al., 1975; Smyth et al., 1977);另外,透长石+多晶石英组合还出现在含金刚石的超高压变质地体榴辉岩的石榴子石和绿辉石中(Okay, 1993; Zhang et al., 1994; Yang J-J et al., 1998; Massonne, 1993; Massonne et al.,

2000),表明该组合也应该代表超高压变质岩石的标志性矿物。

近几年,在祁连地体与柴达木地块之间的大柴旦—都兰北约350 km长的变质带内发现了大量的榴辉岩(杨经绥等, 1998, 2000; 张雪亭等, 1999; Zhang et al., 2000a),榴辉岩的锆石U-Pb年龄为 494.6 ± 6.5 Ma(张建新等, 2000),从而确定了柴达木地块与祁连地体之间存在一条早古生代以高压变质带为特征的板块俯冲带(许志琴等, 1999; Yang et al., 2000a, 2000b)。矿物成分温压计算表明大柴旦地区榴辉岩的形成压力可达2.8~3.3 GPa (Yang et al., 2000b),属出现柯石英的温压稳定区,而作为榴辉岩的围岩副片麻岩锆石中柯石英包裹体的发现(杨经绥等, 2001),充分证明了本区是超高压变质地体,但是榴辉岩中至今尚未找到超高压标志矿物——柯石英。本文报道了都兰地区榴辉岩中发现的透长石+多晶石英(柯石英假像),也可以像柯石英和金刚石一样作为超高压变质作用的最直接证据。

1 榴辉岩产出的地质背景

都兰地区榴辉岩产于都兰县城东北约40 km的野马滩一带,祁连地体与柴达木地块的交界处,属ESE向延伸的柴达木盆地北缘榴辉岩带的东段,也是该榴辉岩带中3个主要的榴辉岩出露地区之一(图1)。榴辉岩呈大小不等的岩块分布于新元古代花岗

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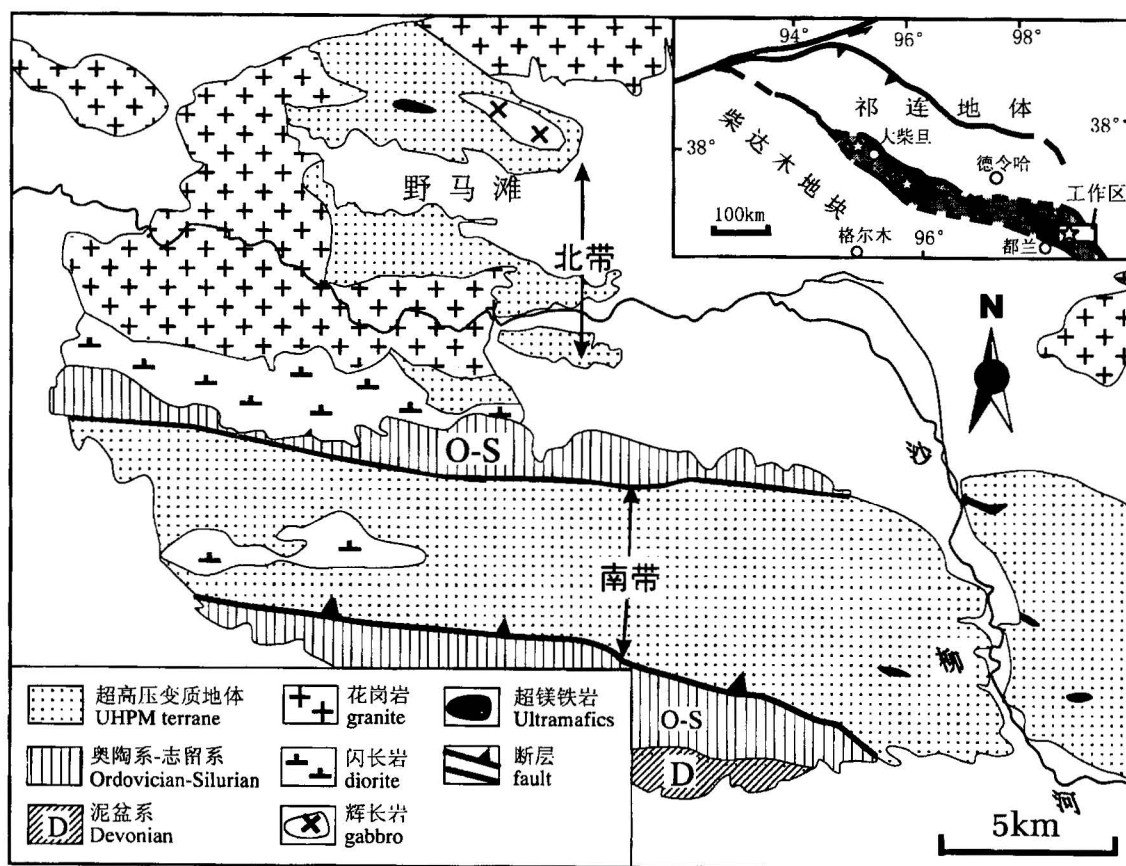


图1 都兰地区北东野马滩—沙柳河一带地质简图及榴辉岩位置(据1:20万都兰幅地质图修改)

Fig. 1 Geological map and locations of eclogite in Yematan—Shaliuhe region, northeastern Dulan, Qinghai Province (after 1:200000 geological map of Dulan, revised)

质片麻岩中,最大的岩块达200 m×300 m,并与一些蛇纹石化橄榄岩岩块及石榴角闪岩岩块相伴产出。

2 榴辉岩的岩石和矿物特征

榴辉岩岩块呈致密坚硬的透镜体突出于花岗质片麻岩中。研究表明,都兰地区榴辉岩可以被划分为差异十分明显的南北两个亚带(杨经绥等,2000)(图1)。南带榴辉岩的峰期变质矿物组合为石榴子石+绿辉石+金红石+蓝晶石+石英,矿物的定向性不明显,绿辉石中 Na_2O 含量相对较低(3%~4%);榴辉岩的退化变质较强,出现大量普通角闪石、斜长石、黝帘石交代早期变质矿物。北带榴辉岩峰期变质矿物组合为石榴子石+绿辉石+金红石+石英(柯石英)+多硅白云母,未见有蓝晶石和斜长石,绿辉石的 Na_2O 含量相对较高,可达6%以上;其中部分岩石可见明显的片理化以及石榴子石和绿辉石被拉长定向的变形结构。后期退化变质的矿物主要为角

闪石和单斜辉石+钠质斜长石后生合晶。

3 透长石+多晶石英(柯石英假像)

被发现的透长石(KAlSi_3O_8)+多晶石英集合体主要呈矿物包裹体产于北带榴辉岩的石榴子石中,包裹体的形态一种为六边形结构,其内部为栅栏状(palyside-shaped)多晶石英构成的柯石英假像,并呈放射状向包裹体的周缘发散,多晶透长石则构成了包裹体的基质(图2a和b);另一种为卵圆形的多晶透长石+多晶石英包裹体,其中石英也是呈栅栏状集中于包裹体的外部边缘(图2c和d)。电子探针分析显示透长石的Or为94%~99%(表1),而激光拉曼光谱分析显示很好的透长石谱线,与苏鲁超高压榴辉岩中透长石包裹体(Yang J-J et al., 1998)一致。其中915~917 cm^{-1} 是石榴子石的峰位(图3)。激光拉曼光谱测试在台湾成功大学实验室进行,型号: DILOR-JOBIN YVON U1000;激光功率1.5W,波长514.54nm。包裹体内透长石与石英的相对含量

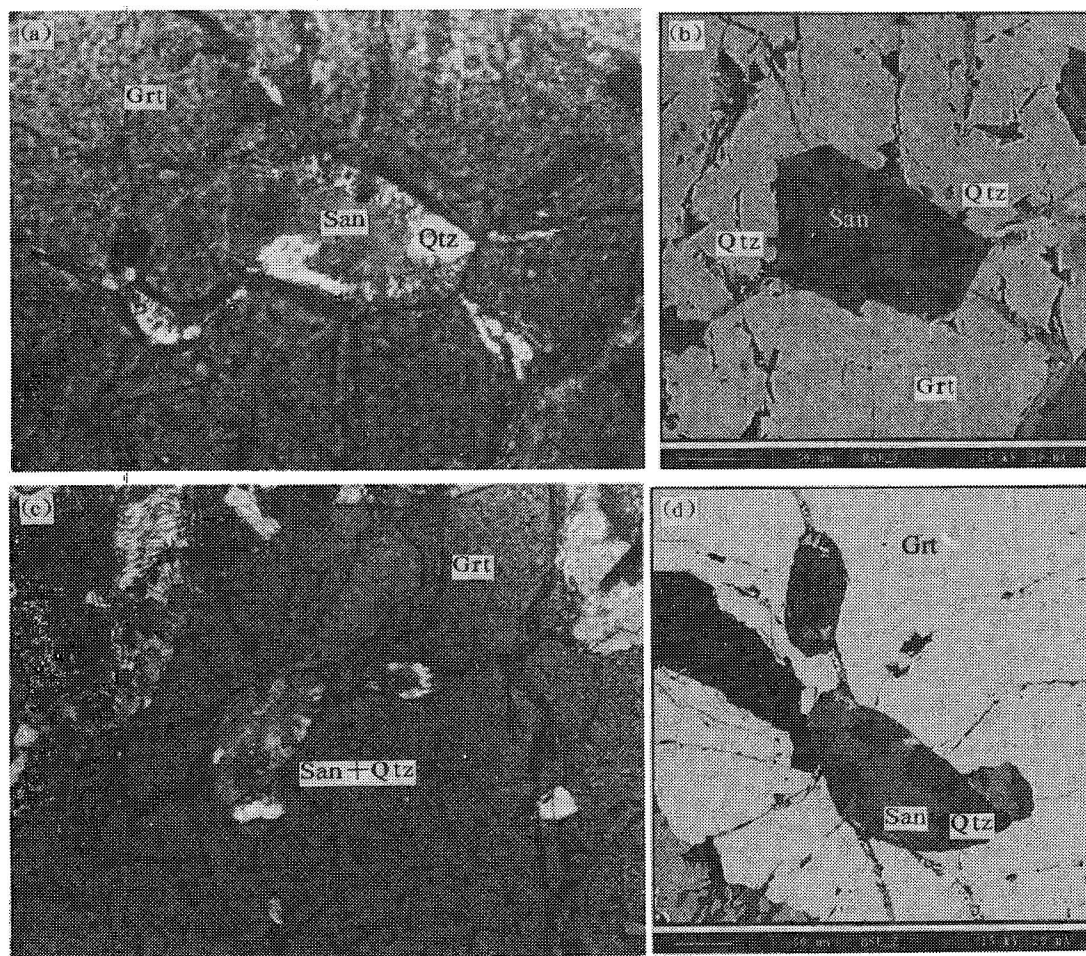


图 2 都兰北带榴辉岩多晶透长石+多晶石英包裹体

Fig. 2 Sanidine with polycrystalline quartz inclusions in eclogitic garnets from the north eclogite belt in Dulan

(a)—石榴子石中六边形透长石+石英包裹体,放射状裂纹发育,石英分布于包裹体的边缘呈栅栏状;(b)—(a)图中包裹体的背散射图像;

(c)—石榴子石中卵圆形透长石+石英(栅栏状)包裹体,放射状裂纹发育;(d)—(c)图中包裹体的背散射图像

(a)—Hexagonal sanidine + quartz inclusion in garnet with fine radial cracks, quartzes are palisade-like on the rim of the inclusion;

(b)—back-scattered image of inclusion (a); (c)—oval-shaped sanidine + quartz (palisade-like) in garnet with radial fractures;

(d)—back-scattered image of inclusion (c)

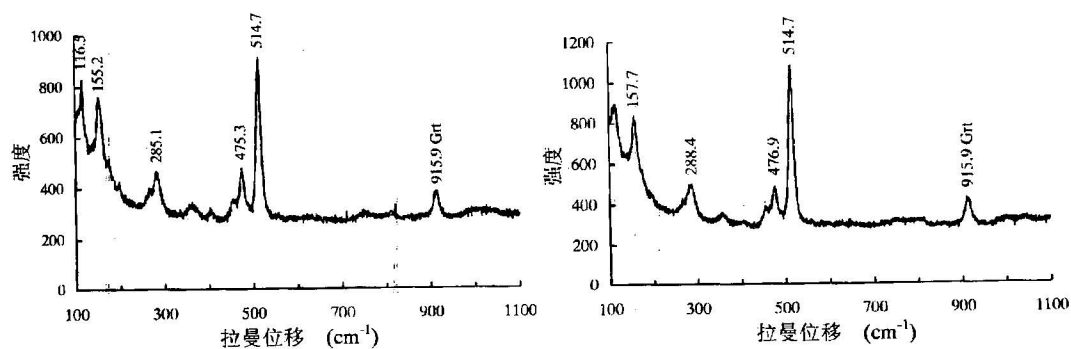


图 3 榴辉岩中透长石包裹体的激光拉曼光谱

Fig. 3 Laser Raman spectrum of sanidine inclusions in garnet

变化较大,有的完全由多晶透长石组成,有的则完全为多晶石英,其寄主矿物石榴子石发育极好的放射状裂纹,这些特征都显示包裹体可能是超高压相矿物组合透长石+柯石英在上升过程中退化变质并发生体积膨胀的结果(Enami et al., 1990; Okay, 1993; Yang J-J et al., 1998)。

表1 石榴子石中透长石和石英包裹体的成分

Table 1 Composition of sanidine and quartz inclusions in garnets

矿物	透长石					石英	
样号	99Y308	Y125-1	Y125-2	Y125-3	99Y134	99Y308	Y125
SiO ₂	65.49	64.94	64.59	63.89	64.16	99.60	100.16
TiO ₂	0.00	0.00	0.03	0.00	0.00	0.00	0.00
Al ₂ O ₃	18.07	18.43	18.28	17.99	18.25	0.17	0.01
Cr ₂ O ₃	0.00	0.04	0.02	0.00	0.00	0.00	0.02
MgO	0.01	0.00	0.00	0.01	0.00	0.00	0.01
CaO	0.01	0.01	0.00	0.01	0.10	0.00	0.02
MnO	0.08	0.00	0.04	0.00	0.00	0.03	0.04
FeO	0.14	0.27	0.11	0.12	0.11	0.21	0.16
BaO	0.24	0.11	0.16			0.01	0.00
Na ₂ O	0.51	0.37	0.05	0.12	0.13	0.01	0.02
K ₂ O	15.62	16.34	16.55	17.59	16.75	0.10	0.01
total	100.17	100.50	99.83	99.73	99.50	100.11	100.45
Or	94.74	96.44	99.20	98.93	98.35		
Ab	4.74	3.31	0.50	1.03	1.16		
An	0.07	0.05	0.00	0.05	0.49		
Cn	0.45	0.20	0.30				

4 讨论与结论

实验岩石学表明,虽然透长石并不是超高压的标志矿物,但其稳定范围可以达到5 GPa(Lindesey, 1966; Kinomura et al., 1975; Yagi et al., 1994),远远超过石英—柯石英的转变线(Bohlen et al., 1982)。金伯利岩中榴辉岩基质中透长石+柯石英共生组合(Smyth et al., 1977)及透长石在金刚石中呈包裹体的出现(Prinz et al., 1975)等证明了实验得出的结论。透长石(有的文献中统称为钾长石)±石英包裹体在不同造山带的超高压榴辉岩中已有发现,例如,Zhang等(1994)认为苏鲁超高压地体的钾长石是早期超高压相矿物 K-cymrite (KAlSi₃O₈·H₂O,稳定范围大于3 GPa)分解的产物;Yang J-J等(1998)则将相同现象解释为超高压环境形成的高钾矿物(钾质硬玉)在降压过程中分解为 KAlSi₃O₈+SiO₂的结果,而其原始矿物稳定的压力范围超过4.0 GPa;Massonne (1993)在研究德国 Erzgebirge 和 Munchberger 超高压地体榴辉岩时曾认为钾长石+石英是由矿物中熔体包裹体重结晶而成,但在后来

的研究中他们根据包裹体的形态和结构否定了硅酸盐熔体成因的假说,而新的认识认为该组合是形成于超高压(UHP)条件下的固体矿物相(Massonne et al., 2000)。

因此,尽管对钾长石+多晶石英在榴辉岩石榴子石和绿辉石中出现的原因有不同解释,但根据其特征的结构和形态,以及以独特的组合产在典型超高压榴辉岩地体之中等事实,以上研究者都一致认为钾长石(或透长石)±石英(柯石英假像)是榴辉岩峰期超高压(压力大于3 GPa)变质作用的产物(Smith et al., 1991; Okay, 1993; Zhang et al., 1994; Yang J S et al., 1998; Massonne et al., 2000)。对于柴达木盆地北缘都兰榴辉岩中出现的透长石+多晶石英(柯石英假像),我们认为可以将其作为榴辉岩早期经历超高压变质过程的标志,尽管如何解释它们的形成目前尚未定论。

产于大陆内部造山带中的高压超高压榴辉岩带被认为代表了一种陆陆碰撞作用的产物(Maruyama et al., 1996),研究这些高压超高压变质作用发生及发展,尤其它们的深部地质过程及其折返机制,是研究大陆增生和认识大陆内部动力事件的关键(Coleman et al., 1995)。都兰地区榴辉岩中透长石+石英聚晶(柯石英假像)包裹体的发现说明该区榴辉岩峰期变质的压力已达到2.8 GPa以上(俯冲深度达80~90 km),从而证明柴达木盆地北缘榴辉岩带应属超高压变质作用的产物。

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Sanidine + Quartz Inclusions in Dulan Eclogites: Evidence for UHP Metamorphism on the North Margin of the Qaidam Basin, NW China

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Abstract

An Early Palaeozoic eclogite-bearing complex belt has been identified as a continental collision suture zone between the Qilian terrain and the Qaidam terrain in recent years. K-feldspar (sanidine) occurs alongside quartz (a pseudomorph after coesite) as microinclusions in garnets of eclogite from the north belt eclogite in Dulan, the east part of the north Qaidam eclogite belt. The mosaic polycrystalline texture of K-feldspar + quartz aggregation and radial fractures around the inclusions indicate that they are the product of peak metamorphism under the ultrahigh-pressure condition. Therefore, the authors consider that the eclogite-bearing complex belt is most probably a UHP belt formed at a depth over 80 km.

Key words: eclogite; sanidine + polycrystalline inclusion; ultrahigh-pressure metamorphism; north margin of the Qaidam basin; Dulan
